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Helicopter-borne magnetic,
electromagnetic and radiometric
geophysical survey in Harstad area,
Troms County.



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Summary:

NGU conducted an airborne geophysical survey in Harstad area in July 2019. This report describes and documents the acquisition, processing and visualization of the recorded datasets. The geophysical survey results consist of approximately 1700 line-km of data, covering an area of 340 km².

The NGU modified Geotech Ltd. Hummingbird frequency domain EM system supplemented by an optically pumped Cesium magnetometer and the Radiation Solutions 1024 channels RSX-5 spectrometer mounted on an AS350-B3 helicopter was used for data acquisition.

The survey was flown with 200 m line spacing, 90° direction. Average speed was 55 km/h, and average height clearance of the EM bird was 63 m, and 93 meters for the spectrometer.

Collected data were processed at NGU using Geosoft Oasis Montaj software. Raw total magnetic field data were corrected for diurnal variation and leveled using standard micro levelling algorithm. Radiometric data were processed using standard procedures recommended by International Atomic Energy Association (IAEA).

EM data were filtered and leveled using both automated and manual levelling procedures. Apparent resistivity was calculated from in-phase and quadrature data for three coplanar frequencies (880Hz, 6.6kHz and 34kHz), and for two coaxial frequencies (980Hz and 7kHz) separately using a homogeneous half space model and gridded using proxy resistivity values.

All data were gridded using cell size of 50x50 meters and presented as 40% transparent grids with shaded relief on top of topographic maps, at the scale of 1:100.000.

Keywords:	Airborne	Geophysics
Magnetic	Gamma spectrometry	Radiometric
Electromagnetic	Technical report	

CONTENTS

1. SURVEY SPECIFICATIONS	7
1.1 Airborne Survey Parameters	7
1.2 Airborne Survey Instrumentation.....	8
1.3 Airborne Survey Logistics Summary	9
2. DATA PROCESSING AND PRESENTATION	9
2.1 Total Field Magnetic Data	9
2.2 Electromagnetic Data	11
2.3 Radiometric data.....	13
3. PRODUCTS	17
4. REFERENCES	17
Appendix A1: Flow chart of magnetic processing.....	18
Appendix A2: Flow chart of EM processing	18
Appendix A3: Flow chart of radiometry processing.....	18

FIGURES

Figure 1: Harstad survey area.....	6
Figure 2: Hummingbird system in air	8
Figure 3: Gamma-ray spectrum with K, Th, U and Total Count windows.....	13
Figure 4: Harstad survey area with flight path	20
Figure 5: Total Magnetic Field	21
Figure 6: Magnetic Horizontal Gradient	22
Figure 7: Magnetic Vertical Gradient	23
Figure 8: Magnetic Tilt Derivative	24
Figure 9: Apparent Resistivity. Frequency 7000 Hz, Coaxial coils	25
Figure 10: Apparent Resistivity. Frequency 6600 Hz, Coplanar coils.....	26
Figure 11: Apparent Resistivity. Frequency 980 Hz, Coaxial coils	27
Figure 12: Apparent Resistivity. Frequency 880 Hz, Coplanar coils.....	28
Figure 13: Apparent Resistivity. Frequency 34133 Hz, Coplanar coils.....	29
Figure 14: Radiometric Total counts.....	30
Figure 15: Potassium ground concentration	31
Figure 16: Uranium ground concentration	32
Figure 17: Thorium ground concentration	33
Figure 18: Radiometric Ternary Image	34

TABLES

Table 1. Instrument Specifications	8
Table 2. Hummingbird EM system, frequency and coil configurations	8
Table 3. Survey Specifications Summary	9
Table 4. Apparent Resistivity and related Proxy Resistivity Values	12
Table 5. Specified channel windows for the 1024 RSX-5 system.....	13
Table 6. Maps in scale 1:100.000, available from NGU on request.....	17

INTRODUCTION

In 2019 NGU received government funds to acquire airborne geophysical data in the Harstad area on Hinnøya island in Troms County. The helicopter survey reported herein amounts to 1700 line-km (340 km²), with the area covered shown in Figure 1.

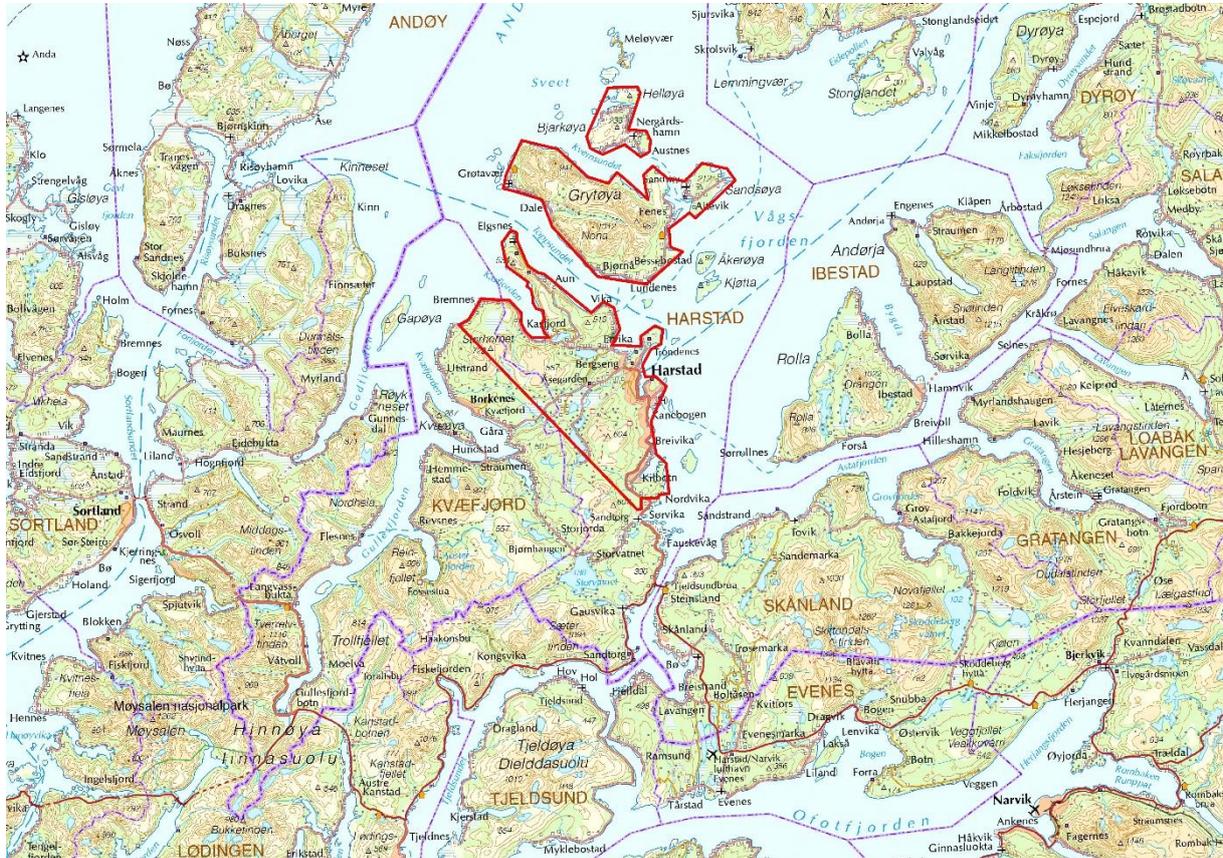


Figure 1: Harstad survey area

The objective of the airborne geophysical survey was to obtain a dense high-resolution magnetic, electromagnetic and radiometric data set over the survey area. This data is required for the enhancement of a general understanding of the regional geology of the area, with adjoining areas covered by other airborne surveys earlier years.

In this regard, the data can be used to map contacts and structural features within the property. It also improves defining the potential of known zones of mineralization, their geological settings, and identifying new areas of interest, as the dataset fills a gap in the previously acquired high-resolution geophysical surveys of the region.

The survey incorporated the use of a Hummingbird™ 5-frequency electromagnetic system supplemented by a high-sensitivity cesium magnetometer, gamma-ray spectrometer and radar altimeter. A GPS navigation computer system with flight path indicators ensured accurate positioning of the geophysical data with respect to the World Geodetic System 1984 geodetic datum (WGS-84).

1. SURVEY SPECIFICATIONS

1.1 Airborne Survey Parameters

NGU used a modified Hummingbird™ electromagnetic and magnetic helicopter survey system designed to obtain low level, slow speed, detailed airborne magnetic and electromagnetic data (Geotech 1997). The system was supplemented by 1024 channel gamma-ray spectrometer, installed under the belly of the helicopter, which was used to map ground concentrations of U, Th and K, and radiation Total Counts.

The airborne survey began on July 2nd, 2019 and ended on July 13th, 2019. A Eurocopter AS350-B3 (LN-OGL) from helicopter company HeliScan AS was used to tow the bird. The survey lines were spaced 200 m apart, with lines oriented at 90°. The magnetic and electromagnetic sensors are housed in a single 7.5 m long bird, flown at an average height of about 63 m above the topographic surface.

Rugged terrain and abrupt changes in topography affected the aircraft pilot's ability to 'drape' the terrain; therefore the average instrumental height was sometimes higher than the standard survey instrumental height, which is defined as 40 m plus a height of obstacles (trees, power lines etc.) for EM and magnetic sensors.

The ground speed of the aircraft varied from 40 to 110 km/h depending on topography, wind direction and its magnitude. On average the ground speed during measurements is calculated to 55 km/h. Magnetic data were recorded at 0.2 second intervals resulting in approximately 3 m average point spacing.

EM data were recorded at 0.1 second intervals resulting in data with a sample increment of 1.5 m along the ground in average. Spectrometry data were recorded every 1 second giving a point spacing of approximately 15 meters. The parameters allow recognizing details in the data to detect subtle anomalies that may represent mineralization and/or rocks of different lithological and petrophysical composition.

A base magnetometer to monitor diurnal variations in the magnetic field was located at the base in Sollia ski resort, near Harstad, UTM33 7632000N 556700E. The GEM GSM-19 station magnetometer data were recorded once every 3 seconds. The CPU clock of the base magnetometer and the helicopter magnetometer were both synchronized to UTC (Universal Time Coordinates) through the built-in GPS receiver to allow correction of diurnals.

Navigation system uses GPS/GLONASS satellite tracking systems to provide real-time WGS-84 coordinate locations for every second. The accuracy achieved with no differential corrections is reported to be ± 5 m in the horizontal directions. The GPS receiver antenna was mounted internally inside the canopy of the helicopter.

For quality control, the electromagnetic, magnetic and radiometric, altitude and navigation data were monitored on four separate windows in the operator's display during flight while they were recorded in three data ASCII streams to the PC hard disk drive. Spectrometry data were also recorded to an internal hard drive of the spectrometer. The data files were transferred to the field workstation via USB flash drive. The raw data files were backed up onto USB flash drive in the field.

1.2 Airborne Survey Instrumentation

Instrument specification is given in Table 1. Frequencies and coil configuration for the Hummingbird EM system is given in Table 2.

