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The Thermal Springs of Bockfjord, Svalbard -  
Hydrochemical Data Report

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Sammendrag:  In July 1996, a team of researchers visited the "Troll" and "Jotun" springs of northern Svalbard to carry out mapping and sampling activities.  The main conclusions of the hydrochemical survey will be reported in scientific articles in the journals "Geothermics" and "NGU Bulletin". This report serves to document the data used in these articles and to present an unpublished assessment of analytical techniques applied to the trace element geochemistry of solid (rock and travertine) samples.			
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## 1. FOREWORD

The following text is based on an expanded version of an article submitted to the journal *Geothermics* and due for publication in 1998.

The Appendices document the raw hydrochemical and geochemical data from the investigation, and present the data in the form of x-y plots and boxplots produced using the geochemical statistical and visualisation software, DAS.

## 2. INTRODUCTION

The two groups of thermal springs at Bockfjord, the Troll springs and the Jotun springs (Fig. 1), at almost 80°N latitude on Svalbard, probably represent the northernmost documented subaerial thermal springs in the World. In July 1996 a multidisciplinary team of scientists visited the spring sites to examine the hydrochemistry, flow regime, isotopic composition, geological situation and microbiology of the springs with the following objectives:

- (i) to document and understand a unique natural phenomenon which, although currently in a near-pristine state, is likely to become increasingly visited by tourists and scientists in the near future (Høgvard et al. 1997).
- (ii) to increase our understanding of groundwater recharge and discharge processes in deeply permafrosted terrains (see e.g. Lauritzen 1991, Haldorsen & Lauritzen 1993).
- (iii) to explore novel ways of obtaining information on past climatic variations on Svalbard from the isotopic signature preserved in the springs' travertine deposits.

This report seeks to document the hydrochemistry of major and selected minor elements in the spring waters, together with possible geothermometric interpretations of the waters' composition. Future papers will address the trace element and isotope chemistry of the waters, host rocks and travertine deposits, and the interpretation of the hydrogeological regime.

### **3. THE THERMAL SPRINGS OF SVALBARD**

In a permafrost terrain, where the annual average surface temperature is  $< 0^{\circ}\text{C}$ , any persistent groundwater spring may be regarded as thermal according to strict definitions. There are typically three types of springs on Svalbard:

Type 1: during summer, melting of permafrost in the active layer leads to the creation of very shallow groundwater systems and springs, discharging at temperatures of near  $0^{\circ}\text{C}$ .

Type 2: springs derived from systems of sub-permafrost groundwater. Under much of Svalbard, permafrost may exceed 300 m thickness (Orvin 1944). Groundwater recharge is believed to take place below warm-based glaciers, where permafrost is absent (Liestøl 1976, Lauritzen 1991, Haldorsen & Lauritzen 1993, Lauritzen & Bottrell 1994). Groundwater flows beneath the thick permafrost layer to discharge at springs, which may be high discharge karstic springs or low discharge springs which may form pingos, where there are open taliks in the permafrost. The location of the base of the permafrost and the temperature of the groundwater depend on a complex interplay of several factors such as surface climate and topography, geothermal heat flux, groundwater flux and frictional heat generated by groundwater flow. Lauritzen (1991) and Haldorsen & Lauritzen (1993) note that much sub-permafrost groundwater discharges at temperatures of around  $4^{\circ}\text{C}$ .

Type 3: Svalbard boasts a number of springs which are more impressively thermal, often with significant discharges (Orvin 1944). Examples of  $\text{H}_2\text{S}$ -rich, saline springs with temperatures of up to  $15^{\circ}\text{C}$  were described at Fisneset ('Fart Point') by Lauritzen & Bottrell (1994). These lie on the Hornsund fault zone of Southern Svalbard and the salinity appears to be derived from a deep thermal brine rather than modern sea water. The Troll and Jotun springs bear many similarities to the Fisneset springs but lie much further north and are considerably hotter.

### **4. THE BOCKFJORD SPRINGS**

#### **4.1 History**

The springs are believed to have been documented first in the late 19<sup>th</sup> century, and they are described in some detail by Hoel & Holtedahl (1911, 1913) and Hoel (1914). These authors measured water temperatures of  $24.5^{\circ}\text{C}$  at the Jotun springs and up to  $28.3^{\circ}\text{C}$  at the Troll springs. They also provide details of some chemical analyses performed at that time (Table 1), which are of interest in comparison with the current analytical suite. Hoel and Holtedahl (1911) also observe that the waters are of dominantly sodium bicarbonate character.

## 4.2 Geological Setting

The springs occur almost on the line of a major fault which appears to follow the western shore of Bockfjord. The fault has a large downthrow on the eastern side and forms the western boundary of a major graben structure, bringing the Devonian 'Old Red' sandstones and siltstones of the Wood Bay formation into contact with the Precambrian Hecla Hoek complex (Gjelsvik 1979, Hjelle 1993). The Hecla Hoek consists of a ca. 2 km wide outcrop of limestone and dolomite marbles in the east (adjacent to the fault), overlying a thick sequence of gneisses and migmatites further west. The complex also includes phyllites, mica schists and metadolerites and is believed to be of mid-late Proterozoic age.

The Troll springs appear to emerge from brecciated fine-grained marbles of the Hecla Hoek complex, although an apparent outcrop of greenish basal Devonian sandstones can be observed in the shoreline several tens of metres east of Troll Spring Area 1. The geology of the Jotun Spring area is more difficult to interpret. Outcrops of Hecla Hoek marble are found a short distance below the springs although much of the frost-shattered material immediately around the springs is mica schist, also from the Hecla Hoek. Gee & Moody-Stuart (1966) describe the occurrence of Devonian sandstones on the coast north of Sverrefjellet.

The Bockfjord area exhibits significant evidence of geologically recent crustal uplift, mantle lithosphere thinning and volcanism (Vågnes & Amundsen 1993). The main phase of uplift and volcanism was in the Neogene, resulting in tholeiitic basalt caps to several of the mountains around Bockfjord. A later phase of alkaline olivine basalt and breccia eruption occurred near Bockfjord in the Quaternary, most notably from the volcanic centre of Sverrefjell, between the Troll and Jotun springs (Fig. 1). The volcanic centres of Sverrefjell and Sigurd fjell also lie on the same graben boundary fault as the springs. These volcanoes may have been active as recently as the last interstadial (Skjelkvåle et al. 1989).

Vågnes and Amundsen (1993) suggest a geothermal heat-flow anomaly extending south from Yermak to Bockfjord, with a heat flow at Bockfjord of some  $130 \text{ mW m}^{-2}$ . Xenolithic evidence from Quaternary volcanism at Bockfjord suggest temperatures of  $550^\circ\text{C}$  at 7 km depth (Amundsen et al. 1988), corresponding to an average temperature gradient of some  $0.079^\circ\text{C/m}$ .

The age of the springs themselves is unknown. W.B. Harland and M.J. Hambrey are recorded by Langangen (1979) as opining that the springs are unlikely to be older than 4-6500 years. However, if Liestøl (1976) is correct in his theory that many springs and pingos on Svalbard first formed below sea-level, the springs might be older. They might even be as old as the volcanoes and may have formed subglacial spring outlets during the maximum of the last glaciation, as well as submarine spring outlets during the periods of high sea-level.

The Troll springs can be seen to occur below a fossil shore line and at the same topographic level as apparently marine clays. One should not, however, conclude that the springs only became 'activated' on their post-glacial emergence from the sea. It is also possible that the springs may have survived marine inundation at times of higher sea-level; submarine thermal springs are known today from the Tempelfjord area of Svalbard (Hoel and Holtedahl 1911, 1913).

### **4.3 Spring Morphology**

The mountains to the south and west of Bockfjord are heavily glaciated. Lobes of glacier extend down almost to sea-level at the southern ends of Bockfjord and Woodfjord, with associated terminal moraine complexes and sandurs with braided river systems. Smaller glacial lobes also extend almost to sea-level from side valleys along the western side of Bockfjord.

There are six discrete areas of travertine deposition associated with groups of active thermal springs at the Troll springs area. The spring areas are numbered 1 to 6 from the south. The springs typically deposit wide shallowly concave basin complexes of travertine, each basin with a low wall. The hottest waters emerge at the top of the basin complex (the so called 'A'-position - Fig. 2) and cascade down through the basin complex, often seeping into the travertine near the base of the complex. Below the basin complex sheets of massive travertine breccia tend to occur. This presumably represents frost-shattered material cemented by the carbonate-saturated thermal groundwaters seeping through it. The thermal waters may re-emerge on their way down to the fjord at 'B'- and 'C' positions or they may not re-emerge until the level of the sandur at the southern end of the fjord ('D' position). The spring nomenclature is such that 3A2 means 'the second spring of the third spring area from the south, A-position spring'.

The top of the travertine deposits of Spring Area 6 is topographically higher than areas 1-5, but no active spring emerges from the A-position. Two thermal springs emerge lower down the travertine deposits in what is strictly a 'B'-position. As these are highest level waters of spring area 6, these are, for the purposes of this paper, designated 6A1 and 6A2.

A seventh fossil spring area is observed between areas 5 & 6 at a considerably higher topographic level. It is argued that the two northern Spring Areas (6 and 7) were once active at times of greater groundwater head in the thermal system enabling travertine systems to develop at a higher topographic level than today. It is of course possible, that the effect is due to post glacial emergence of the Svalbard coast and the actual head relative to sea-level has not changed. Nevertheless, in several spring areas, the size of the travertine basin complexes



often appears disproportionately large relative to the size of the current springs. This would indicate that groundwater heads and flow have genuinely been larger in the past, a conclusion supported by Hoel & Høltedahl (1911, 1913).

The Jotun springs have not developed any terraced travertine basins. They have rather developed huge convex mounds of travertine. At Spring Area 1 (east), three small springs (1JA1-3) emerge from the flattened top surface of the mound. At Spring Area 2 (west), no springs emerge from the top although one emerges from the flanks of the mound (designated 1JB). The remains of a third travertine area can be found at an apparently higher level between Jotun Spring Areas 1 and 2.

Several springs appear to evolve gas bubbles regularly. The composition of the gas has not been determined, although at 3A1 the smell of H<sub>2</sub>S can clearly be distinguished.

The flow from the springs is impossible to assess as much of the run-off from the A-springs infiltrates directly into the permeable travertine as groundwater. It may be that only a minor proportion of the total spring flow is observed as surficial overflow from the springs.

Another thermal spring area has been reported by Otto Salvigsen, in a position between the Troll and Jotun springs, on the west of Bockfjord. Hitherto, it has not been documented in any detail.

## **5. SAMPLING PROGRAM**

At almost all A-springs and a number of B-, C- and D-springs, pH, temperature and Eh were measured in the field (Appendix 4). At selected springs (seven A-springs at Troll, 2 D-springs at Troll, and Jotun springs 1JA3 and 2JB), field analyses of alkalinity were carried out and samples were collected for laboratory analysis. The selected A springs include the hottest springs of each spring area (except area 4). Additionally a cold spring between Spring Areas 3 and 4 at Troll was sampled (Sample K1). This is believed to be a permanent spring, as indicated by a verdant growth of spring moss, and is also believed to represent a discharge of non-thermal sub-permafrost groundwater (Type 2 - see above), an interpretation confirmed by the rather mature hydrochemistry of the water. A sample was also taken from a thaw seepage from the active layer between Spring Area 4 and 5 (Type 1, sample K2).

Field measurements of pH, temperature and Eh were made using a Yellow Springs YSI 3500 portable meter, calibrated regularly against buffer solutions of pH 4, 7 and 10. Alkalinity

titrations were performed in the field using an Aquamerck kit with a pH=4.3 mixed indicator. An average of three separate titrations was taken.

On the first day 10 ml water samples for radon analysis were taken by injection into pre-filled vials containing a scintillation liquid (for method see Reimann et al. 1996). These were returned by helicopter to be analysed at the National Radiation Protection Authority within three days, using scintillation counting. Duplicate samples were taken, whose analysis indicated the very good reproducibility of the technique, even at low radon concentrations.