Table 1. Instrument Specifications

Instrument	Producer/Model	Accuracy / Sensitivity	Sampling frequency / interval
Magnetometer	Scintrex Cs-2	<2.5nT throughout range / 0.0006nT $\sqrt{\text{Hz}}$ rms	5 Hz
Base magnetometer	GEM GSM-19	0.1 nT	3 s
Electromagnetic	Geotech Hummingbird	1 – 2 ppm	10 Hz
Gamma spectrometer	Radiation Solutions RSX-5	1024 ch's, 16 liters down, 4 liters up	1 Hz
Radar altimeter	Bendix/King KRA 405B	$\pm 3\%$ 0 – 500 feet $\pm 5\%$ 500 – 2500 feet	1 Hz
Pressure/temperature	Honeywell PPT	$\pm 0.03\%$ FS	1 Hz
Navigation	Topcon GPS-receiver	± 5 meter	1 Hz
Acquisition system	NGU custom software		

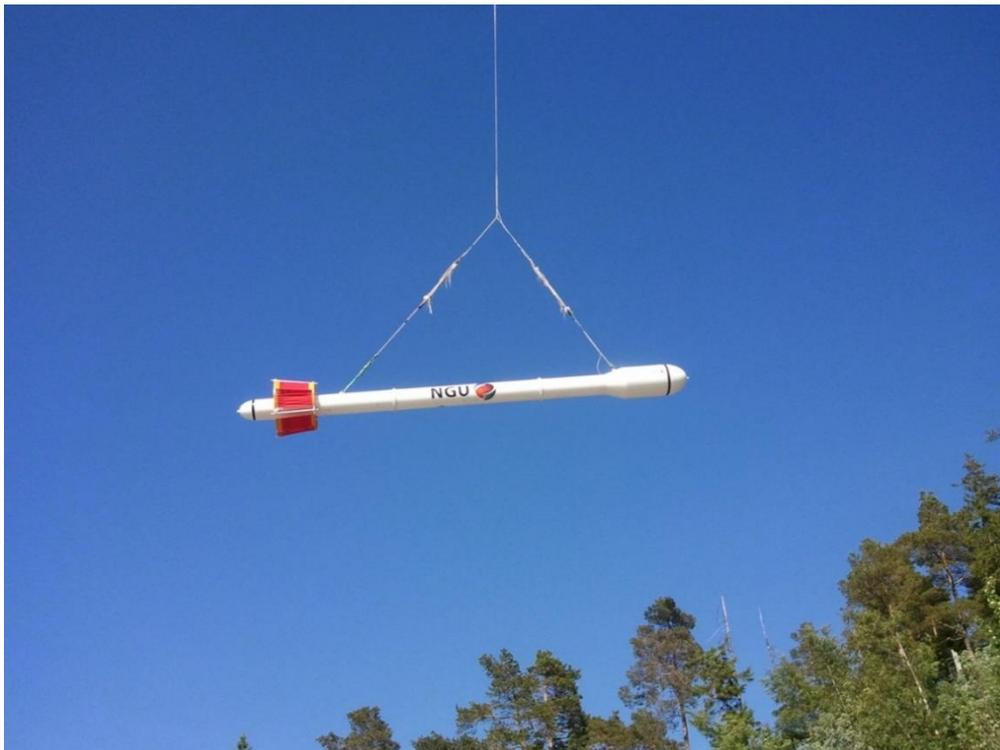


Figure 2: Hummingbird system in air

Table 2. Hummingbird EM system, frequency and coil configurations

Coils	Frequency	Orientation	Separation
A	7701 Hz	Coaxial	6.30 m
B	6606 Hz	Coplanar	6.30 m
C	980 Hz	Coaxial	6.025 m
D	880 Hz	Coplanar	6.025 m
E	34133 Hz	Coplanar	4.90 m

1.3 Airborne Survey Logistics Summary

A summary of the survey specifications is shown in Table 3.

Table 3. Survey Specifications Summary

Parameter	Specifications
Traverse (survey) line spacing	200 meters
Traverse line direction	E-W (90°)
Nominal aircraft ground speed	40 - 110 km/h
Average aircraft ground speed	55 km/h
Average sensor terrain clearance Mag	63 m
Average sensor terrain clearance Rad	93 m
Sampling rates:	
Magnetometer	0.2 seconds
EM	0.1 seconds
Spectrometer, GPS, altimeter	1.0 second
Base Magnetometer	3.0 seconds

2. DATA PROCESSING AND PRESENTATION

The magnetic, spectrometry and resistivity data were processed by Frode Ofstad at NGU. The ASCII data files were loaded into three separate Oasis Montaj databases. All three datasets were processed consequently according to processing flow charts shown in Appendix A1, A2 and A3.

2.1 Total Field Magnetic Data

At the first stage the raw magnetic data was visually inspected, and spikes were removed manually. Non-linear filter was also applied to airborne raw data to eliminate short-period spikes. Typically, several corrections must be applied to magnetic data before gridding - heading correction, lag correction and diurnal correction.

Diurnal Corrections

The temporal fluctuations in the magnetic field of the earth affect the total magnetic field readings recorded during the airborne survey. This is commonly referred to as the magnetic diurnal variation. These fluctuations can be effectively removed from the airborne magnetic dataset by using a stationary reference magnetometer that records the magnetic field of the earth simultaneously with the airborne sensor at given short time interval.

Diurnal variation channel was inspected for spikes, and spikes were removed manually if necessary. Magnetic diurnals that were recorded on the base station magnetometer were within the standard NGU specifications during the entire survey (Rønning 2013).

Diurnal variations were measured with GEM GSM-19 magnetometer. The base station computer clock was continuously synchronized with GPS clock. The average datum base level (\bar{B}_B) was set to 53535 nT for the Harstad area. The recorded data are merged with the airborne data and the diurnal correction is applied according to equation (1).

$$\mathbf{B}_{Tc} = \mathbf{B}_T + (\bar{B}_B - \mathbf{B}_B) \quad (1)$$

where:

$$\begin{aligned} \mathbf{B}_{Tc} &= \text{Corrected airborne total field readings} \\ \mathbf{B}_T &= \text{Airborne total field readings} \\ \bar{B}_B &= \text{Average datum base level} \\ \mathbf{B}_B &= \text{Base station readings} \end{aligned}$$

Corrections for Lag and heading

Neither a lag nor cloverleaf tests were performed before the survey. According to previous reports the lag between logged magnetic data and the corresponding navigational data was 1-2 fids. Translated to a distance it would be no more than 10 m - the value comparable with the precision of GPS. A heading error for a towed system is usually either very small or non-existent. No lag and heading corrections were applied.

Magnetic data processing, gridding and presentation

The total field magnetic anomaly data (\mathbf{B}_{TA}) were calculated from the diurnal corrected data (\mathbf{B}_{Tc}) after subtracting the IGRF for the surveyed area calculated for the data period (eq.2)

$$\mathbf{B}_{TA} = \mathbf{B}_{Tc} - IGRF \quad (2)$$

IGRF 2015 model was employed in these calculations, to ensure that the 2019 Harstad data would better match the previously processed and published surrounding data set from Hinnøya area, flown in 2013.

The total field anomaly data were split into lines and then were gridded using a minimum curvature method with a grid cell size of 50 meters. This cell size is exactly one quarter of the 200m average line spacing. In order to remove small line-to-line levelling errors that were detected on the gridded magnetic anomaly data, the Geosoft Micro-levelling technique was applied on the flight line based magnetic database. Then, the micro-leveled channel was gridded using minimum curvature method with 50 m grid cell size.

The processing steps of magnetic data presented so far, were performed on point basis. The following steps are performed on grid basis.

The Horizontal and Vertical Gradient along with the Tilt Derivative of the total magnetic anomaly were calculated from the stitched micro-leveled total magnetic anomaly grid. The magnitude of the horizontal gradient was calculated according to equation (3)

$$HG = \sqrt{\left(\frac{\partial(B_{TA})}{\partial x}\right)^2 + \left(\frac{\partial(B_{TA})}{\partial y}\right)^2} \quad (3)$$

where \mathbf{B}_{TA} is the micro-leveled total field anomaly field. The vertical gradient (VG) was calculated by applying a vertical derivative convolution filter to the micro-leveled \mathbf{B}_{TA} field. The Tilt derivative (TD) was calculated according to the equation (4)

$$TD = \tan^{-1}\left(\frac{VG}{HG}\right) \quad (4)$$

A 3x3 convolution filter was applied to smooth the resulted magnetic grids. The results are presented in a series of colored shaded relief maps (1:100.000). The maps are:

- A. Total field magnetic anomaly
- B. Horizontal gradient of total magnetic anomaly
- C. Vertical gradient of total magnetic anomaly
- D. Tilt Derivative (or Tilt angle) of the total magnetic anomaly

These maps are representative of the distribution of magnetization over the surveyed areas. The list of the produced maps is shown in Table 6.

2.2 Electromagnetic Data

The EM system transmits five fixed frequencies and records an in-phase and a quadrature response for each of the five coil sets of the electromagnetic system. The received signals are processed and used for computation of an apparent resistivity.

In-phase and quadrature data were filtered with 15 fiducial non-linear filter to eliminate spherical spikes, which were represented as irregular noise of large amplitude in records and high frequency noise of bird electronics. Then, a 20-fiducial low-pass filter was applied to suppress instrumental and cultural noise. These filters were not able to suppress all the noise completely, due to irregular nature of noise. Also, shifts of 7000 Hz IP and Q records, with amplitude of 5-10 ppm, was observed in some flights. Shifts were edited manually where possible.

In order to remove the effects of instrument drift caused by gradual temperature variations in the transmitting and receiving circuits, background responses are recorded during each flight. To obtain a background level, the bird is raised to an altitude of at least 1000 ft above the topographic surface so that no electromagnetic responses from the ground or seawater are present in the recorded traces.

The EM traces observed at this altitude correspond to a background (zero) level of the system. If these background levels are recorded at 20-30 minutes interval, then the drift of the system (assumed to be linear) can be removed from the data by resetting these points to the initial zero level of the system. The drift must be removed on a flight-by-flight basis, before any further processing is carried out. Geosoft HEM module was used for applying drift correction. Residual instrumental drift, usually small, but non-linear, was manually removed on a line-to-line basis.

When levelling of the EM data was complete, apparent resistivity was calculated from in-phase and quadrature EM components using a homogeneous half space model of the earth (Geosoft HEM module) for 6600, 7000, 980, 880 and 34133 Hz. A threshold value of 2 ppm was set for inversion, with a starting value of 500 ohm-m.

Electromagnetic field decays rapidly with the distance (height of the sensors) – as z^{-2} – z^{-5} depending on the shape of the conductors and, at certain height, signals from the ground sources become comparable with instrumental noise. Levelling errors or precision of levelling can lead sometimes to appearance of artificial resistivity anomalies when data were collected at high instrumental altitude.

For better visual appearance of the data in the gridding process, proxy resistivity values between 1 to 13 was used, according to the specified intervals given in Table 4 below.

Table 4. Apparent Resistivity and related Proxy Resistivity Values

Apparent Resistivity (Ω m)	Proxy Resistivity Value		Apparent Resistivity (Ω m)	Proxy Resistivity Value
< 3	1		200-500	8
3-5	2		500-1000	9
5-10	3		1000-2000	10
10-20	4		2000-5000	11
20-50	5		5000-10000	12
50-100	6		Over 10000	13
100-200	7			

Application of threshold allows excluding such data from an apparent resistivity calculation, though not completely. It's particularly noticeable in low frequencies datasets. Resistivity data were visually inspected; artificial anomalies associated with high altitude measurements were manually removed.

Data recorded at the height above 150 m were considered as non-reliable and removed from presentation. Proxy resistivity were gridded with a cell size 50 m. Power lines strongly affected low frequency data – 880 and 980 Hz channels, and the most prominent noise from power lines were filtered manually.

2.3 Radiometric data

Airborne gamma-ray spectrometry measures the abundance of Potassium (K), Thorium (eTh), and Uranium (eU) in rocks and weathered materials by detecting gamma-rays emitted due to the natural radioelement decay of these elements. The data analysis method is based on the IAEA recommended method for U, Th and K (International Atomic Energy Agency, 1991; 2003). A short description of the individual processing steps of that methodology as adopted by NGU is given below.

Energy windows

The Gamma-ray spectra were initially reduced into standard energy windows corresponding to the individual radio-nuclides K, U and Th. Figure 3 shows an example of a Gamma-ray spectrum and the corresponding energy windows and radioisotopes (with peak energy in MeV) responsible for the radiation.

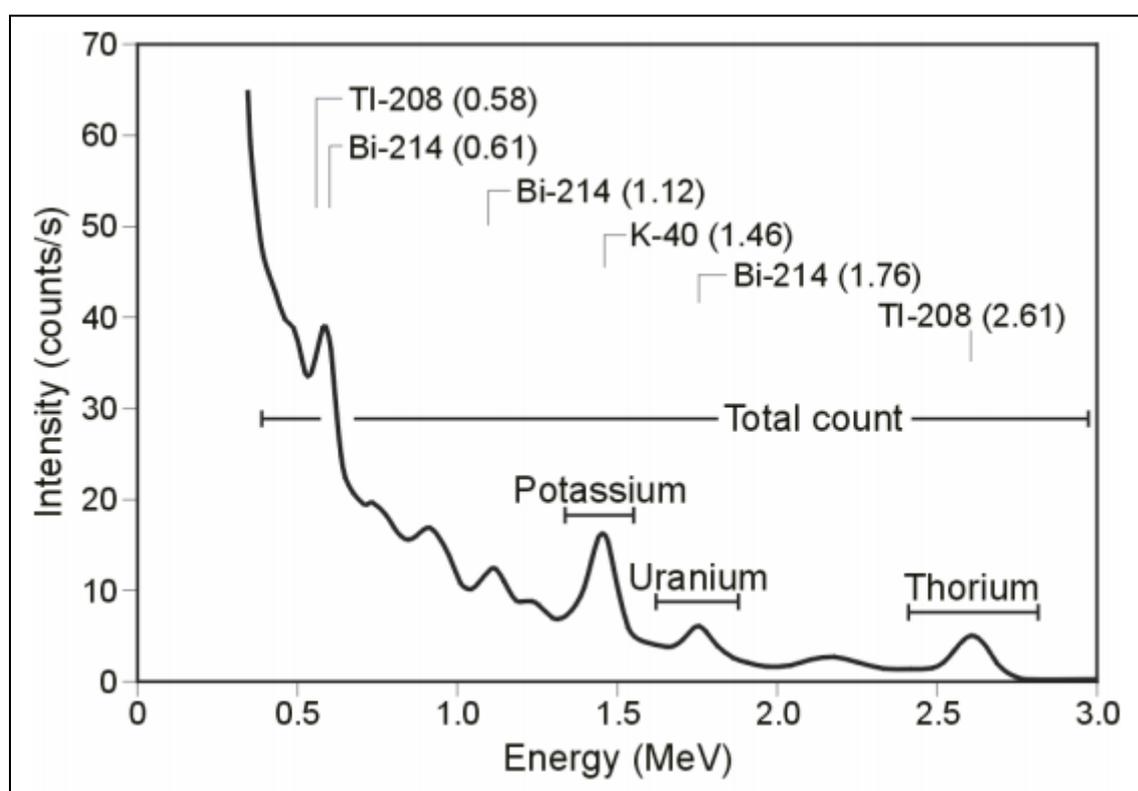


Figure 3: Gamma-ray spectrum with K, Th, U and Total Count windows.

Table 5. Specified channel windows for the 1024 RSX-5 system

Gamma-ray spectrum	Cosmic	Total count	K	U	Th
Down	1023	135-935	455-522	552-618	802-935
Up	1023			552-618	
Energy windows (MeV)	>3.07	0.41-2.81	1.37-1.57	1.66-1.86	2.41-2.81

The RSX-5 is a 1024 channel system with four downward and one upward looking detector, which means that the actual Gamma-ray spectrum is divided into 1024 channels. The first channel is reserved for the “Live Time” and the last for the Cosmic rays. Table 5 shows the channels that were used for the reduction of the spectrum.

Live Time correction

The data were corrected for live time. “Live time” is an expression of the relative length of time the instrument was able to register new pulses per sample interval. On the other hand, “dead time” is an expression of the relative length of time the system was unable to register new pulses per sample interval. The relation between “dead time” and “live time” is given by the equation (5)

$$\text{“Live time”} = \text{“Real time”} - \text{“Dead time”} \quad (5)$$

where the “real time” or “acquisition time” is the elapsed time over which the spectrum is accumulated (about 1 second).

The live time correction is applied to the total count, Potassium, Uranium, Thorium, upward Uranium and cosmic channels. The formula used to apply the correction is as follows:

$$C_{LT} = C_{RAW} \cdot \frac{\text{Acquisition Time}}{\text{Live Time}} \quad (6)$$

where C_{LT} is the live time corrected channel in counts per second, C_{RAW} is the raw channel data in counts per second, while Acquisition Time and Live Time are in microseconds.

Cosmic and aircraft correction

Background radiation resulting from cosmic rays and aircraft contamination was removed from the total count, Potassium, Uranium, Thorium, upward Uranium channels using the following formula:

$$C_{CA} = C_{LT} - (a_c + b_c \cdot C_{Cos}) \quad (7)$$

where C_{CA} is the cosmic and aircraft corrected channel, C_{LT} is the live time corrected channel a_c is the aircraft background for this channel, b_c is the cosmic stripping coefficient for this channel and C_{Cos} is the low pass filtered cosmic channel.

Radon correction

The upward detector method, as discussed in IAEA (1991), was applied to remove the effects of the atmospheric radon in the air below and around the helicopter. Using spectrometry data over-water, where there is no contribution from the ground sources, enables the calculation of the coefficients (a_c and b_c) for the linear equations that relate the cosmic corrected counts per second of Uranium channel with that of total count, Potassium, Thorium and Uranium upward channels over water. Data over-land was used in conjunction with data over-water to calculate the a_1 and a_2 coefficients used in equation (8) for the determination of the Radon component in the downward uranium window:

$$Radon_U = \frac{U_{up_{CA}} - a_1 \cdot U_{CA} - a_2 \cdot Th_{CA} + a_2 \cdot b_{Th} - b_U}{a_U - a_1 - a_2 \cdot a_{Th}} \quad (8)$$

where $Radon_U$ is the radon component in the downward Uranium window, $U_{up_{CA}}$ is the filtered upward uranium, U_{CA} is the filtered Uranium, Th_{CA} is the filtered Thorium, a_1 ,

a_2 , a_U and a_{Th} are proportional factors and b_U and b_{Th} are constants determined experimentally.

The effects of Radon in the downward Uranium are removed by simply subtracting $Radon_U$ from U_{CA} . The effects of radon in the other channels are removed using the following formula:

$$C_{RC} = C_{CA} - (a_C \cdot Radon_U + b_C) \quad (9)$$

where C_{RC} is the Radon corrected channel, C_{CA} is the cosmic and aircraft corrected channel, $Radon_U$ is the Radon component in the downward uranium window, a_C is the proportionality factor and b_C is the constant determined experimentally for this channel from over-water data.

Compton Stripping

Potassium, Uranium and Thorium Radon corrected channels, are subjected to spectral overlap correction. Compton scattered gamma rays in the radio-nuclides energy windows were corrected by window stripping using Compton stripping coefficients determined from measurements on calibrations pads (Grasty et al, 1991) at the Geological Survey of Norway in Trondheim (see values in Appendix A2).

The stripping corrections are given by the following formulas:

$$A_1 = 1 - (g \cdot \gamma) - (a \cdot \alpha) + (a \cdot g \cdot \beta) - (b \cdot \beta) + (b \cdot \alpha \cdot \gamma) \quad (10)$$

$$U_{ST} = \frac{Th_{RC} \cdot ((g \cdot \beta) - \alpha) + U_{RC} \cdot (1 - b \cdot \beta) + K_{RC} \cdot ((b \cdot \alpha) - g)}{A_1} \quad (11)$$

$$Th_{ST} = \frac{Th_{RC} \cdot (1 - (g \cdot \gamma)) + U_{RC} \cdot (b \cdot \gamma - a) + K_{RC} \cdot ((a \cdot g) - b)}{A_1} \quad (12)$$

$$K_{ST} = \frac{Th_{RC} \cdot ((\alpha \cdot \gamma) - \beta) + U_{RC} \cdot ((a \cdot \beta) - \gamma) + K_{RC} \cdot (1 - (a \cdot \alpha))}{A_1} \quad (13)$$

where U_{RC} , Th_{RC} , K_{RC} are the radon corrected Uranium, Thorium and Potassium and a , b , g , α , β , γ are Compton stripping coefficients. U_{ST} , Th_{ST} and K_{ST} are stripped values of U, Th and K.

Reduction to Standard Temperature and Pressure

The radar altimeter data were converted to effective height (H_{STP}) using the acquired temperature and pressure data, according to the expression:

$$H_{STP} = H \cdot \frac{273.15}{T + 273.15} \cdot \frac{P}{1013.25} \quad (14)$$

where H is the smoothed observed radar altitude in meters, T is the measured air temperature in degrees Celsius and P is the measured barometric pressure in millibars.

Height correction

Variations caused by changes in the aircraft altitude relative to the ground was corrected to a nominal height of 60 m. Data recorded at the height above 150 m were considered as non-reliable and removed from processing. Total count, Uranium,

Thorium and Potassium stripped channels were subjected to height correction according to the equation:

$$C_{60m} = C_{ST} \cdot e^{C_{ht} \cdot (60 - H_{STP})} \quad (15)$$

where C_{ST} is the stripped corrected channel, C_{ht} is the height attenuation factor for that channel and H_{STP} is the effective height.

Conversion to ground concentrations

Finally, corrected count rates were converted to effective ground element concentrations using calibration values derived from calibration pads (Grasty et al, 1991) at the Geological Survey of Norway in Trondheim (see values in Appendix A2). The corrected data provide an estimate of the apparent surface concentrations of Potassium, Uranium and Thorium (K, eU and eTh). Potassium concentration is expressed as a percentage, equivalent Uranium and Thorium as parts per million (ppm). Uranium and Thorium are described as “equivalent” since their presence is inferred from gamma-ray radiation from daughter elements (^{214}Bi for Uranium, ^{208}Tl for Thorium). The concentration of the elements is calculated according to the following expressions:

$$C_{CONC} = C_{60m} / C_{SENS_60m} \quad (16)$$

where C_{60m} is the height corrected channel, C_{SENS_60m} is experimentally determined sensitivity reduced to the nominal height (60m).

Spectrometry data gridding and presentation

Gamma-rays from Potassium, Thorium and Uranium emanate from the uppermost 30 to 40 centimeters of soil and rock in the crust (Minty, 1997). Variations in the concentrations of these radioactive elements are largely related to changes in the mineralogy and geochemistry of the Earth’s surface.

The spectrometry data were stored in a database and the ground concentrations were calculated following the processing steps. A list of the parameters used in these steps is given in Appendix A3.

Then the data were split in lines and ground concentrations of the three main natural radio-elements Potassium, Thorium and Uranium and total gamma-ray flux (total count) were gridded using a minimum curvature method with a grid cell size of 50 meters. In order to remove small line-to-line levelling errors appeared on those grids, the data were micro-leveled as in the case of the magnetic data, and re-gridded with the same grid cell size.

Quality of the radiometric data was within standard NGU specifications (Rønning 2013). For further reading regarding standard processing of airborne radiometric data, we recommend the publications from Minty et al. (1997).

A 3x3 convolution filter was applied to smooth the concentration grids. A list of the produced maps is shown on Table 6.

3. PRODUCTS

Processed digital data from the survey are presented as:

1. Geosoft XYZ files: Harstad_Mag.xyz, Harstad_EM.xyz, Harstad_Rad.xyz, Coloured maps at the scale 1:100.000 available from NGU on request.
2. Grid-files in Geotiff format

Table 6. Maps in scale 1:100.000, available from NGU on request.

Map #	Name
2020.009-00	Survey Flight Path
2020.009-01	Total magnetic field
2020.009-02	Magnetic Horizontal Gradient
2020.009-03	Magnetic Vertical Gradient
2020.009-04	Magnetic Tilt Derivative
2020.009-05	Apparent Resistivity, Frequency 7000 Hz, coaxial coils
2020.009-06	Apparent Resistivity, Frequency 6600 Hz, coplanar coils
2020.009-07	Apparent Resistivity, Frequency 980 Hz, coaxial coils
2020.009-08	Apparent Resistivity, Frequency 880 Hz, coplanar coils
2020.009-09	Apparent Resistivity, Frequency 34133 Hz, coplanar coils
2020.009-10	Radiometric Total Counts
2020.009-11	Potassium ground concentration
2020.009-12	Uranium ground concentration
2020.009-13	Thorium ground concentration
2020.009-14	Radiometric Ternary Map

Downscaled images of the maps are shown in figures 4 to 17.

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Appendix A1: Flow chart of magnetic processing

Meaning of parameters is described in the referenced literature.

Processing flow:

- Quality control.
- Visual inspection of airborne data and manual spike removal
- Merge basemag data with EM database
- Import of diurnal data
- Correction of data for diurnal variation
- IGRF removed
- Splitting flight data by lines
- Gridding
- Microlevelling
- 3x3 convolution filter

Appendix A2: Flow chart of EM processing

Meaning of parameters is described in the referenced literature.