For each sampled spring 2 x 100 ml samples of filtered water were taken in 100 ml polyethylene bottles. These were rinsed thrice with spring water prior to sampling and twice with filtered spring water. Samples were filtered with 0.45 µm Millipore filters. On arrival at the Geological Survey of Norway (NGU) laboratory, one flask of each sample was acidified with 0.5 ml/100 ml concentrated Suprapure (Merck) HNO<sub>3</sub> to remobilize and preserve any precipitated or adsorbed metals. This sample was then analysed for 32 parameters at NGU by Inductively Coupled Plasma spectroscopy (ICP-AES - Ødegård and Andreassen 1987); Appendix 1. The second flask of each sample was sent to the Norwegian Agricultural University (NLH) for anion analysis by Ion Chromatography (IC): Appendix 2. The flask was then returned to NGU for confirmatory IC analysis: Appendix 1.

## 6. RESULTS

XRD analysis of travertine material from three spring areas (Troll springs, areas 1, 3 and 5) confirmed it to be composed of almost pure calcite (Appendix 6). No significant trace of either dolomite, aragonite or other major minerals was detected. The possible presence of minor amounts of monohydrocalcite, quartz and mica were indicated; the latter two could be detrital minerals.

The results of the water analyses are largely presented graphically in Figures 3-9 and in Tables 1 and 2. Ionic balances were found to be very good, all within 5 %. At the Jotun springs, alkalinity could not be determined in the field for logistical reasons. In the diagrams and calculations, alkalinity has been estimated from the ion balance.

Three groups of thermal waters can be distinguished purely on the basis of the ions Na, Ca, Mg, Cl<sup>-</sup>, SO<sub>4</sub><sup>=</sup> and HCO<sub>3</sub><sup>-</sup>.

- (i) the southern Troll Spring Areas 1, 2 and 3 with Na-(Ca)-HCO<sub>3</sub><sup>-</sup> composition

- (ii) the northern Troll Spring Areas 5 and 6 with Na-HCO<sub>3</sub><sup>-</sup> waters.
- (iii) the Jotun springs with Na-Cl<sup>-</sup> waters.

The waters within each group all have similar ionic composition in terms of proportions of major ions. Figure 3 shows pie diagrams of representatives of each water type, while Figure 4 plots the waters on a Durov diagram. The northern Troll springs are more similar to the southern Troll springs than the Jotun springs. It is also noteworthy that the sub-permafrost non-thermal spring K1 displays a Ca-(Na)-HCO<sub>3</sub><sup>-</sup> composition which bears some resemblance to the southern Troll thermal waters. The thaw water K2 has a Ca-HCO<sub>3</sub><sup>-</sup> composition reflecting the dominant lithology in the active layer - shattered marble fragments.

Figures 5-9 and Table 2 show the relationship between various parameters and temperature and will not be commented in detail, although the reader's attention is drawn to the following observations:

- (i) pH and Eh tend to decrease with increasing temperature.
- (ii) B, C and D springs generally exhibit lower temperatures and higher Eh values than A springs. Plots of Na vs. Cl<sup>-</sup> (Fig. 8: and also Mg and Sr vs. Cl<sup>-</sup>, not shown) indicate a limited degree of proportionate dilution from warmer (A) to colder (D) springs. A plot of Ca vs. Cl<sup>-</sup> (Fig. 8) indicates, however, that calcium is being disproportionately attenuated to a much greater degree than chloride from the A to D springs.
- (iii) alkalinities of all the thermal Troll springs are high (around 13 meq/l). The calculated alkalinities of the Jotun springs are somewhat lower, around 7-10 meq/l.
- (iv) the Na, K, SO<sub>4</sub><sup>=</sup> and Si concentrations all show extremely good correlations with each other, and to a lesser extent with Cl<sup>-</sup>. Concentrations of these parameters in all the Troll springs are similar, independent of temperature, although elevated relative to K1 and K2. 2JB contains higher concentrations of these parameters, 1JA3 the highest of all.
- (v) the Na/Cl ratio in K2 is similar to that in sea-water (i.e. marine salts in precipitation). The other springs, including K1, exhibit significant non-marine excesses of Na.
- (vi) Ca and, to a lesser extent Mg, show a better positive correlation with temperature than the parameters named in (iv)
- (vii) Br/Cl mass ratios for the thermal waters range from 0.0018 in 1JA3 to 0.0029 in 2A4i. In sea-water, the ratio varies from 0.00325-0.0043 (Horne 1969, Michard et al. 1996), with a typical value of 0.00347 being used by Edmunds (1996).
- (viii) Significantly elevated levels of Fe and Mn occur in the hotter A springs (this may be related to low Eh)
- (ix) Rn levels are replicable and are significantly lower in the thermal springs than in K1
- (x) All thermal waters exhibit reasonably consistent but elevated Na/Si (ca. 10) and Mg/Ca (ca. 0.3) mass ratios compared with K1 and K2 (Figure 9). The same applies to Sr/Ca (ca. 0.011) and Na/Ca (ca. 3) mass ratios (Table 2).

(xi) The thermal springs exhibit (Fig. 9) three groups of Na/K mass ratio; the Jotun springs at around 13.5 to 15.5, the southern Troll springs between 16 and 16.5 and the northern Troll springs at around 17.5. This grouping suggests that Na/K ratios may be a particularly interesting tool in the context of geothermometry.

## 7. INTERPRETATION

The thermal waters share some of the features exhibited by K1. This may imply either that some of the processes seen in the thermal waters can be explained by normal hydrochemical evolution occurring in the sub-permafrost zone, or that K1 includes a component of thermal water.

Observation (i) above suggests that 3A1 is the most 'original' of the Troll springs: it is the hottest, has the lowest Eh and the lowest pH and evolves large quantities of H<sub>2</sub>S. The relationships of pH and Eh to temperature may suggest that the primary waters are of sub-neutral pH. The elevated pH at lower temperature may be related to degassing of CO<sub>2</sub>. The primary waters may also be reducing, although it is also conceivable that reducing, H<sub>2</sub>S-rich conditions are imposed during the water's slow percolation through biologically active sediments associated with the spring (although unless the biota refreshes itself, this process is likely to be self-limiting).

Observation (ii) suggests that the B, C and D springs represent re-emergences of the thermal waters after they have been subject to a limited degree of dilution by surficial waters and/or shallow groundwaters and lost heat to the air and to the travertine mass. The Ca vs. Cl<sup>-</sup> plot suggests that calcium is being attenuated by other processes than dilution, probably by calcite precipitation. The fact that neither Mg nor Sr exhibit a similar degree of attenuation is compatible with the lack of dolomite or strontianite in the XRD spectrum.

Observations (iv), (x) and (xi) suggest that the waters may be partially derived from a deep-seated saline thermal water whose composition in terms of Na, K, Si and SO<sub>4</sub><sup>=</sup> is determined by a water-rock equilibrium. Michard et al. (1996) suggest that sodium concentrations can be taken as a good indicator of the extent of water-rock interaction in thermal waters of the crystalline basement of Switzerland. The lesser degree of correlation of these parameters with Cl<sup>-</sup> and the variable, significant sodium-to-chloride excess (observation v), suggest that chloride is not wholly derived from water-rock interaction. This hypothesis is supported by observation (vi): Nordstrom et al. (1989), Michard et al. (1996) and Edmunds (1996) regard an excess Br/Cl<sup>-</sup> ratio as being indicative of chloride-derivation from "exotic" origins such as water-rock interaction (fluid inclusions, radiogenesis) or atmospheric pollution. The Br/Cl<sup>-</sup>

ratios found in the Svalbard springs indicate a slight bromide deficit relative to sea-water. Such a deficit may indicate at least partial derivation of chloride salinity from halite evaporites or conceivably preferential adsorption of bromide on organic matter (Edmunds 1996). It is also conceivable that palaeo-seawater may have had a differing Br<sup>-</sup>/Cl<sup>-</sup> ratio to today.

Overall, the evidence suggests that the composition of the saline water body may result from a fossil sea-water modified by dilution, water-rock interaction, rather than purely from water-rock interaction. Another possibility is that the saline water may be derived from evaporite minerals. Evaporite deposits are not known on North Svalbard, but it cannot be excluded that they may exist at depth in the Devonian succession.

## 7.1 Geothermometry

At particular temperatures, common assemblages of minerals will tend towards equilibrium with a given water chemistry. Several researchers have noted that, for certain parameters or ratios of parameters, the relationship between temperature and composition will be stable and predictable. These parameters or ratios of parameters are known as geothermometers. In order for these to work, one has to assume that effects of dilution are insignificant and that thermodynamic equilibrium has been attained at reservoir temperature. In general, geothermometers based on ratios will be more resistant to dilution effects than those based on absolute concentrations. In addition, it should be realised that the temperature indicated by the geothermometer is not necessarily the maximum temperature of the water, but the temperature at which mineral and water phases were last in equilibrium with respect to the phases in question.

Several well-known geothermometers are summarised by Fournier (1981), who recommends three as being particularly appropriate to most low temperature thermal systems, namely the silica, the Na/K and the Na-Ca-K thermometers (Table 3). The simplest of these is the silica thermometer, which is simply based on the solubility of silica phases at differing temperatures. The geothermometer, being based on an absolute concentration, is susceptible to dilution and is also sensitive to assumptions made about the equilibrium silica phase (quartz or chalcedony, for example). In our case the quartz geothermometer yields temperatures of some 100 - 110°C for the Troll springs and up to 150°C for the Jotun springs.

Truesdell's (1976) Na/K geothermometer yields temperatures of 133 - 143°C for the Troll springs and around 150°C for the Jotun springs. Fournier's (1981) Na/K thermometer yields values between 30 and 40°C higher.

Finally, the more complex Na-K-Ca thermometer (Fournier & Truesdell 1973) indicates temperatures of around 60°C for both the Troll and Jotun springs, following correction for Mg content as indicated by Fournier & Potter (1979). The Na-K-Ca thermometer may, however, be susceptible to the effects of dilution if the proportion of "original" hot water in a mixture is 30% or less (which it may be in this case - see below). It is especially sensitive to possible errors if the uncorrected calculated temperature is around 100°C, where the  $\beta = 1/3$  solution switches to the  $\beta = 4/3$  solution (see Fournier 1981). The Na-K-Ca thermometer may also not be applicable in the Bockfjord case as Na and K concentrations are likely to be controlled by a feldspar equilibrium at depth in silicate-dominated Hecla Hoek or older rocks. It is possible that the Ca-Mg signature is picked up at a shallower depth and lower temperature during the water's transit through carbonate-rich rocks of the upper Hecla Hoek. The Na-K-Ca thermometer is therefore of dubious applicability in this study.

The degree of dilution can be accounted for by the method of Fournier and Truesdell (1974). In essence this involves the solution of two simultaneous equations to find the factor X (the proportion of diluting cold water), and the original temperature of the hot saline water T:

$$X = (Si_T - Si_{spr}) / (Si_T - Si_{cold})$$

$$X = (H_T - H_{spr}) / (H_T - H_{cold})$$

where  $Si_T$ ,  $Si_{spr}$ ,  $Si_{cold}$  are concentrations of silicon in the original saline water at temperature T, in the sampled spring and in the cold dilution water respectively.  $H_T$ ,  $H_{spr}$  and  $H_{cold}$  are the enthalpies of the same waters. In all cases, K1, the supposed sub-permafrost groundwater, is taken as the cold diluting water. As  $Si_T$  is a direct function of temperature (the solubility of a silica phase such as quartz), as is  $H_T$ , the equations may be solved. By plotting X for a range of T, the intersection of the two lines gives the solution of the equations. In fact, a third equation can be added to further constrain the result. We can write similar equations for Na and K:

$$X = (Na_T - Na_{spr}) / (Na_T - Na_{cold})$$

$$X = (K_T - K_{spr}) / (K_T - K_{cold})$$

also:  $Na_T / K_T = R_T$

where R is a ratio depending on temperature T, according to, for example, relationships derived by Truesdell (1976) or Fournier (1981).

Thus:

$$X = (Na_{spr} - R_T \cdot K_{spr}) / (Na_{cold} - R_T \cdot K_{cold})$$

In fact, the solution to this system of equations is very insensitive to dilution X and gives a value of T similar to that initially derived from the Na/K geothermometers.