Processing flow:

- Filtering of in-phase and quadrature channels with non-linear and low pass filters
- Automated leveling
- Quality control
- Visual inspection of data.
- Splitting flight data by lines
- Manual removal of remaining part of instrumental drift
- Calculation of an apparent resistivity using both - in-phase and quadrature channels
- Gridding

Appendix A3: Flow chart of radiometry processing

Underlined processing stages are not only applied to the K, U and Th window, but also to the total count. Meaning of parameters is described in the referenced literature.

- Airborne and cosmic correction (IAEA, 2003)
Used parameters: determined by high altitude calibration flights (1500-9000 ft) at Frosta in 2013

Channel	Background	Cosmic
K	7.3314	0.0617
U	0.8981	0.0454
Th	0.8881	0.0647
Uup	0.0423	0.0423
Total counts	36.291	1.0379

- Radon correction using upward detector method (IAEA, 2003)
Used parameters determined from survey data over water and land at Harstad in July 2019.

Coefficient	Value	Coefficient	Value
a_u	0.23881	b_u	0.01707
a_K	0.5335	b_K	2.98762
a_{Th}	0.08776	b_{Th}	0.28904
a_{TC}	16.26412	b_{TC}	10.88806
a_1	0.06722439	a_2	0.02561496

- Stripping corrections (IAEA, 2003)
Used parameters determined from measurements on calibrations pads at NGU on April 2015

Coefficient	Value
a	0.049355
b	0
c	0
α	0.305403
β	0.47203
γ	0.828857

- Height correction to a height of 60 m
Parameters determined by high altitude calibration flights (100 – 700 ft). The average values from tests performed at Frosta, 2013 and 2014 were used. Attenuation factors in 1/m:

Channel	Attenuation factor
K	-0.008884
U	-0.006528
Th	-0.006617
TC	-0.007331

- Converting counts at 60 m heights to element concentration on the ground
Used parameters determined from measurements on calibrations pads at NGU on April 2015

Channel	Sensitivity
K (%/count)	0.007544793
U (ppm/count)	0.088909372
Th (ppm/count)	0.151433049

- Microlevelling using Geosoft menu and smoothening by a convolution filtering

Microlevelling parameters	Value
De-corrugation cutoff wavelength (m)	1200
Cell size for gridding (m)	50
Naudy (1968) Filter length (m)	800

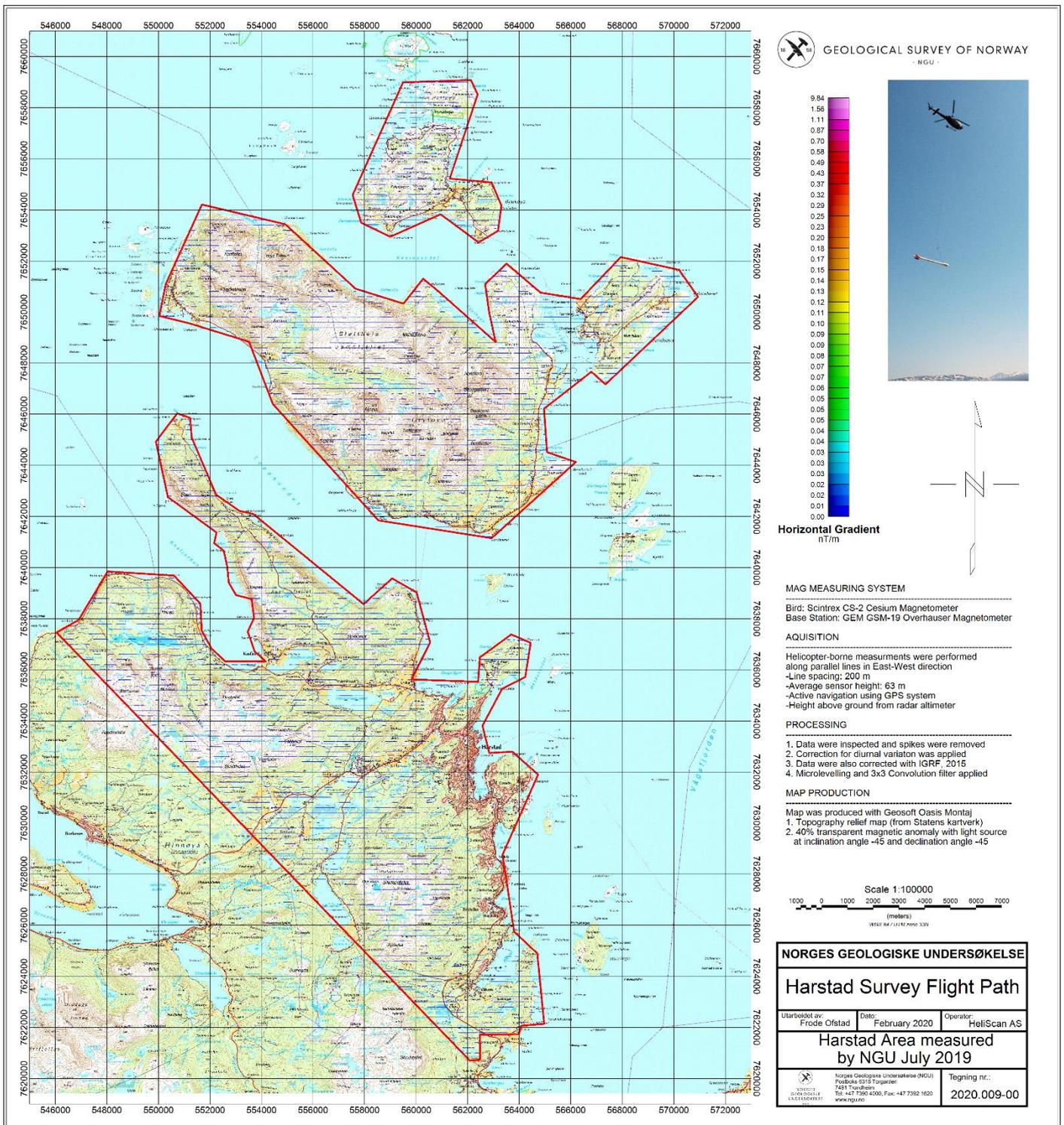


Figure 4: Harstad survey area with flight path

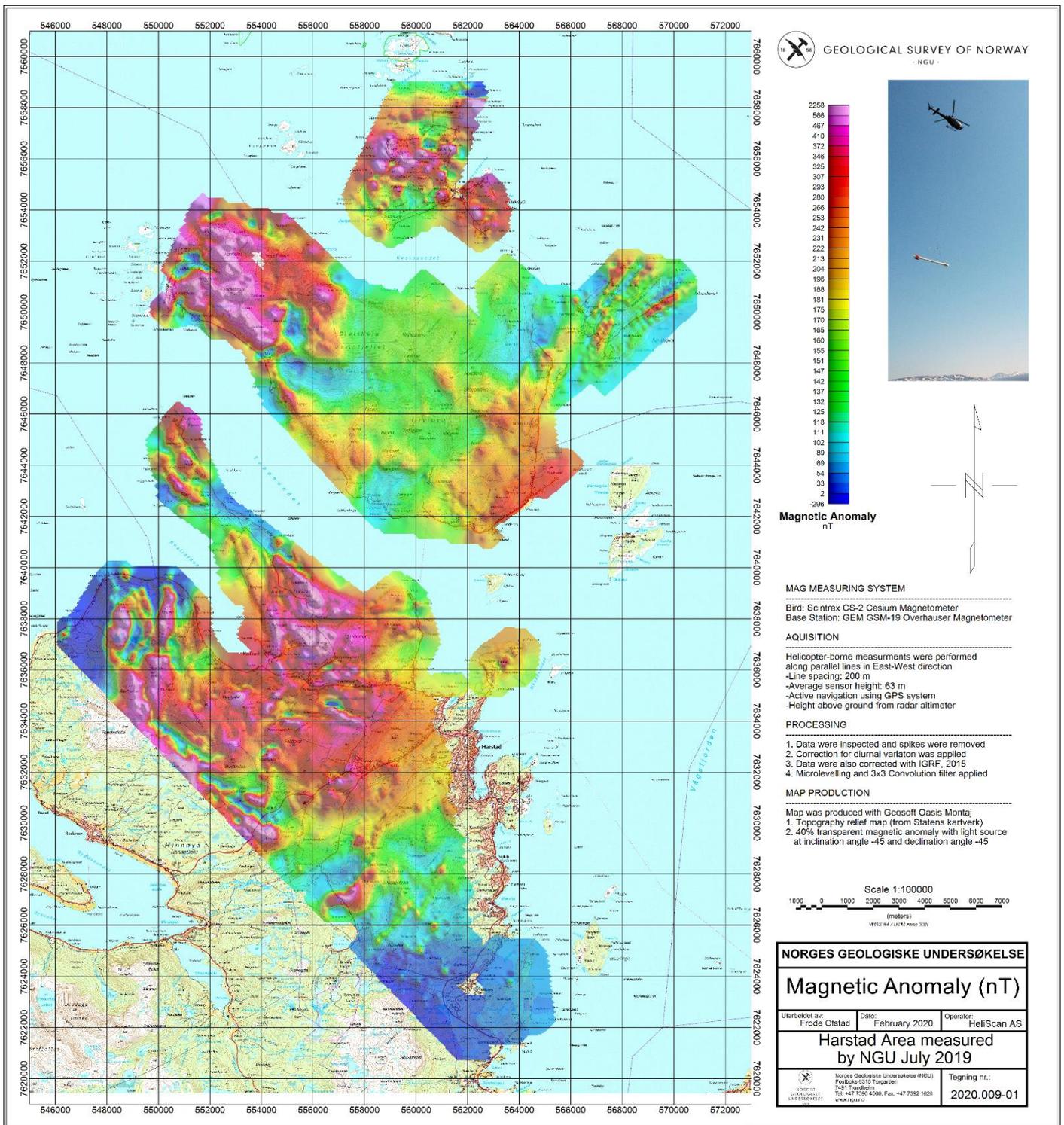


Figure 5: Total Magnetic Field

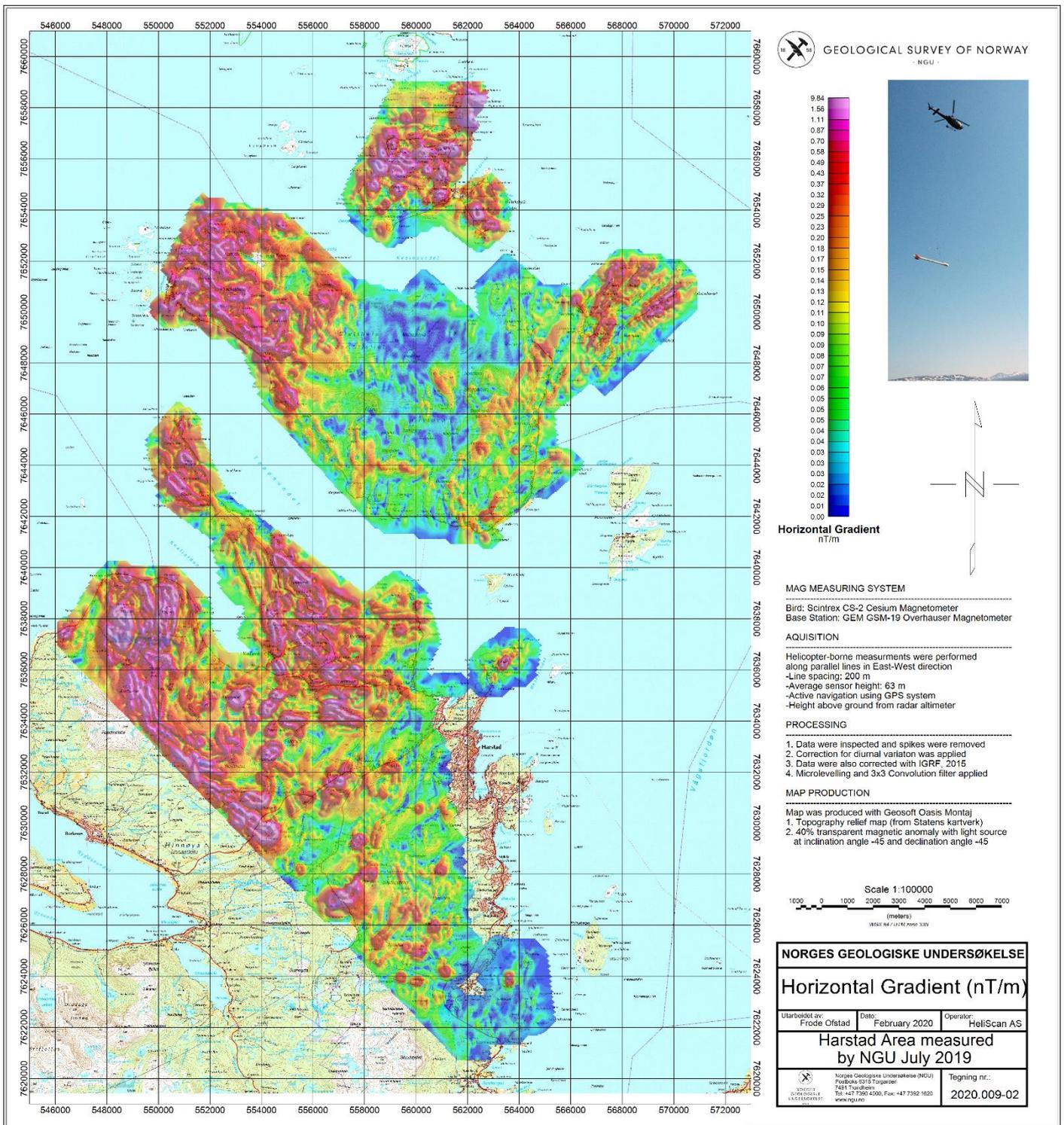


Figure 6: Magnetic Horizontal Gradient

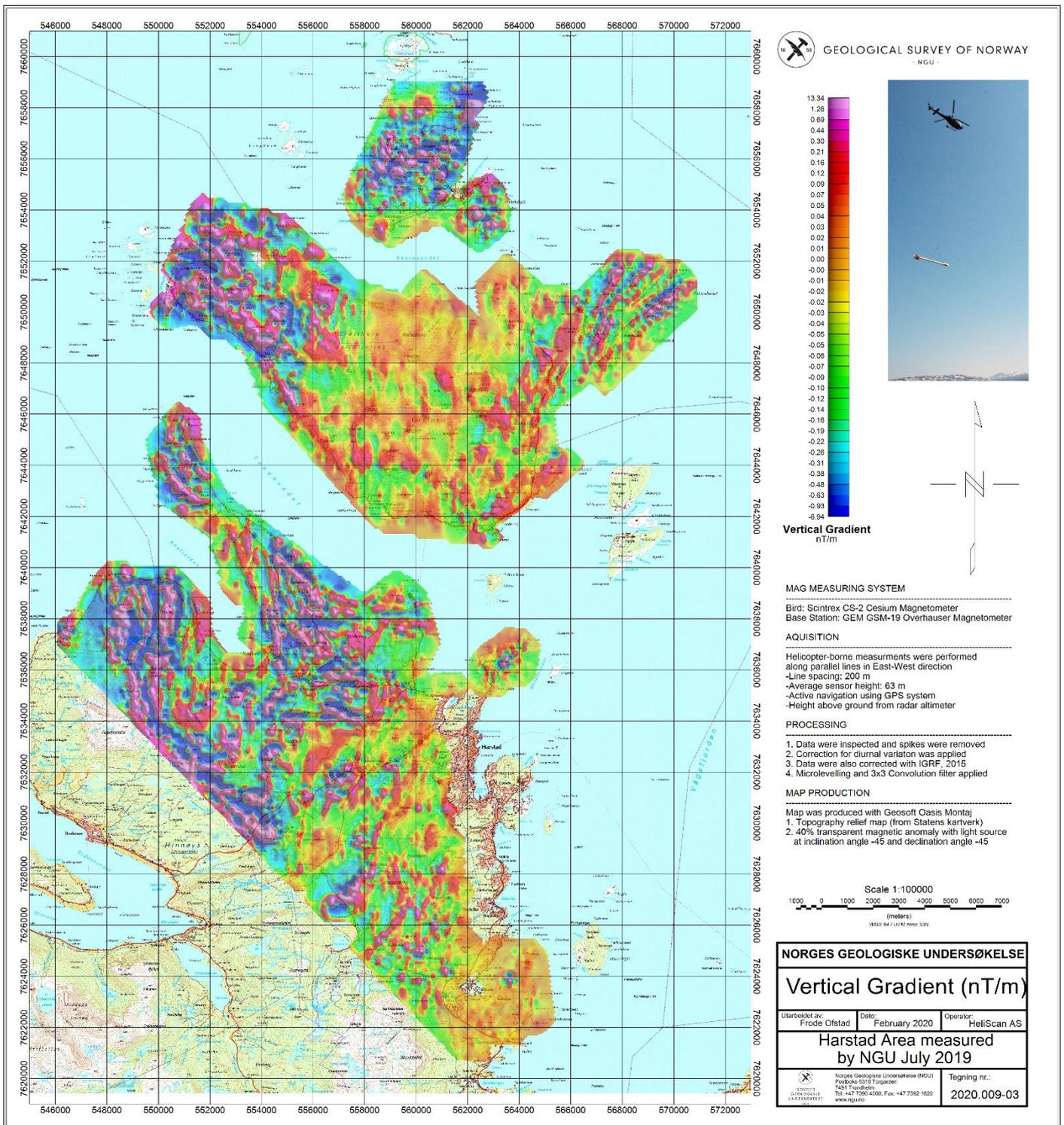


Figure 7: Magnetic Vertical Gradient

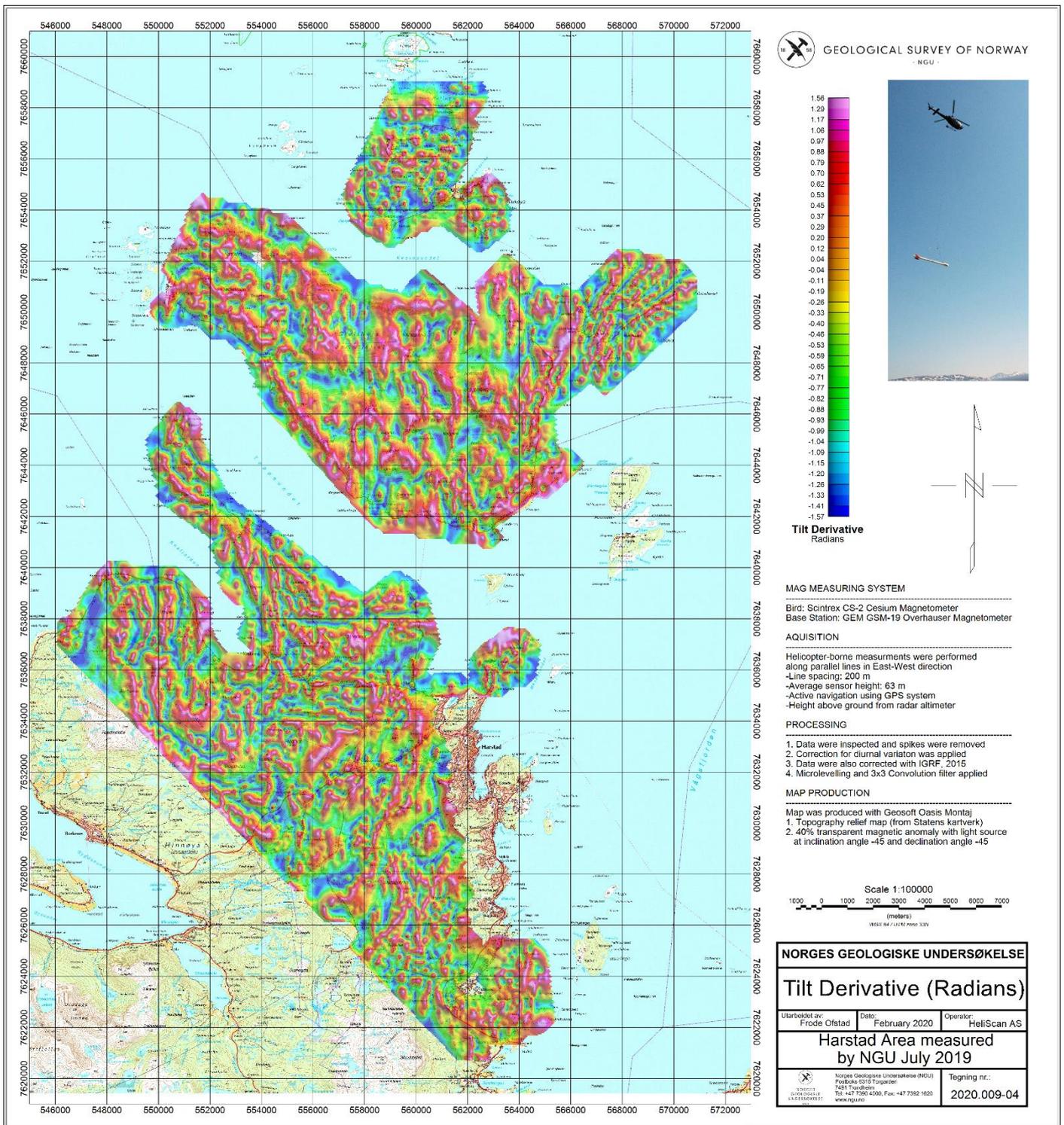


Figure 8: Magnetic Tilt Derivative