The triangle defined by the lines for Na/K, Si and H on Figure 10a define a likely area for the original geothermal water based on information from spring 3A1. Fournier & Truesdell (1974) note that in many systems, the H curve will overestimate T as it assumes no heat loss due to steam loss or factors other than dilution. Thus, the intersection of the Na/K and Si curves is a promising indicator of T and X. In the case of spring 3A1 and Fournier's (1981) geothermometer for Na/K, a temperature of some 180°C with a ratio of some 80%:20% cold:hot water is indicated. Truesdell's (1976) geothermometer results in a larger X:T triangular window, but a temperature of 140°C and X of less than 0.6 are a possibility. Similar results (for Si and enthalpy) can be obtained by using the plot in Fig. 4.10 of Fournier (1981), reproduced here as Figure 12.

The technique is applied to waters 5A3 and 1JA3 in figures 10b and 10c. The possible X:T windows are larger than for 3A1 but yield not dissimilar results. 1JA3 suggests somewhat higher minimum temperatures than the waters from the Troll springs. This is not surprising as the Jotun springs lie a considerable distance from the Troll springs and probably represent a different hydrothermal system. They are also nearer to Sverrefjell volcano.

In summary, the most plausible geothermometric techniques suggest that the Troll thermal springs are derived from a hot saline water of original temperature between 130 and 180°C, probably diluted by between 70 and 90% by cold, sub-permafrost groundwater of the type observed in K1 during its passage to the surface. Given an average geothermal gradient of 0.079°C/m, the geothermometer temperatures correspond to a depth of some 1.6 - 2.3 km.

## 7.2 Critique of Geothermometry

The use of geothermometers, especially for high CO<sub>2</sub>, cool springs, is fraught with difficulties:

- (i) the Na-Ca-K thermometer, with magnesium correction, is susceptible to interference by high-level dissolution of Mg and Ca by cool, aggressive, high-CO<sub>2</sub> water from readily soluble carbonates of the Hecla Hoek Marbles.
- (ii) the Na-K thermometer may also be susceptible to interference from the same source (i.e. dissolution of Na and K from high-level marbles). In defence of the thermometer, the concentrations of Na and K in the marbles are very low (particularly in Sv5R - see Table 4), but the ratio Na/K is much lower in the marbles than the water. A Na/K fingerprint from the marbles on the waters would thus tend to artificially lower the aqueous Na/K ratio, resulting in misleadingly high calculated temperatures.
- (iii) it may thus be suggested that the waters do not contain a high-temperature component at all, but reflect waters circulating at shallow depths with little dilution. Nevertheless, dissolved

silicon concentrations (using the chalcedony thermometer, which may be more appropriate for low T waters) suggest minimum equilibration temperatures of at least 70°C for the Troll springs and 90-130°C for the Jotun springs, even without taking into account possible dilution.

- (iv) even the chalcedony geothermometer is open to question, however, in cool, aggressive waters which are rich in CO<sub>2</sub>. Such waters may attack silicate minerals other than silica. Silicon concentrations may then accumulate in the water in concentrations greater than that reflecting quartz/chalcedony saturation (due to the slow kinetics of silica precipitation). Use of silica geothermometers in such situations will thus overestimate temperatures.

### 7.3 Speciation and Saturation Modelling

The composition of the sampled spring waters was modelled using the code MINTEQA2. Iron and manganese redox couples were activated. Selected results of the modelling are indicated in Figures 11a-c. In the following discussion, the term 'saturation' is used to apply to waters within  $\pm 0.3$  of a saturation index (SI) of zero. The following features can be noted:

- (i) The thaw water K2 is saturated with respect to calcite and quartz, but significantly undersaturated with respect to dolomite, magnesite, strontianite, microcline, albite, anorthite, fluorite and barite. It is supersaturated with respect to goethite, kaolinite and montmorillonite. It is a typical shallow groundwater from a calcitic rock, where normal subaerial carbonate and silicate weathering processes are occurring.
- (ii) The cold sub-permafrost groundwater K1 is saturated with respect to chalcedony, but supersaturated with respect to quartz, kaolinite, montmorillonite and goethite. The water is undersaturated with respect to calcite, dolomite, magnesite, strontianite, fluorite, barite, and the feldspar phases. The quartz oversaturation and calcite undersaturation are slightly unusual; the former may possibly be ascribed to mixing with a deeper, thermal water.
- (iii) The hottest thermal water 3A1 is approximately saturated with respect to a number of mineral phases including kaolinite, microcline, aragonite, barite, calcite, dolomite. It is supersaturated with respect to fluorite, chalcedony and quartz and undersaturated with respect to the feldspar phases, magnesite and strontianite. These observations confirm that 3A1 is in approximate carbonate equilibrium with respect to dolomite and calcite (comprising the bulk of the surrounding dolomitic marble bedrock). The supersaturation with respect to the silica phases confirms that the precipitation of silica is kinetically slow relative to the implied rapid ascent of the thermal waters. The other hot (> c. 20°C) waters exhibit similar saturation features to 3A1.



- (iv) The cooler thermal waters, which have presumably been subject to more dilution during their upward ascent or (in the case of the 'D' springs) through the secondary travertine tend towards a greater degree of supersaturation with respect to the carbonate and silicate minerals. The waters between 10 and 15°C tend to exhibit calcite, aragonite and dolomite supersaturation and in two cases approximate magnesite saturation. These minerals are thus all candidates for possible deposition in travertine. XRD spectra indicate however that calcite is the unique detectable carbonate constituent in travertine samples. The cooler waters also tend to become supersaturated with respect to microcline and albite. The increasing saturation with respect to carbonate minerals can be ascribed to progressive degassing of carbon dioxide as water nears the surface environment, this leads to an increase in pH affecting the saturation of carbonate phases. Phases such as quartz, chalcedony, fluorite and barite show little change in saturation status relative to the hot waters, the solubility of these minerals not being directly pH dependent.
- (v) The minerals gypsum, rhodochrosite, strontianite and siderite are undersaturated in all samples. The lack of strontianite saturation suggests that strontium is a mobile element and a useful indicator of water-rock interaction and dilution.

## 8. CONCLUSION

The Bockfjord thermal spring waters can be divided into three groups of similar ionic composition. Troll spring areas 1-3 yield waters of Na-(Ca)-HCO<sub>3</sub><sup>-</sup> composition, areas 5 and 6 waters of Na-HCO<sub>3</sub><sup>-</sup> composition and the Jotun springs waters of Na-Cl<sup>-</sup> composition. There is a negative correlation between temperature and Eh and temperature and pH. The warmest water, at 25.6°C, has an Eh of -270 mV, a pH of 6.83 and contains hydrogen sulphide. It is thus suggested that the "pristine" thermal water source has a sub-neutral pH and may be reducing (although the low Eh may alternatively result from biological activity in near-surface sediments). The thermal waters are also likely to be CO<sub>2</sub>-bearing; although no gas samples were taken, the negative relationship of pH and temperature is indicative of degassing of CO<sub>2</sub>.

Common geothermometers yield widely varying temperatures for the thermal water source. A combination of Na/K, silica and enthalpy calculations suggests temperatures at depth of 130 - 180°C for the Troll springs, with 10-30% thermal water diluted by 70-90% cold water, although these techniques are open to criticism. These temperatures correspond to depths of 1.6 - 2.3 km. If no dissolution is assumed, the chalcedony geothermometer alone indicates temperatures of around 70°C. The Jotun springs appear to be derived from a thermal water

source some few tens of °C hotter. However, each geothermometer is open to serious criticism in cool, high pCO<sub>2</sub> waters and the real circulation depth and temperature could be significantly lower. Drilling would be required to unambiguously evaluate water temperatures at depth.

Br<sup>-</sup>/Cl<sup>-</sup> ratios suggest that the water's chloride salinity is not derived from water-rock interaction but rather from fossil sea-water or, conceivably, evaporitic deposits in the Devonian sandstones. The sodium, potassium, sulphate and silicon concentrations appear to be controlled by near-equilibrium water-rock interactions at depth.

## 9. ACKNOWLEDGEMENTS

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## 10. ABBREVIATIONS USED

mg/l = milligrams per litre

Bq/l = becquerels per litre

meq/l = milliequivalents per litre

μS/cm = microsiemens per centimetre

km = kilometres

μm = micrometres

°C = degrees Centigrade

mV = millivolts

mW.m<sup>-2</sup> = milliwatts per square metre

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**Table 1. Composition of Troll and Jotun springs, Bockfjord, according to Hoel & Holtedahl (1911)**

	<b>Troll springs No. 3</b>	<b>Jotun Spring No.1</b>	<b>Jotun Spring No.2</b>
Total mineral content (mg/l)	988	2410	1456
Non-volatile mineral content (mg/l)	780	2178	1256
SO <sub>4</sub> <sup>-</sup> (mg/l)	36	104	61
Cl <sup>-</sup> (mg/l)	122	779	419
Ca (mg/l)	109	114	100
Mg (mg/l)	41	34	28

**Table 2. The chemical composition of the sampled springs.**

All units are in mg/l unless otherwise stated. Values in rounded parentheses represent values calculated from the ionic balance.

For radon determinations, a detection limit of 10 Bq/l is cited, although actual duplicate determinations (shown in square parentheses) indicate that this limit is probably too high.

Spring	T (C)	Alkalinity (meq/l)	Eh (mV)	pH	EC ( $\mu$ S/cm)	Rn (Bq/l) duplicates	F <sup>-</sup>	Cl <sup>-</sup>	Br <sup>-</sup>	NO <sub>3</sub> <sup>-</sup>	SO <sub>4</sub> <sup>=</sup>	Si	Mg	Ca	Na	K	Fe	Mn	Ba	Sr	B	Li
1A5	21.6	13.2	-220	6.80	16070		3.35	147	0.408	<0.05	52.9	23.8	33.5	111	214	13.1	0.031	0.011	0.112	1.1	0.226	0.350
2A4i	22.5	13.2	-240	6.68	14960		3.32	144	0.412	<0.05	49.2	23.4	33.0	111	212	13.0	0.025	0.014	0.113	1.1	0.228	0.350
2A4ii	13.4	13.2	-60	7.19	15810	<10 [2]	3.29	147	0.413	<0.05	49.5	21.7	33.0	104	211	13.1	0.029	0.013	0.107	1.1	0.218	0.346
3A1	25.6	13.1	-273	6.83	16300	<10 [9, 8]	3.26	150	0.386	<0.05	50.7	23.5	33.1	110	212	13.2	0.093	0.014	0.112	1.1	0.212	0.348
3Di	9.5	11.3	216	7.65	14080		3.21	135	0.365	0.18	46.4	21.4	29.7	81.3	200	12.7	<0.01	<0.001	0.097	0.96	0.210	0.324
5A3	19.4	13.5	133	6.78	19420	<10 [7, 7]	3.93	228	0.631	0.21	55.3	29.3	29.2	99.0	288	16.6	<0.01	0.017	0.105	1.1	0.331	0.456
5A4	14.5	13.6	159	7.47	18380		3.87	221	0.534	0.23	57.2	28.6	31.1	88.6	282	16.3	0.010	<0.001	0.096	1.1	0.325	0.454
5Dii	11.8	12.9	160	7.18	18300		3.92	216	0.552	0.18	56.0	28.9	29.1	84.2	276	16.2	<0.01	<0.001	0.096	1.0	0.323	0.438
6A1	20.0	13.7	70	6.69	19140		3.97	220	0.581	<0.05	56.6	28.3	28.4	94.8	283	16.0	0.063	0.013	0.110	1.1	0.311	0.443
1JA3	23.0	(10.4)	-100	6.67			4.56	958	1.73	<0.05	108	60.7	29.3	112	706	46.5	0.183	0.011	0.099	1.4	1.06	1.17
2JB	13.8	(7.3)	200	7.08			3.19	495	1.02	0.15	73.8	32.0	25.2	96.4	354	25.9	<0.01	<0.001	0.053	0.94	0.601	0.643
K1	3.6	5.4	203	7.00	16120	65, 69	0.92	29.1	<0.1	0.23	15.8	6.5	12.3	68.8	46.7	3.7	<0.01	<0.001	0.037	0.32	0.045	0.071
K2	0.2	2.2	230	7.97	2420		0.10	4.76	<0.1	<0.05	7.16	0.70	4.1	41.4	2.4	<0.5	<0.01	<0.001	0.008	0.051	<0.01	<0.005