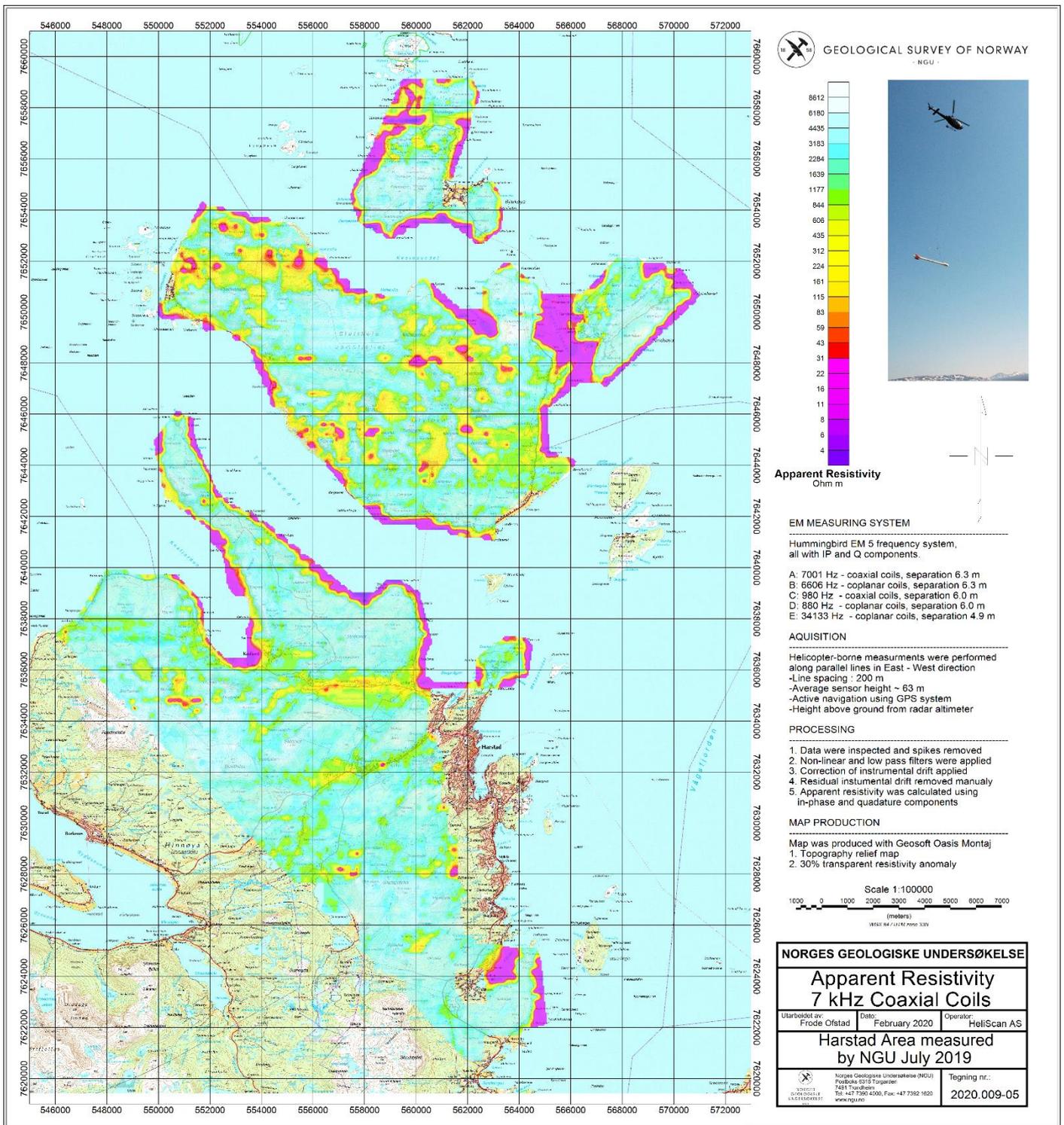


Figure 9: Apparent Resistivity. Frequency 7000 Hz, Coaxial coils

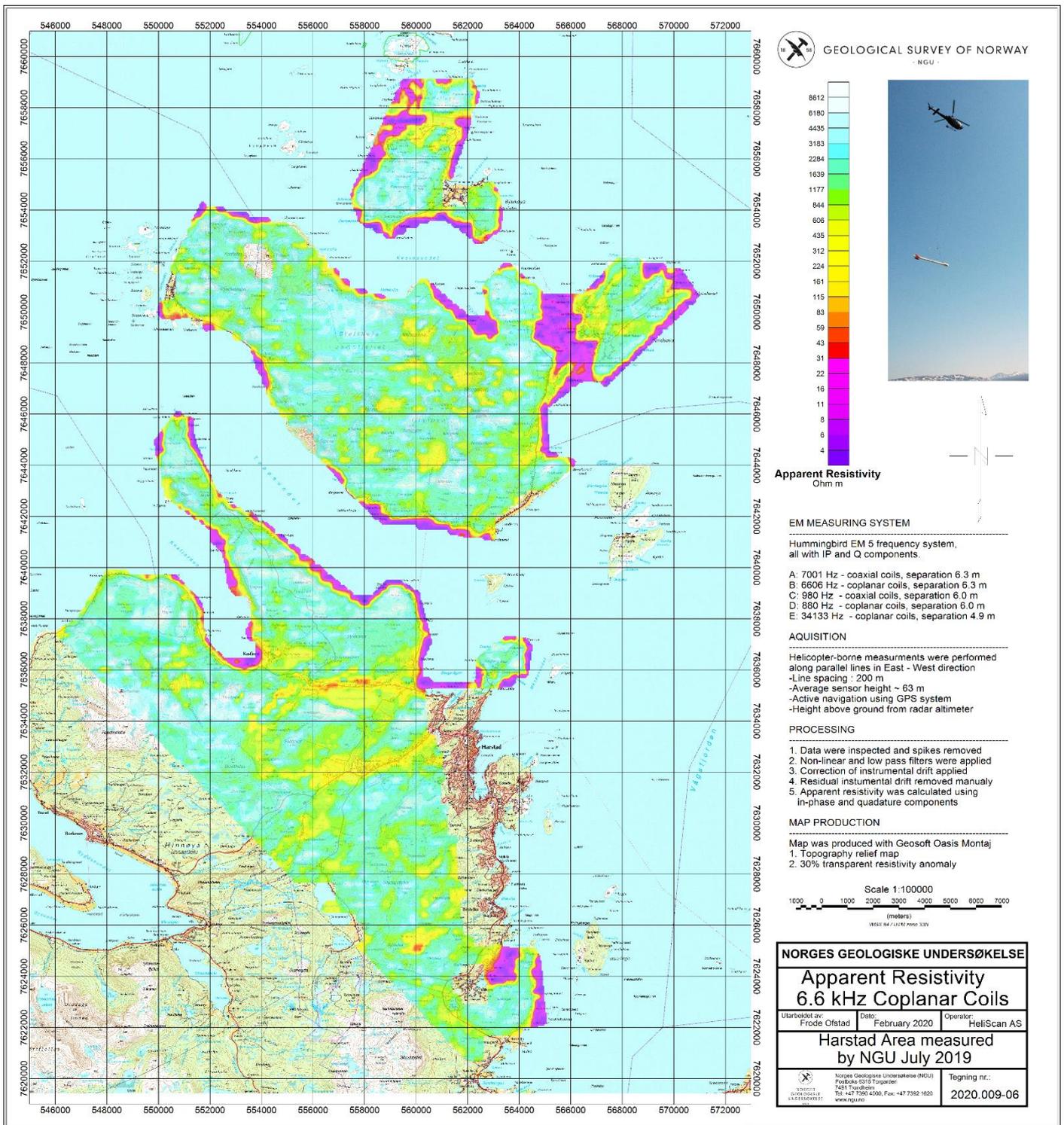


Figure 10: Apparent Resistivity. Frequency 6600 Hz, Coplanar coils

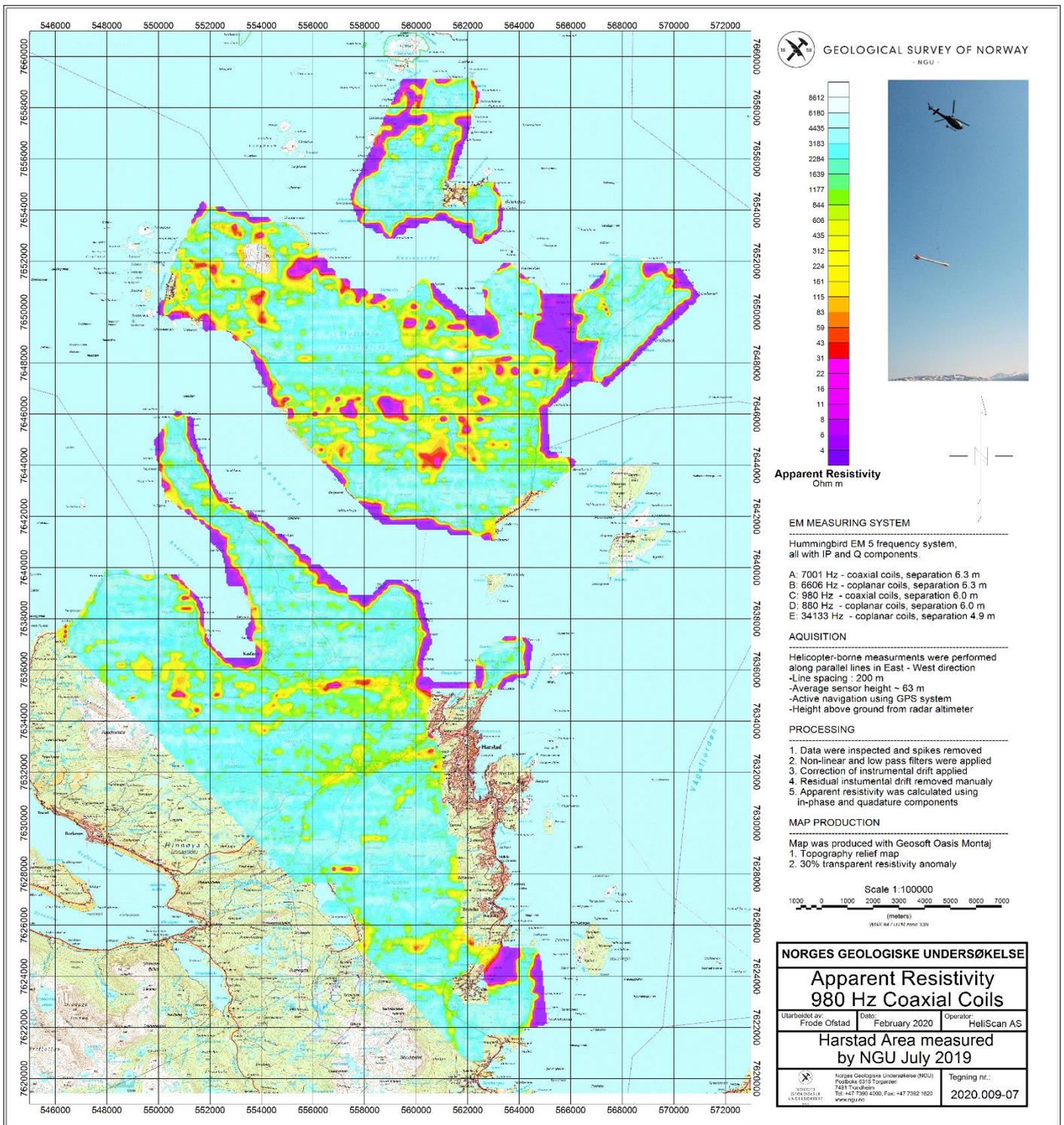


Figure 11: Apparent Resistivity. Frequency 980 Hz, Coaxial coils

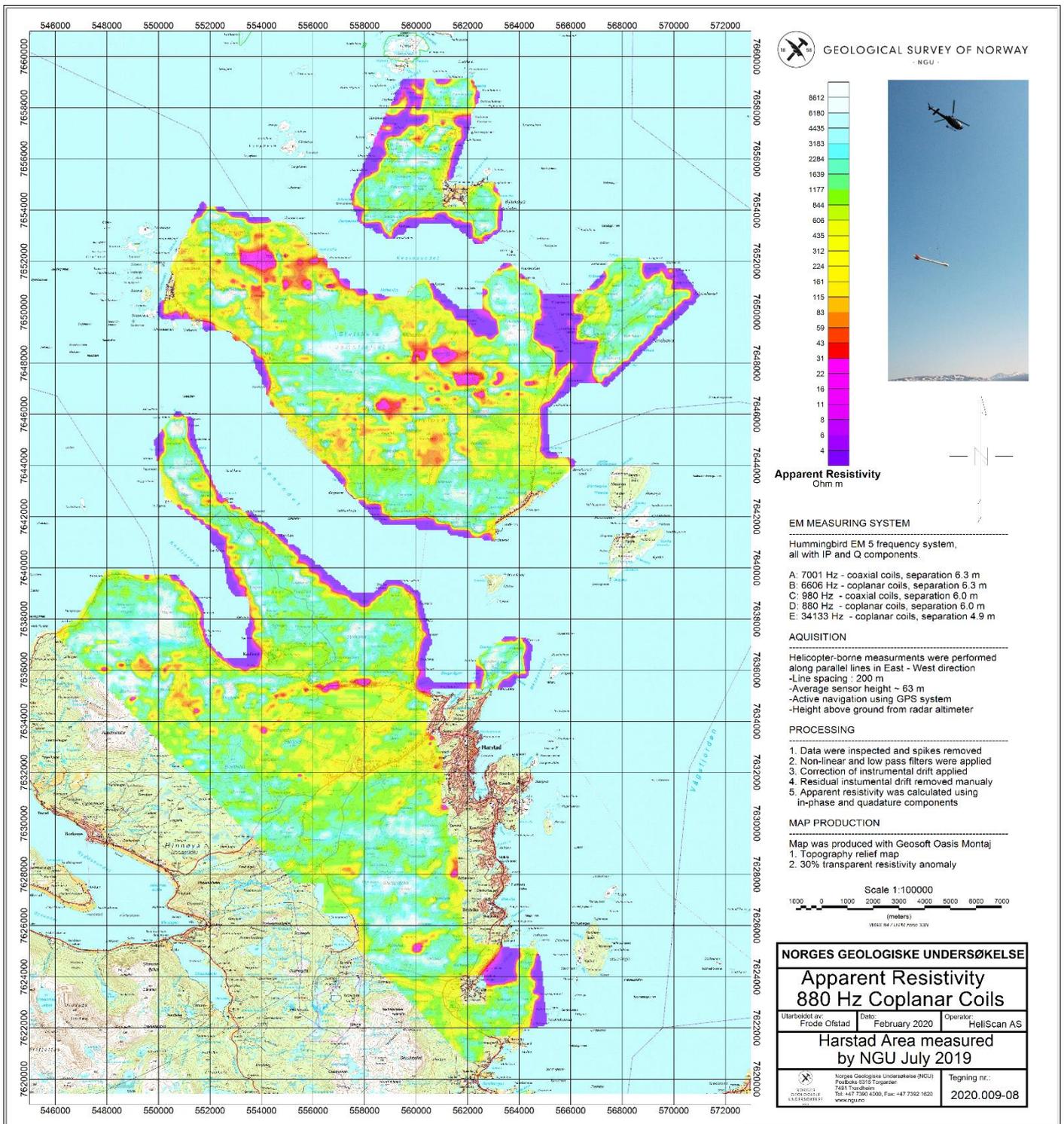


Figure 12: Apparent Resistivity. Frequency 880 Hz, Coplanar coils

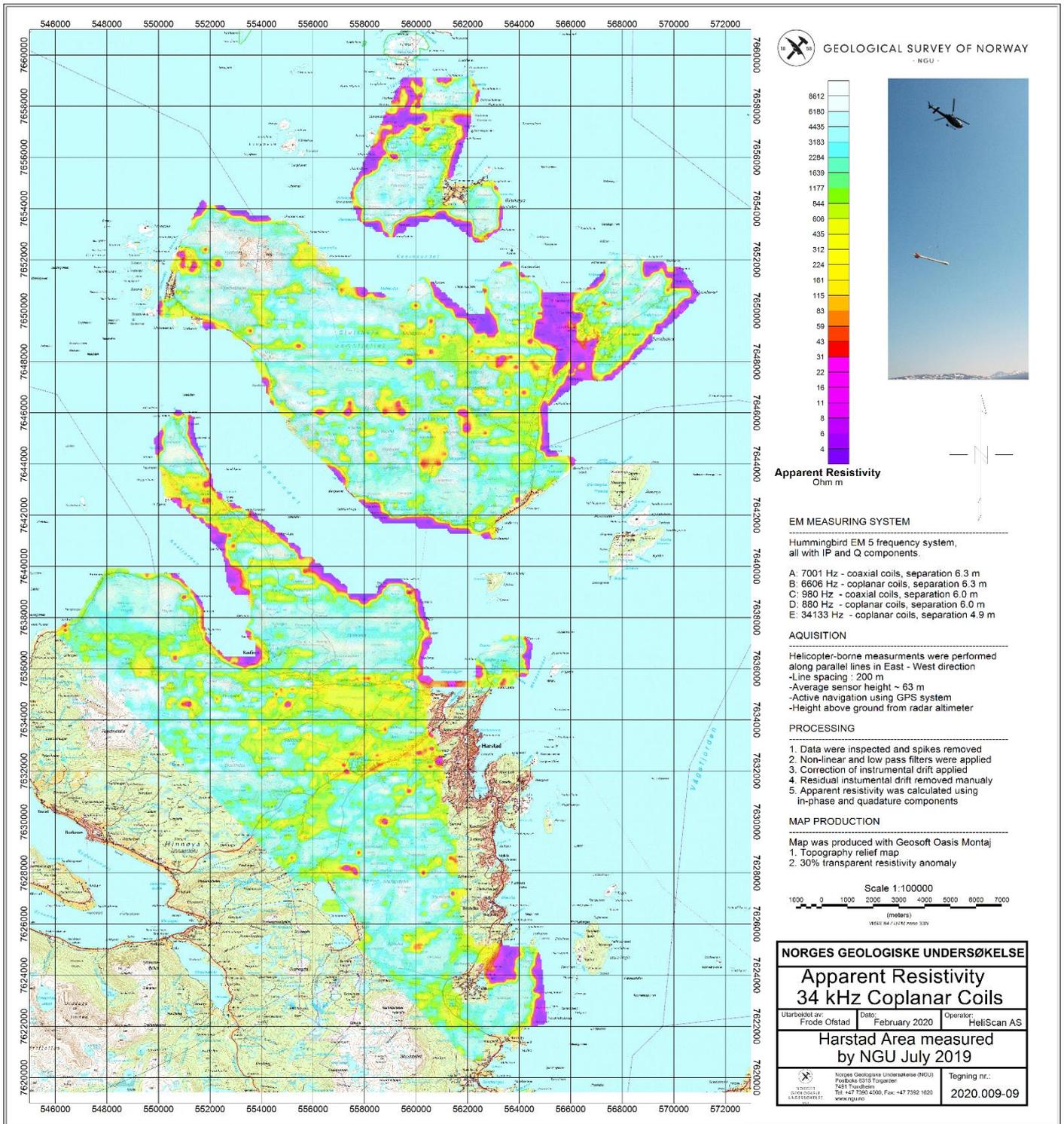


Figure 13: Apparent Resistivity. Frequency 34133 Hz, Coplanar coils

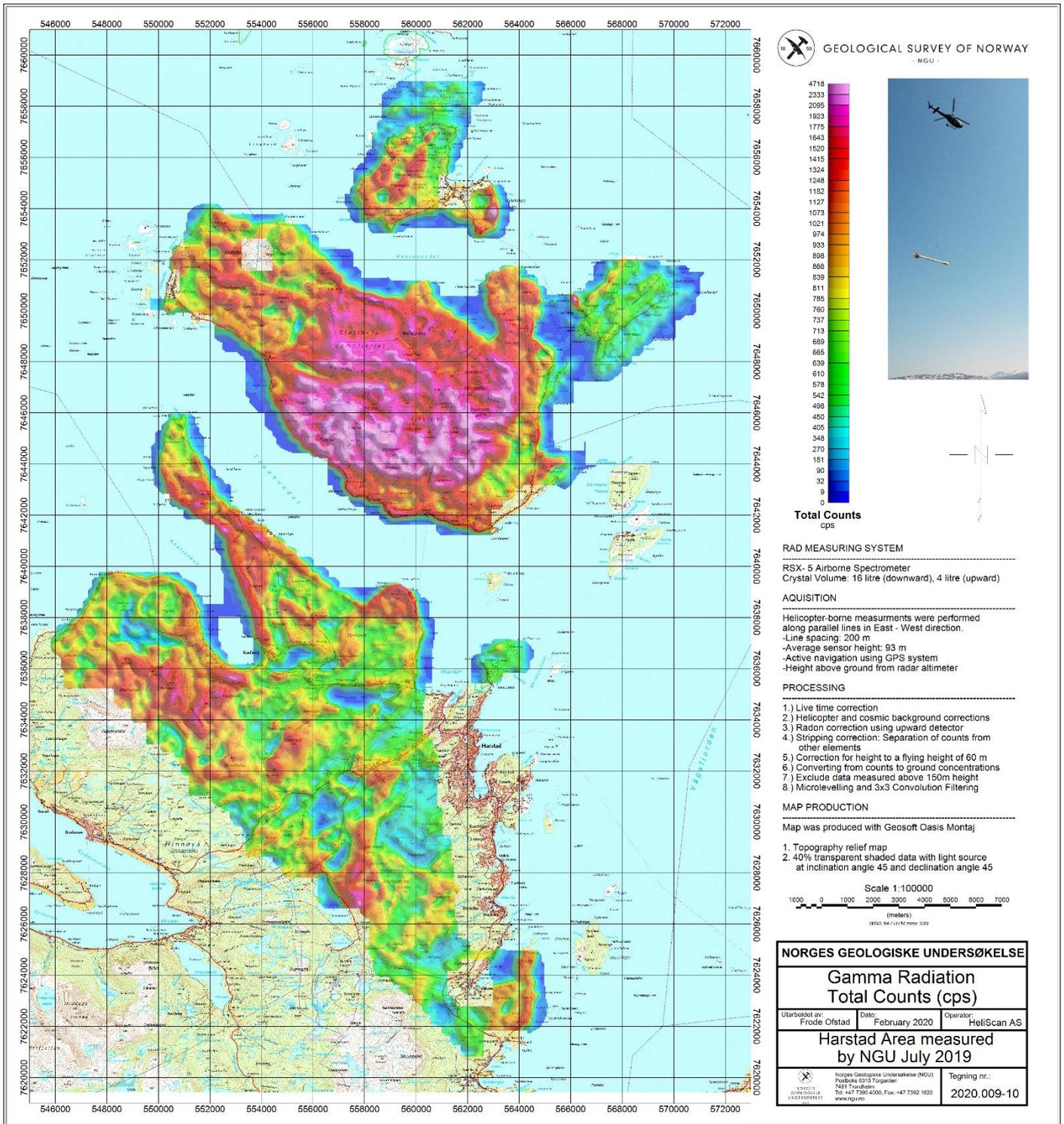


Figure 14: Radiometric Total counts

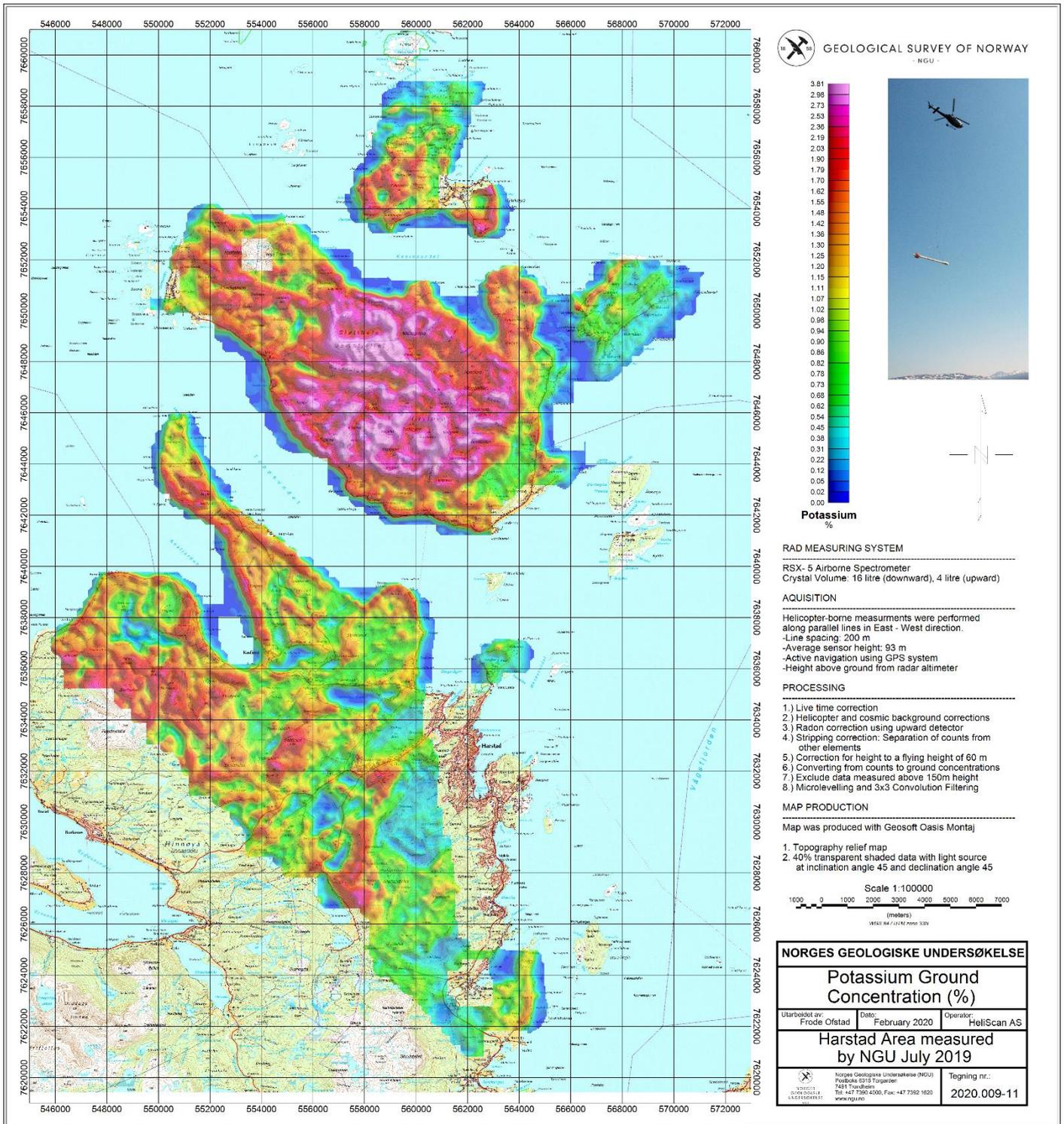


Figure 15: Potassium ground concentration

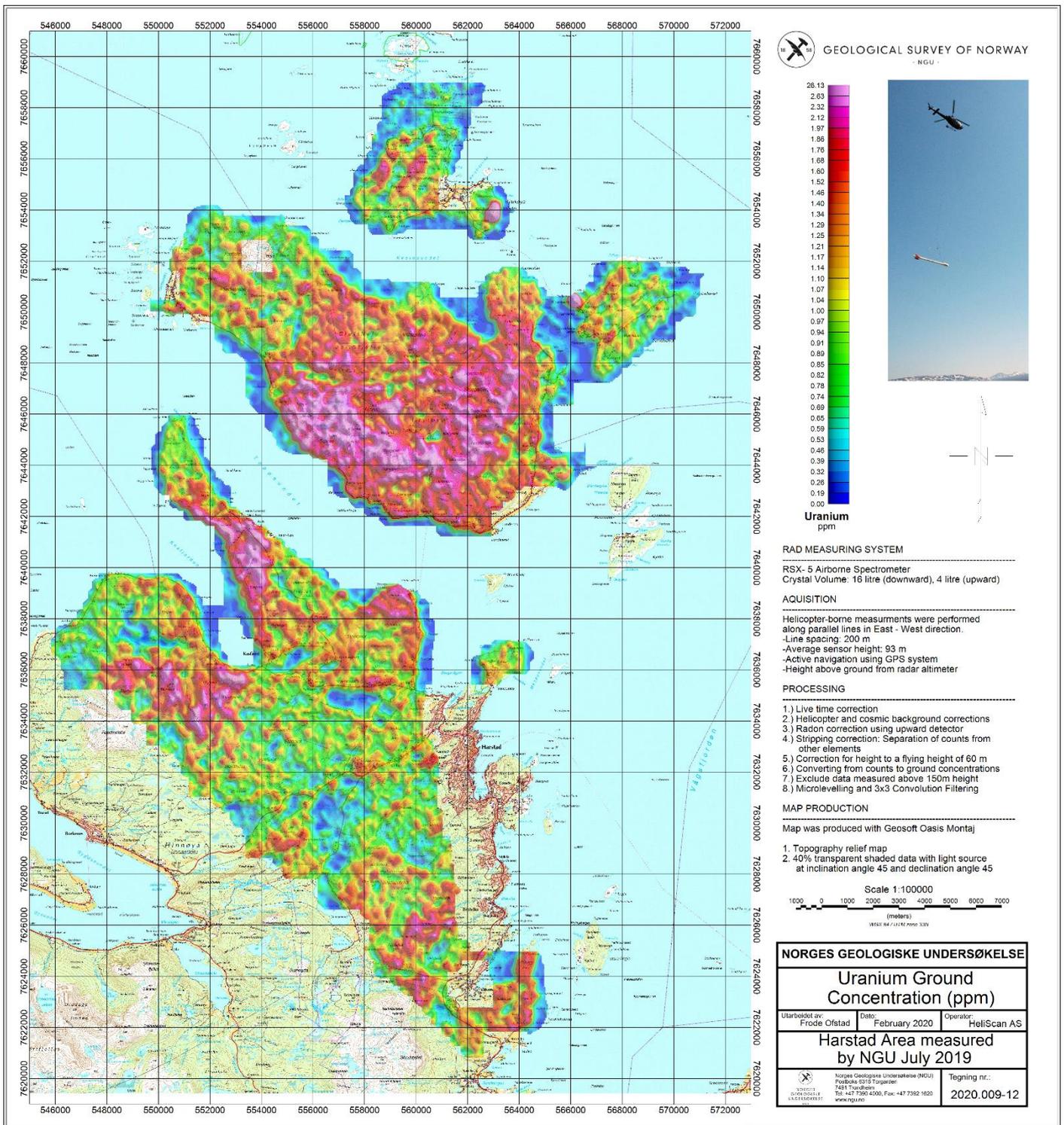


Figure 16: Uranium ground concentration