**Table 3. Calculated temperatures (°C) by the application of common geothermometers to the sampled waters. The geothermometers used are those summarised by Fournier (1981). For the Na-K-Ca thermometer, the  $\beta=1/3$  value is used (bold type) if the  $\beta=4/3$  value is  $> 100^\circ\text{C}$ .**

<i>Spring</i>	<i>T (C)</i>	<i>Na/K</i> <i>- Fournier</i> <i>(1981)</i>	<i>Na/K</i> <i>-</i> <i>Truesdell</i> <i>(1976)</i>	<i>Quartz</i> <i>- no</i> <i>steam</i> <i>loss</i>	<i>Chalcedony</i>	<i>Na-K-Ca</i> <i>(<math>\beta=4.3</math>)</i> <i>Fournier &amp;</i> <i>Truesdell (1973)</i>	<i>Na-K-Ca</i> <i>(<math>\beta=1.3</math>)</i> <i>Fournier &amp;</i> <i>Truesdell (1973)</i>	<i>T corr for</i> <i>Mg content</i> <i>Fournier &amp;</i> <i>Potter (1979)</i>	<i>Na-K-Ca</i> <i>(corr.)</i> <i>Fournier &amp;</i> <i>Potter (1979)</i>
1A5	21.6	178	140	103	73	<b>78</b>	145	14	64
2A4i	22.5	178	140	102	72	<b>78</b>	145	13	65
2A4ii	13.4	179	141	98	68	<b>80</b>	146	18	62
3A1	25.6	179	141	102	72	<b>79</b>	146	14	64
3Di	9.5	181	143	98	67	<b>83</b>	149	29	55
5A3	19.4	174	135	113	84	<b>92</b>	148	31	61
5A4	14.5	174	135	111	82	<b>94</b>	149	41	53
5Dii	11.8	175	137	112	83	<b>95</b>	150	41	54
6A1	20.0	172	133	111	82	<b>92</b>	147	31	61
1JA3	23.0	184	147	152	127	142	<b>168</b>	104	64
2JB	13.8	192	156	117	88	112	<b>163</b>	103	60

**Table 4. X-Ray Fluorescence (XRF) analytical results for Mg, Ca, Na and Ca (as oxide %) for selected rock samples, and calculated molar and mass ratios. Analyses performed by Ulrich Siewers at the Federal Institute for Geosciences and Natural Resources (BGR), Hannover, Germany. See appendix 5.**

<i>Sample No.</i>	<i>Rock type</i>	<i>MgO</i>	<i>CaO</i>	<i>Na<sub>2</sub>O</i>	<i>K<sub>2</sub>O</i>	<i>Mg/Ca</i>	<i>Na/K</i>	<i>Mg/Ca</i>	<i>Na/K</i>
		%	%	%	%	molar	molar	mass	mass
Sv1R	Fine grained grey Hecla Hoek marble. Troll.	0.74	52.41	0.13	0.26	0.020	0.76	0.012	0.45
Sv5R	Fine grained buff Hecla Hoek marble. Troll.	17.99	29.27	0.01	0.01	0.855	1.52	0.519	0.89
Sv17R	Coarser Hecla Hoek marble. Jotun.	0.75	52.62	0.29	0.07	0.020	6.29	0.012	3.70
Sv19R	Mica schist. Hecla Hoek. Jotun.	2.13	1.79	2.37	4.71	1.655	0.76	1.004	0.45
Sv28R	Devonian Old Red Sandstone. Troll.	0.87	2.05	0.34	1.51	0.590	0.34	0.358	0.20
Sv3R	Recent travertine - Troll area 5A4	0.84	50.91	0.11	0.03	0.023	5.57	0.014	3.28
Sv10R	Fossil travertine Troll area 1	1.13	51.77	0.13	0.1	0.030	1.98	0.018	1.16
Sv11Ra	Fossil travertine Troll area 3	0.67	54.23	0.09	0.01	0.017	13.67	0.010	8.04
Sv11Rb	Fossil travertine Troll area 3	1.01	53.79	0.1	<0.01	0.026	>15.2	0.016	>8.9
Sv14R	Fossil travertine - Troll area 7	0.77	53.43	0.05	0.05	0.020	1.52	0.012	0.89
Sv15R	Recent travertine - Troll area 3A2	0.55	52.58	0.2	0.02	0.015	15.19	0.009	8.94



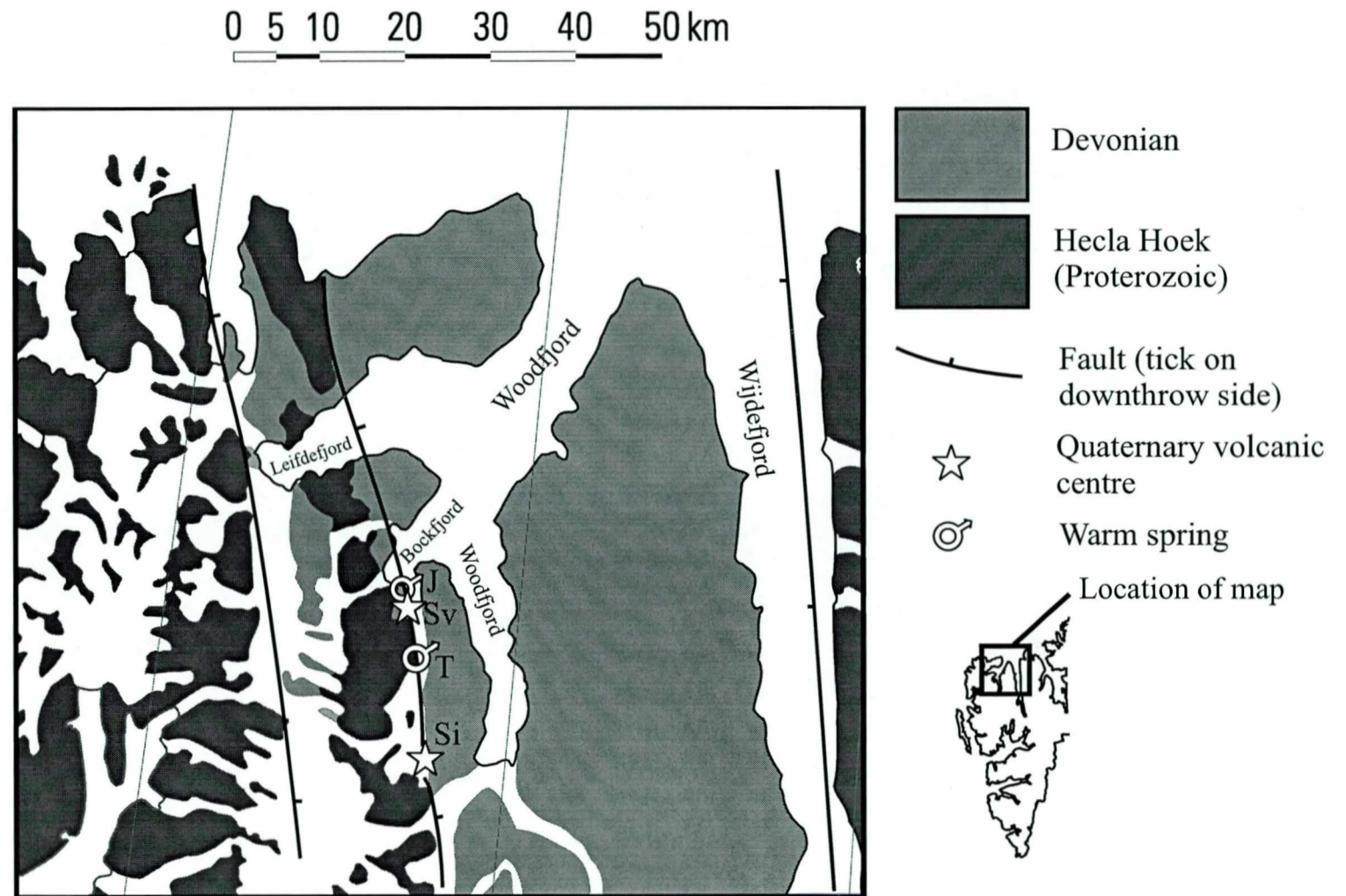


Figure 1. Sketch map of the Woodfjord area of Northern Svalbard. J = Jotun springs, Sv = Sverrefjellet, T = Troll springs, Si = Sigurdfjellet. The white areas represent glacier covered terrain or terrain covered by Quaternary (moraine / sandur) deposits. Inset shows map of Spitsbergen with location of main map.

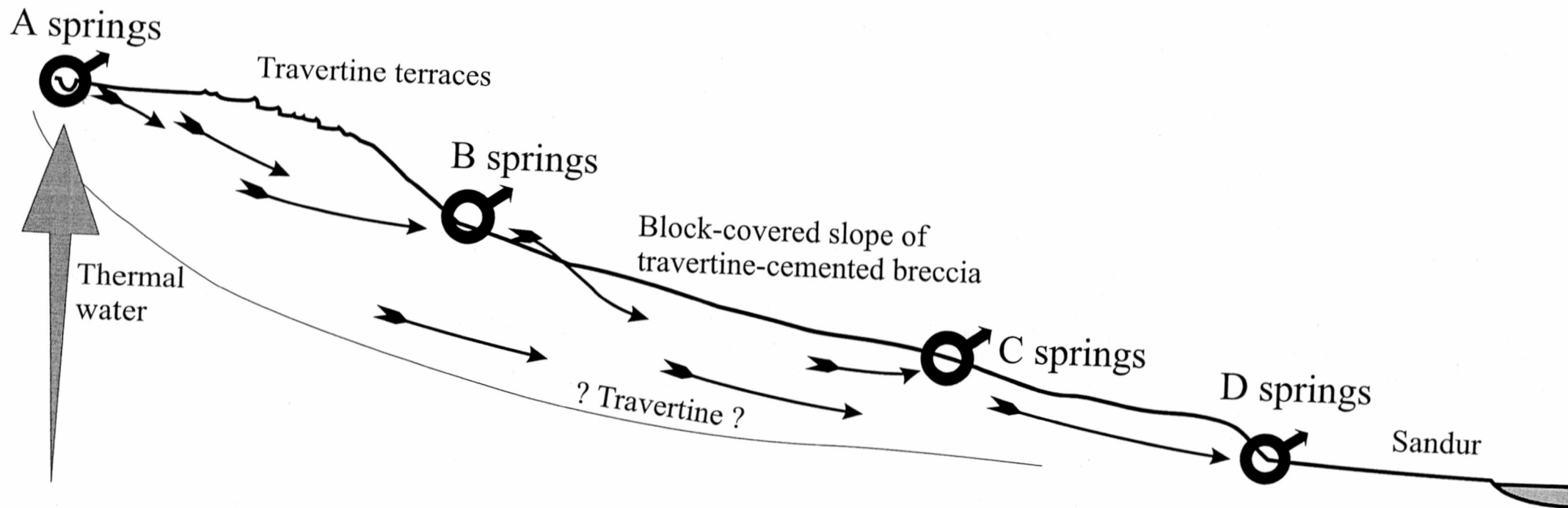


Figure 2. Schematic cross-section of the structure and probable groundwater flow at the Troll Springs site, Svalbard.