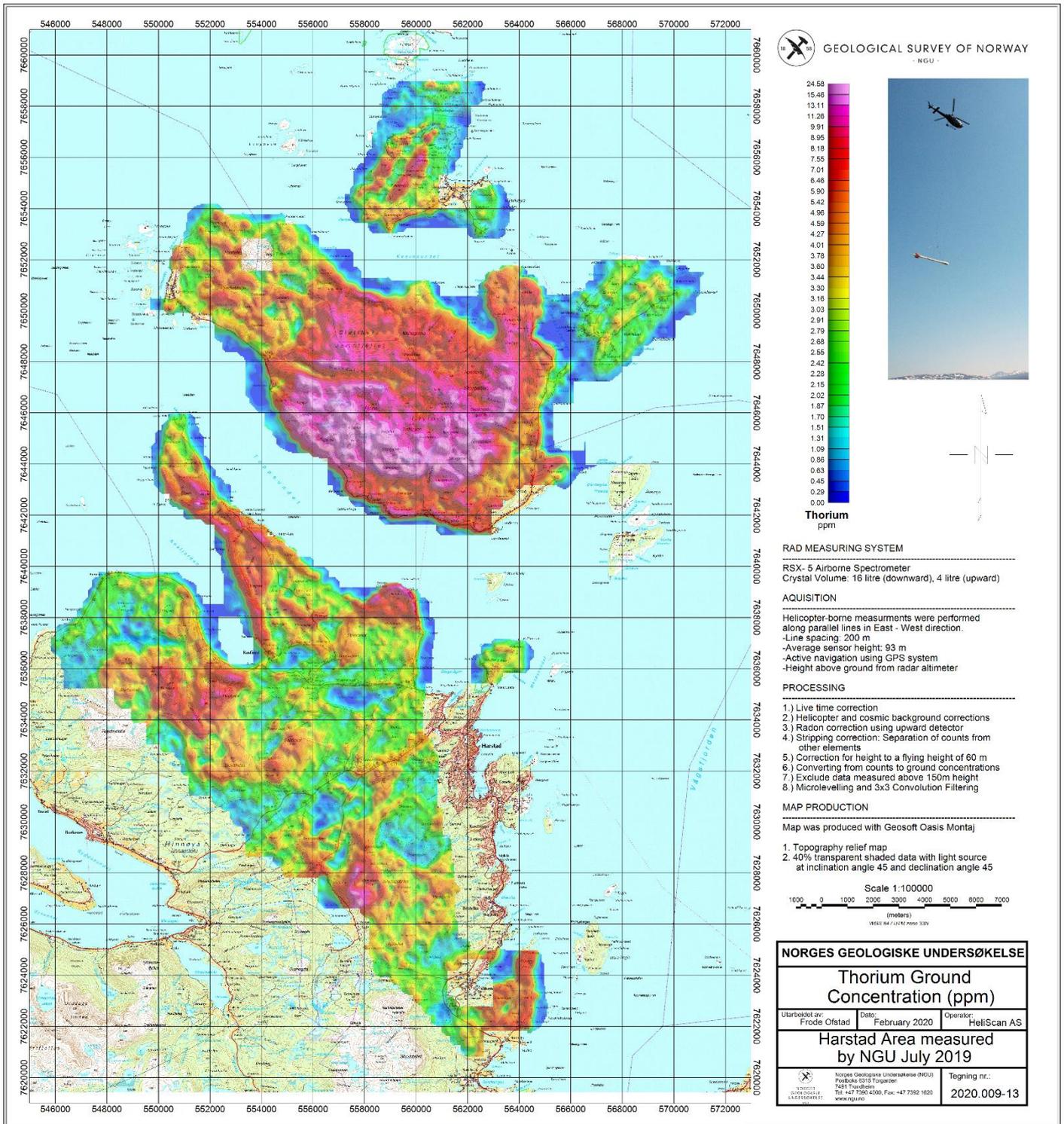


Figure 17: Thorium ground concentration

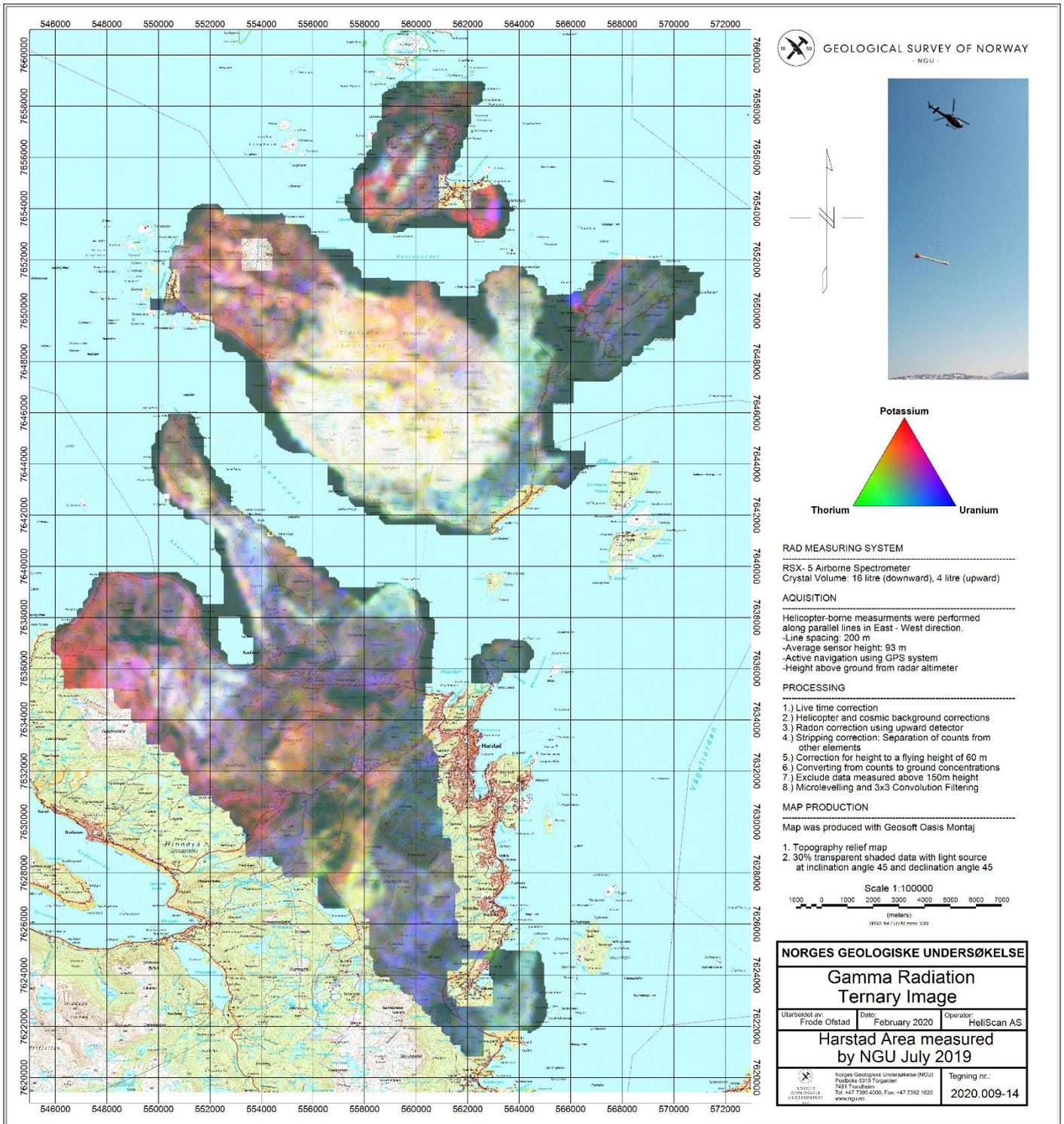


Figure 18: Radiometric Ternary Image



GEOLOGICAL
SURVEY OF
NORWAY

· NGU ·

Geological Survey of Norway
PO Box 6315, Sluppen
N-7491 Trondheim, Norway

Visitor address
Leiv Eirikssons vei 39
7040 Trondheim

Tel (+ 47) 73 90 40 00
E-mail ngu@ngu.no
Web www.ngu.no/en-gb/