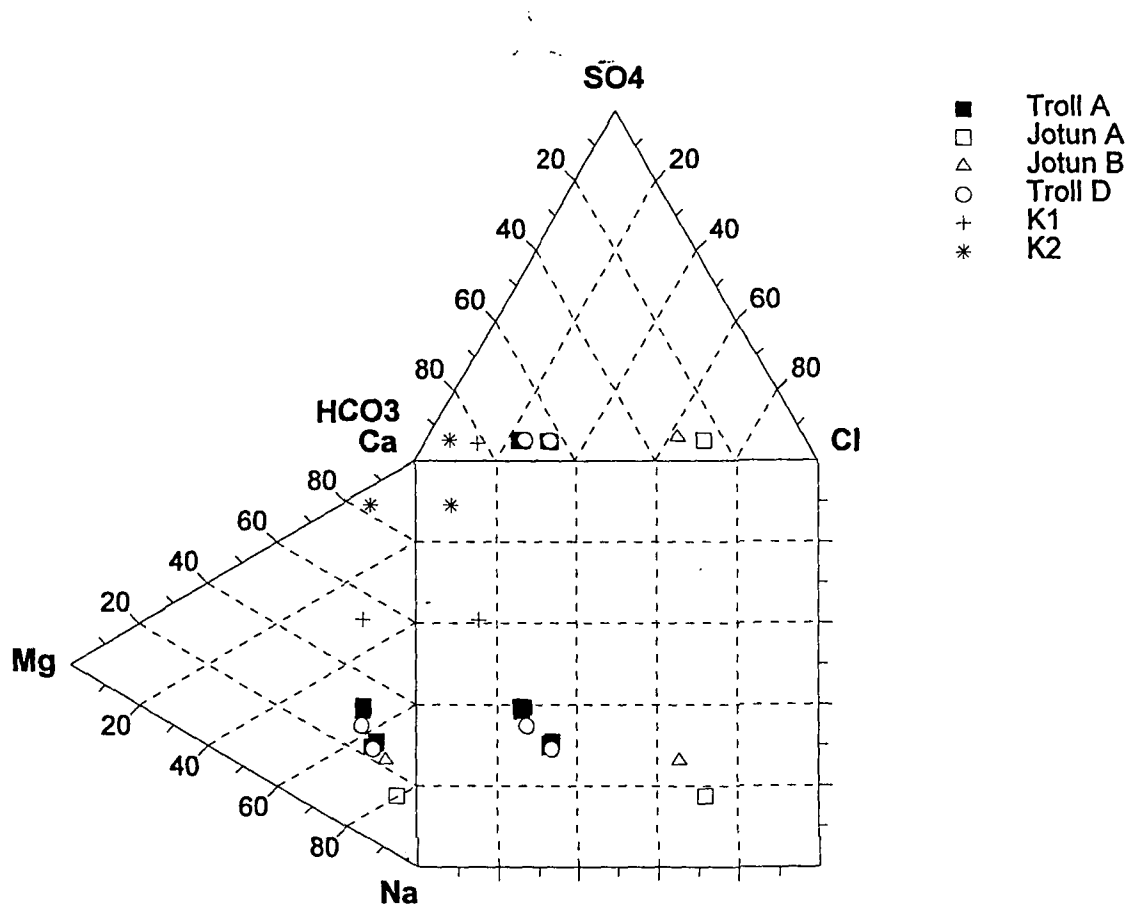


Figure 4. Durov diagram showing the major-ion composition of groundwaters (based on meq/l) from the Troll and Jotun spring areas.

Figure 5. Plots showing the relationship of temperature to Eh and pH, Bockfjord springs, Svalbard. In the box-plots of the lower two diagrams, the "box" represents the interquartile range, with the median shown by a horizontal line. The whiskers represent the extraquartile data. Box-plots are not calculable for samples of less than 5: in such cases the data points are shown as lines.

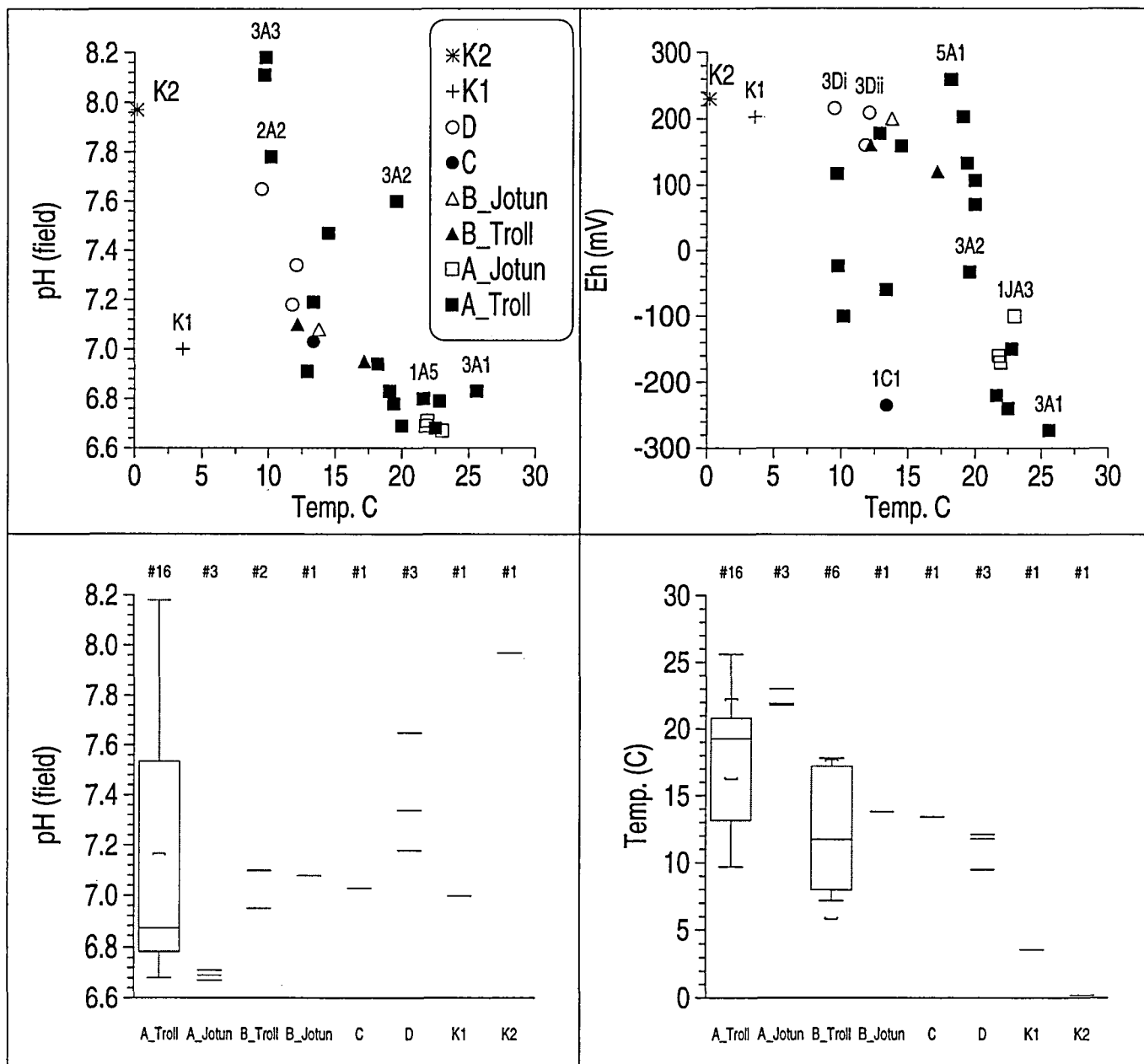


Figure 6. Plots showing the relationship of selected parameters to temperature. In cases where the parameter is below the analytical detection limit, a value of half the detection limit is plotted.

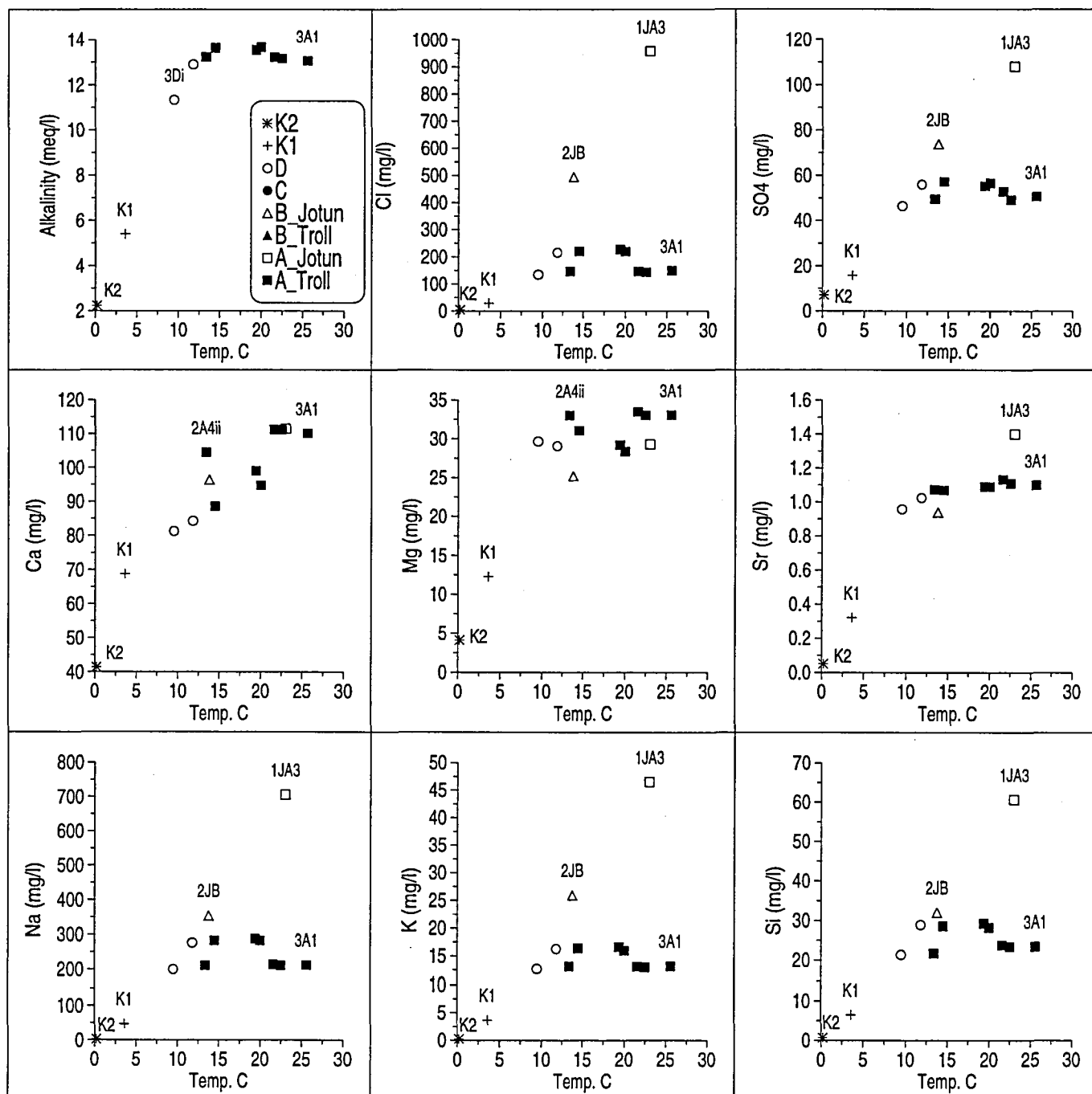


Figure 7. Plots showing the relationship of selected parameters (F<sup>-</sup>, Fe, Mn and Rn) to temperature. In cases where the parameter is below the analytical detection limit, a value of half the detection limit is plotted.

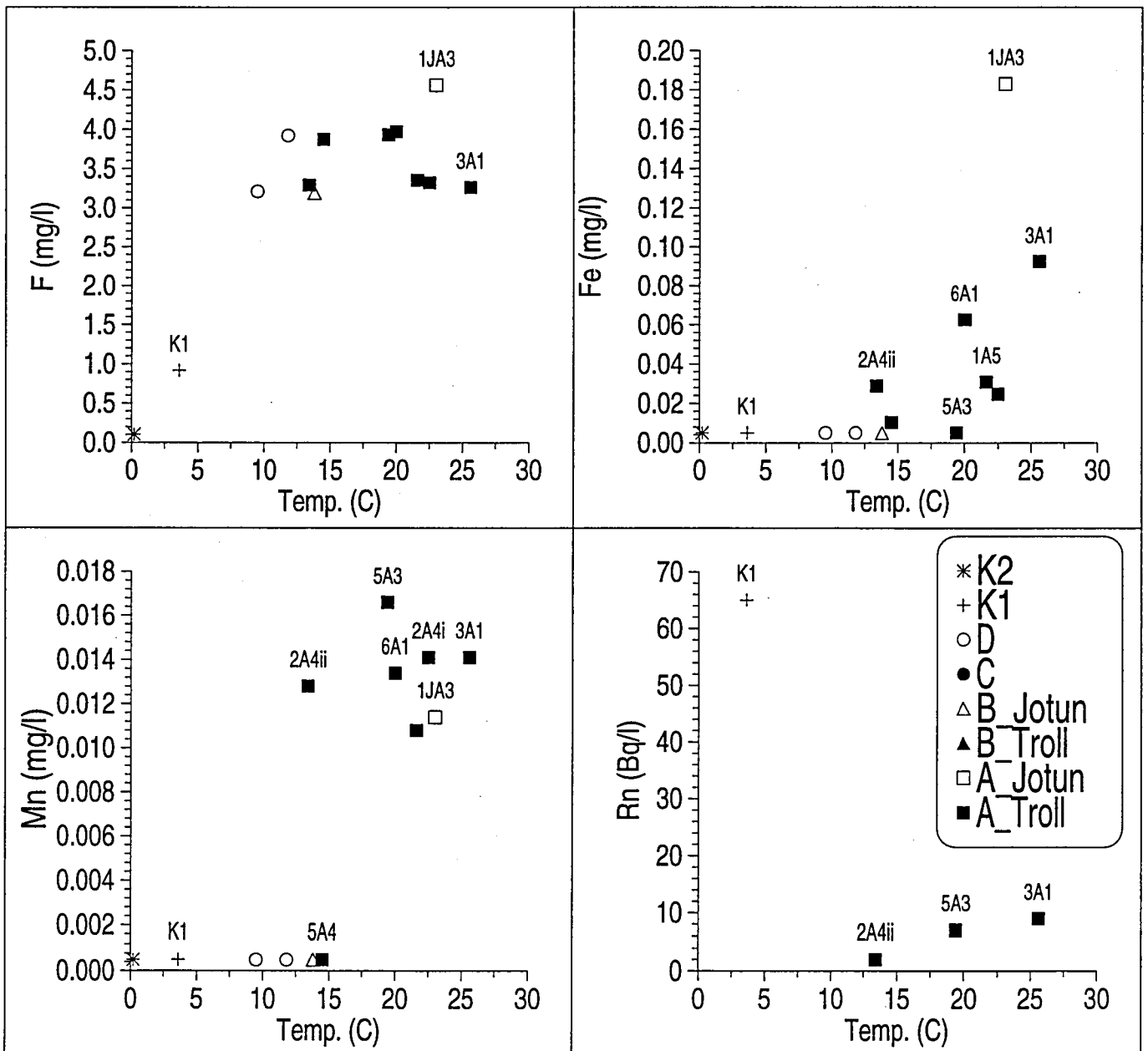


Figure 8. Plots showing the inter-relationship of selected major ions (Na, Ca, K, Cl<sup>-</sup> and SO<sub>4</sub><sup>2-</sup>). In cases where the parameter is below the analytical detection limit, a value of half the detection limit is plotted.

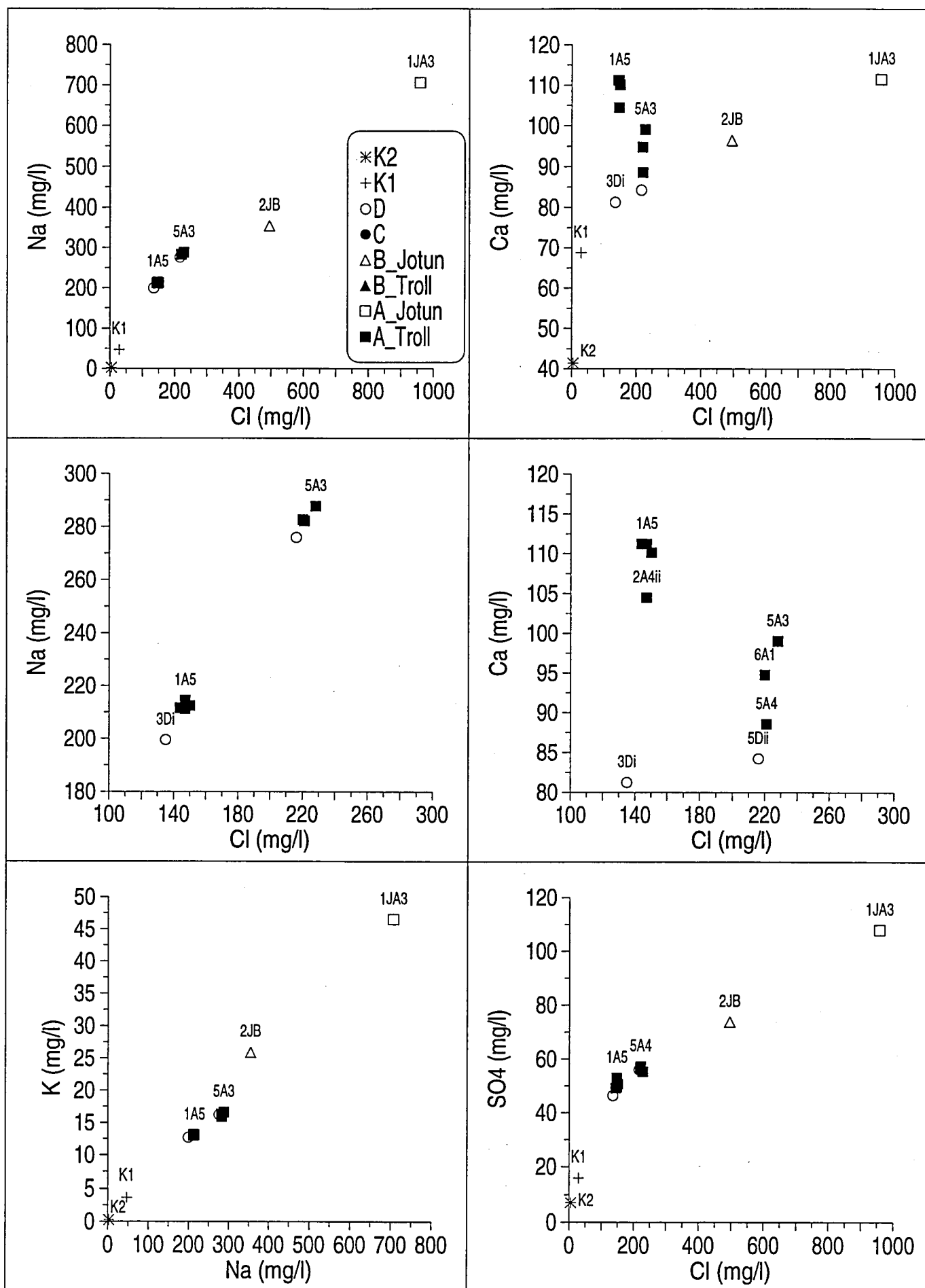
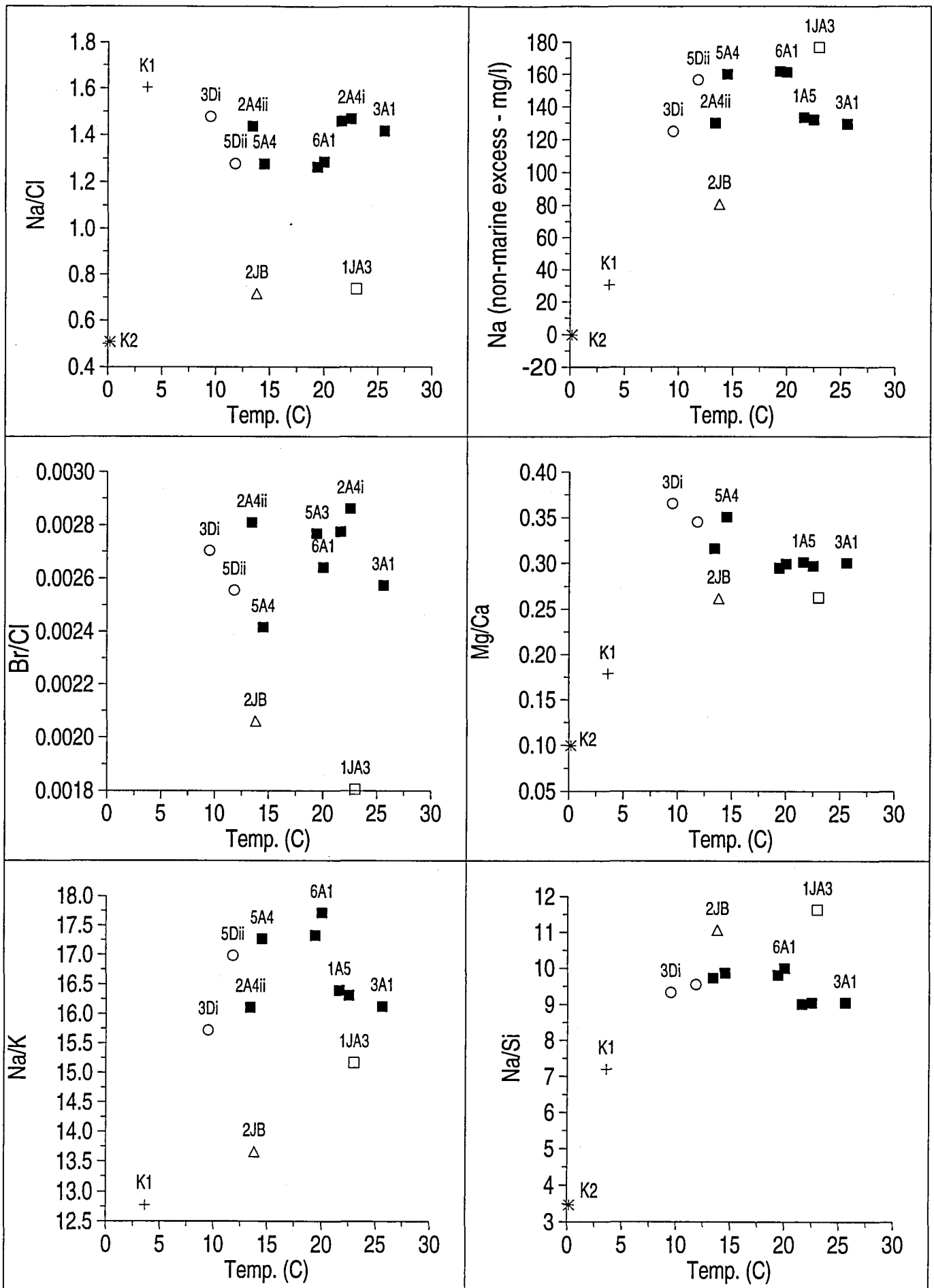
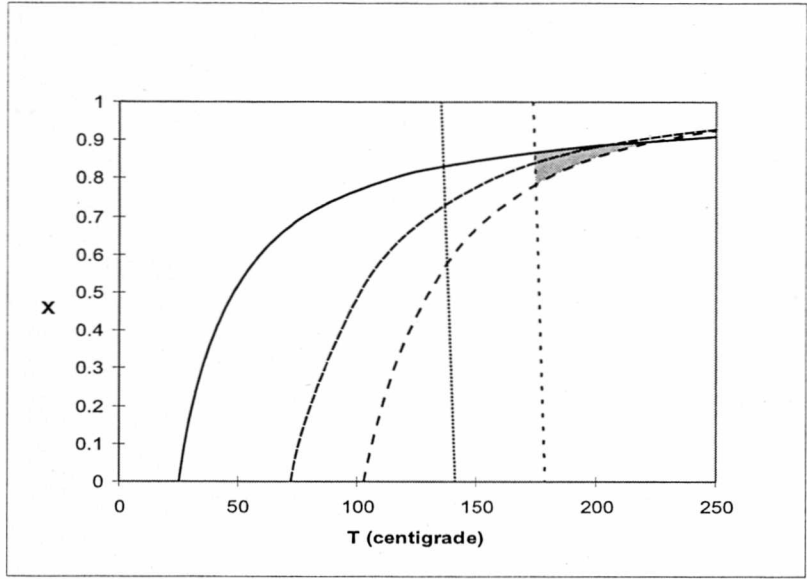


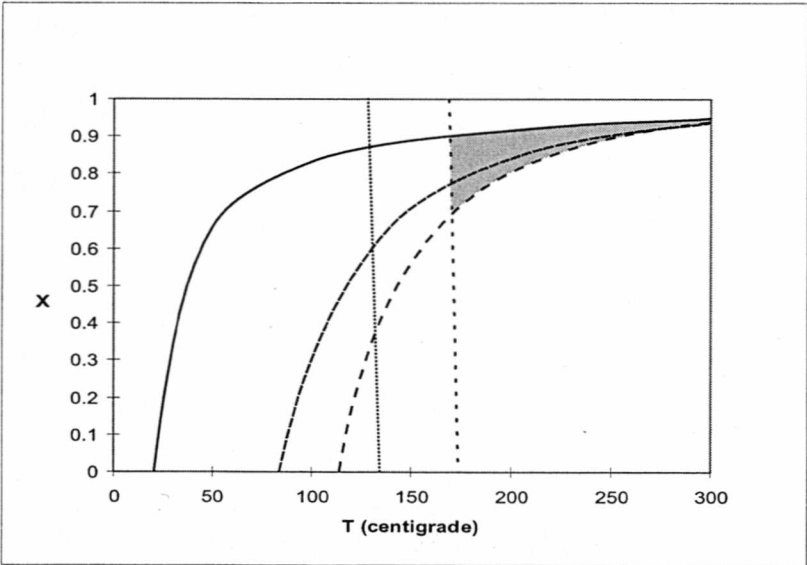


Figure 9. Plots showing the relationship of selected mass ratios and non-marine sodium excess ( $\text{Na}^*$ ) to temperature. In cases where one parameter of the ratio is below the analytical detection limit, no point is plotted.

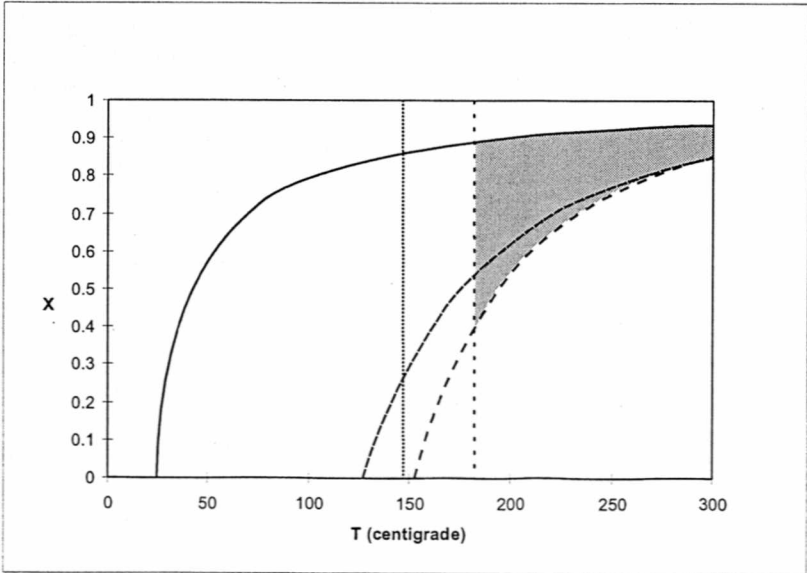




3A1



5A3

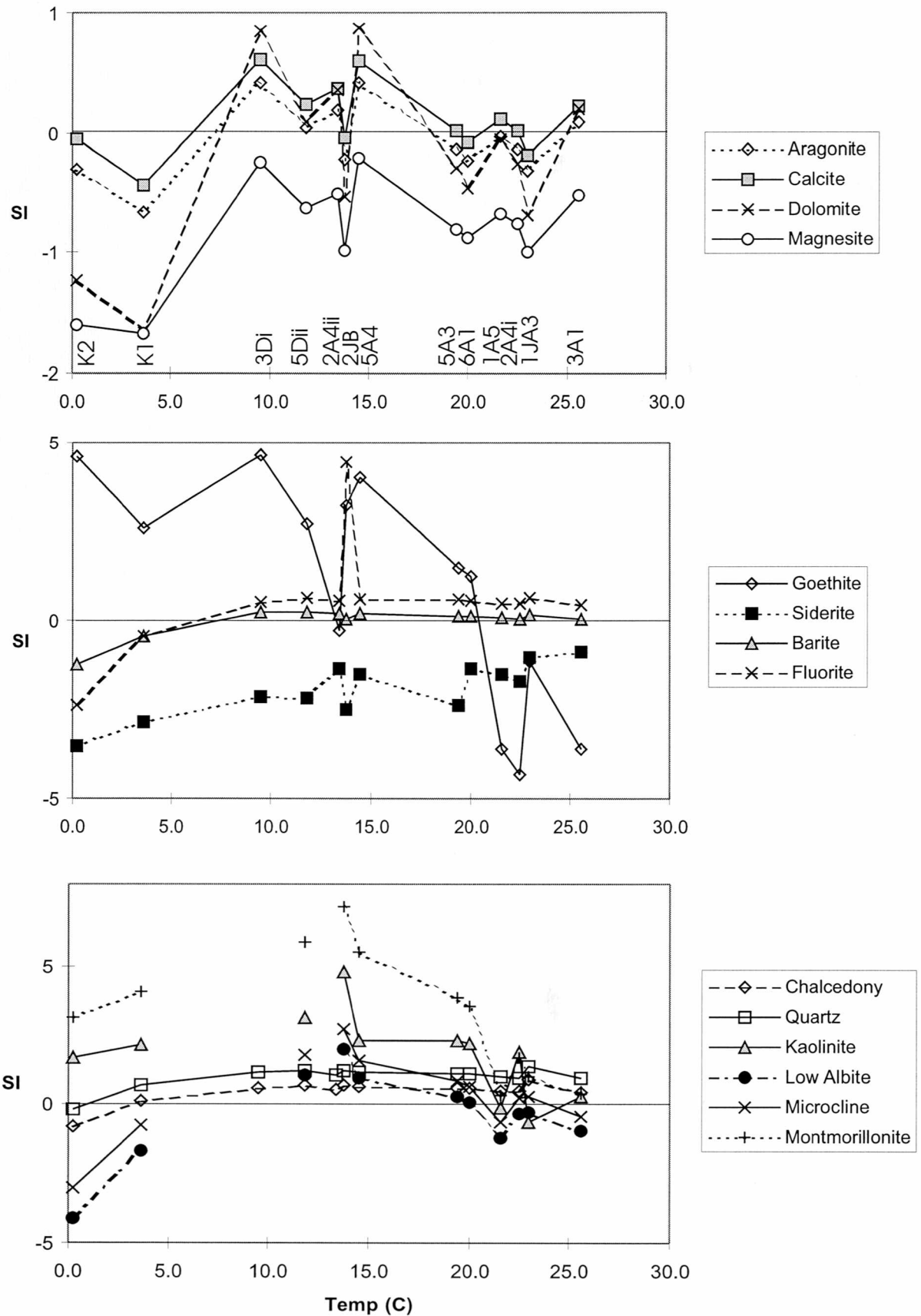


1JA3

- X(Quartz)
- - - X(Chalcedony)
- X(H)
- . - X(Na/K(Fournier))
- ..... X(Na/K(Truesdell))

Figure 10. Mixing curves plotted for X (proportion of cold water) vs. temperature T, using quartz solubility, chalcedony solubility, enthalpy and the Na/K thermometers of Fournier (1981) and Truesdell (1976). The shaded area represents the area enclosed by Fournier's (1981) Na/K thermometer, quartz solubility and the enthalpy dilution line.

Figure 11. Plots showing the relationship of various saturation indices to temperature. In the few cases where a parameter is below the analytical detection limit, a value of half the detection limit has been used as input to MINTEQA2. In several samples where Al is below detection limit for ICP-AES, values derived from ICP-MS analyses carried out at BGR, Hannover, were used.



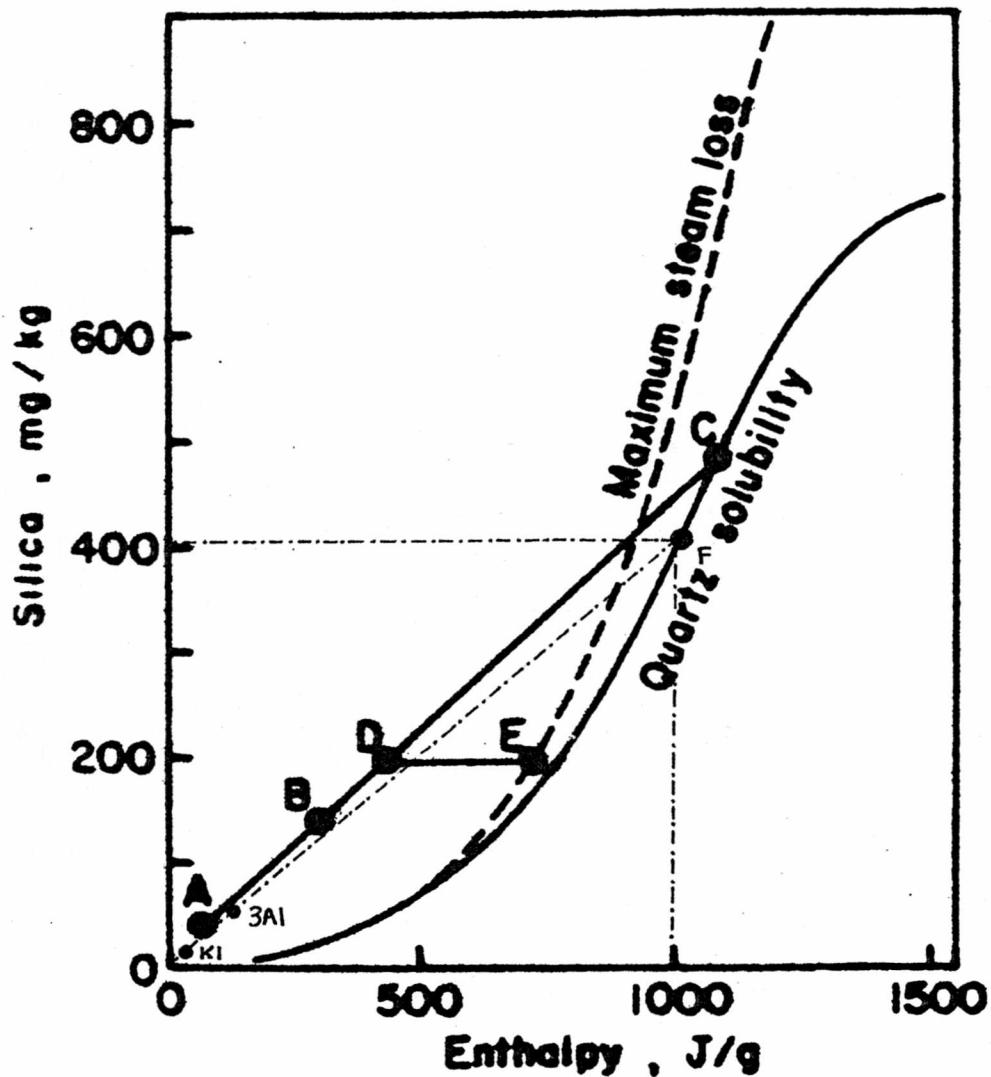


Figure 12. Estimation of temperature of 'original' thermal water using Fournier's (1981) silica-enthalpy diagram. A line drawn through K1 (Si = 6.5 mg/l, T = 3.6 °C; i.e. SiO<sub>2</sub> = 13.9 mg/l, H = 15.1 J/g) and 3A1 (Si = 23.5 mg/l, T = 25.6 °C; i.e. SiO<sub>2</sub> = 50.3 mg/l, H = 107.2 J/g), intersects the quartz solubility curve at point F; i.e. H = 1000 J/g (T=232°C) and SiO<sub>2</sub> = 403 mg/l (Si = 188 mg/l). The latter implies a dilution factor (proportion of cold water) of around 0.9. This is similar to the results obtained for SiO<sub>2</sub> and enthalpy in Figure 10. The results for Svalbard are shown as a dotted line.

## **Appendix 1**

### **Analytical Results from NGU**

**Ion chromatography (filtered [0.45 $\mu$ m], unacidified water samples)**

**ICP-AES (filtered [0.45 $\mu$ m], acidified water samples)**

ANALYSEKONTRAKT NR.: 1996.0121  
NGU PROSJEKT NR.: 2717.00

OPPDRAGSGIVER: NGU, Varme kilder Svalbard

ADRESSE:

TLF.: 310

KONTAKTPERSON: David Banks

PRØVETYPE: Vann

ANTALL PRØVER: 12

IDENTIFIKASJON AV PRØVER: Iflg. liste fra oppdragsgiver

PRØVER MOTTATT: 30.07.96

ANMERKNINGER: Ingen

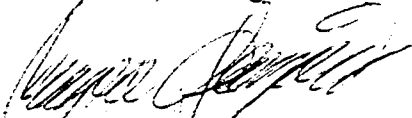
SPESIFIKASJON AV OPPDRAGET I HENHOLD TIL ANALYSEKONTRAKT:

METODE	DOKUMENTASJON *)	OMFATTES AV AKKREDITERING
ICP-AES vann	NGU-SD 3.1	Ja
IC	NGU-SD 3.4	Ja

Denne rapporten inneholder i alt 7 sider. Rapporten må ikke gjengis i utdrag uten skriftlig godkjenning fra NGU-Lab.

Alle forhold ved prøvetaking, behandling og transport av prøvene før innlevering til NGU-Lab er underlagt oppdragsgivers ansvar. Analyseresultater framlagt i denne rapporten refererer derfor kun til det prøvematerialet som er mottatt av NGU-Lab.

Trondheim, 28. august 1996



Magne Ødegård (e.f.)

\*) Fortegnelse over dokumentasjon finnes i NGU-Labs Kvalitetshåndbok, NGU-SD 0.1, som kan rekvireres fra NGU-Labs sekretariat.

**INSTRUMENT TYPE :** Thermo Jarrell Ash ICP 61

**NEDRE BESTEMMELSESGRENSER VANNANALYSER**

(For vannprøver som tyndes, blir deteksjonsgrensene automatisk omregnet).

Si ppb	Al ppb	Fe ppb	Ti ppb	Mg ppb	Ca ppb	Na ppb	K ppb	Mn ppb	P ppb
20.-	20.-	10.-	5.-	50.-	20.-	50.-	500.-	1.-	100.-
Cu ppb	Zn ppb	Pb ppb	Ni ppb	Co ppb	V ppb	Mo ppb	Cd ppb	Cr ppb	Ba ppb
5.-	2.-	50.-	20.-	10.-	5.-	10.-	5.-	10.-	2.-
Sr ppb	Zr ppb	Ag ppb	B ppb	Be ppb	Li ppb	Sc ppb	Ce ppb	La ppb	Y ppb
1.-	5.-	10.-	10.-	1.-	5.0	1.-	50.-	10.-	1.-

**ANALYSEUSIKKERHET:** ± 20 rel. % for K, Pb, Cd, Li, Ce.  
± 10 rel. % for Si, Al, Na, Mo, Cr, Zr, Ag, B og La.  
± 5 rel. % for Fe, Ti, Mg, Ca, Mn, P, Cu, Zn, Ni, Co, V, Ba, Sr, Be, Sc, Y.

**PREISJON :** Det kjøres rutinemessig kontrollprøver, som føres i kontrollidiagram (X-diagram). Disse kan forevises om ønskelig.

**ANTALL PRØVER:** 12

**ANMERKNINGER:** Verdi for Na i prøve 1JA3F er bestemt separat på fortynnet (1:9) prøve, ellers ingen.

Rapporten må ikke gjengis i utdrag uten skriftlig godkjenning fra NGU-Lab.

Ferdig analysert	06.08.96	Brit Inger Vongraven
	Dato	OPERATØR

	Sv1A5F	Sv2A41	Sv2A42	Sv3A1F	Sv3DiF	Sv5A3F	Sv5A4F	Sv5Dii	SvK1F	SvK2F
Si	23.8ppm	23.4ppm	21.7ppm	23.5ppm	21.4ppm	29.3ppm	28.6ppm	28.9ppm	6.5ppm	697ppb
Al	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb
Fe	30.8ppb	24.7ppb	28.8ppb	92.5ppb	<10.0ppb	<10.0ppb	10.3ppb	<10.0ppb	<10.0ppb	<10.0ppb
Ti	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb
Mg	33.5ppm	33.0ppm	33.0ppm	33.1ppm	29.7ppm	29.2ppm	31.1ppm	29.1ppm	12.3ppm	4.1ppm
Ca	111ppm	111ppm	104ppm	110ppm	81.3ppm	99.0ppm	88.6ppm	84.2ppm	68.8ppm	41.4ppm
Na	214ppm	212ppm	211ppm	212ppm	200ppm	288ppm	282ppm	276ppm	46.7ppm	2.4ppm
K	13.1ppm	13.0ppm	13.1ppm	13.2ppm	12.7ppm	16.6ppm	16.3ppm	16.2ppm	3.7ppm	< 500ppb
Mn	10.8ppb	14.1ppb	12.8ppb	14.1ppb	<1.00ppb	16.6ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb
P	< 100ppb	< 100ppb	< 100ppb	< 100ppb	< 100ppb	< 100ppb	< 100ppb	< 100ppb	< 100ppb	< 100ppb
Cu	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	9.7ppb
Zn	< 2.0ppb	< 2.0ppb	< 2.0ppb	< 2.0ppb	< 2.0ppb	< 2.0ppb	< 2.0ppb	< 2.0ppb	< 2.0ppb	5.0ppb
Pb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb
Ni	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb	<20.0ppb
Co	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb
V	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb
Mo	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb
Cd	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb
Cr	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb
Ba	112ppb	113ppb	107ppb	112ppb	96.6ppb	105ppb	96.3ppb	95.6ppb	36.5ppb	8.2ppb
Sr	1.1ppm	1.1ppm	1.1ppm	1.1ppm	957ppb	1.1ppm	1.1ppm	1.0ppm	321ppb	51.1ppb
Zr	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb	< 5.0ppb
Ag	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb
B	226ppb	228ppb	218ppb	212ppb	210ppb	331ppb	325ppb	323ppb	44.7ppb	<10.0ppb
Be	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb
Li	350ppb	350ppb	346ppb	348ppb	324ppb	456ppb	454ppb	438ppb	71.2ppb	< 5.0ppb
Sc	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb
Ce	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb	<50.0ppb
La	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb	<10.0ppb
Y	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb	<1.00ppb



	1JA3F	2JB1F
Si	60.7ppm	32.0ppm
Al	<20.0ppb	<20.0ppb
Fe	183ppb	<10.0ppb
Ti	< 5.0ppb	< 5.0ppb
Mg	29.3ppm	25.2ppm
Ca	112ppm	96.4ppm
Na	706ppm	354ppm
K	46.5ppm	25.9ppm
Mn	11.4ppb	<1.00ppb
P	< 100ppb	< 100ppb
Cu	< 5.0ppb	< 5.0ppb
Zn	< 2.0ppb	< 2.0ppb
Pb	<50.0ppb	<50.0ppb
Ni	<20.0ppb	<20.0ppb
Co	<10.0ppb	<10.0ppb
V	< 5.0ppb	< 5.0ppb
Mo	<10.0ppb	<10.0ppb
Cd	< 5.0ppb	< 5.0ppb
Cr	<10.0ppb	<10.0ppb
Ba	99.1ppb	53.0ppb
Sr	1.4ppm	939ppb
Zr	< 5.0ppb	< 5.0ppb
Ag	<10.0ppb	<10.0ppb
B	1.1ppm	601ppb
Be	<1.00ppb	<1.00ppb
Li	1.2ppm	643ppb
Sc	<1.00ppb	<1.00ppb
Ce	<50.0ppb	<50.0ppb
La	<10.0ppb	<10.0ppb
Y	<1.00ppb	<1.00ppb

7 ANIONER : F<sup>-</sup>, Cl<sup>-</sup>, NO<sub>2</sub><sup>-</sup>, Br<sup>-</sup>, NO<sub>3</sub><sup>-</sup>, PO<sub>4</sub><sup>3-</sup>, SO<sub>4</sub><sup>2-</sup>

INSTRUMENT TYPE : DIONEX IONEKROMATOGRAF 2120i

**NEDRE BESTEMMELSESGRENSER**

ION	F <sup>-</sup>	Cl <sup>-</sup>	NO <sub>2</sub> <sup>-*</sup>	Br <sup>-</sup>	NO <sub>3</sub> <sup>-</sup>	PO <sub>4</sub> <sup>3-</sup>	SO <sub>4</sub> <sup>2-</sup>
Nedre bestemmelsesgrense - mg/l	0.05	0.1	0.05	0.1	0.05	0.2	0.1

ANALYSEUSIKKERHET : 10 % rel. for alle ionene

PREISJON : Det kjøres rutinemessig kontrollprøver, som føres i kontrolldiagram (X-diagram). Disse kan forevises om ønskelig.

ANTALL PRØVER: 2

ANMERKNINGER:

\* NGU-LAB er ikke akkreditert for NO<sub>2</sub><sup>-</sup> \*

Rapporten må ikke gjengis i utdrag uten skriftlig godkjenning fra NGU-Lab.

Ferdig analysert	28. august 1996	Egil Kvam
	Dato	OPERATØR

Prøve Id.	F <sup>-</sup> [mg/l]	Cl <sup>-</sup> [mg/l]	NO <sub>2</sub> <sup>-</sup> [mg/l]	Br <sup>-</sup> [mg/l]	NO <sub>3</sub> <sup>-</sup> [mg/l]	PO <sub>4</sub> <sup>3-</sup> [mg/l]	SO <sub>4</sub> <sup>2-</sup> [mg/l]
1JA3F	4.56	958	< 0.1	1.73	< 0.05	< 0.2	108
2JB1F	3.19	495	< 0.1	1.02	0.145	< 0.2	73.8

ANALYSEKONTRAKT NR.: 1996.0188  
NGU PROSJEKT NR.: 2717.00

OPPDRAKSGIVER: NGU, Varme kilder på Svalbard

ADRESSE:

TLF.: 310

KONTAKTPERSON: David Banks

PRØVETYPE: Vann

ANTALL PRØVER: 1

IDENTIFIKASJON AV PRØVER: Iflg. liste fra oppdragsgiver

PRØVER MOTTATT: 13.09.96

ANMERKNINGER: Ingen

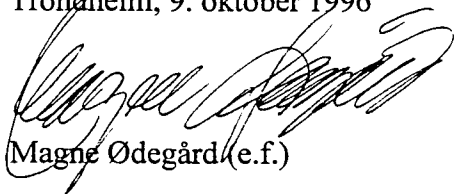
**SPESIFIKASJON AV OPPDRAGET I HENHOLD TIL ANALYSEKONTRAKT:**

METODE	DOKUMENTASJON *)	OMFATTES AV AKKREDITERING
ICP-AES vann	NGU-SD 3.1	Ja
IC	NGU-SD 3.4	Ja

Denne rapporten inneholder i alt 6 sider. Rapporten må ikke gjengis i utdrag uten skriftlig godkjenning fra NGU-Lab.

Alle forhold ved prøvetaking, behandling og transport av prøvene før innlevering til NGU-Lab er underlagt oppdragsgivers ansvar. Analyseresultater framlagt i denne rapporten refererer derfor kun til det prøvematerialet som er mottatt av NGU-Lab.

Trondheim, 9. oktober 1996



Magne Ødegård (e.f.)

\*) Fortegnelse over dokumentasjon finnes i NGU-Labs Kvalitetshåndbok, NGU-SD 0.1, som kan rekvireres fra NGU-Labs sekretariat.

**INSTRUMENT TYPE :**

**Thermo Jarrell Ash ICP 61**

**NEDRE BESTEMMELSESGRENSER VANNANALYSER**

(For vannprøver som tynnes, blir deteksjonsgrensene automatisk omregnet).

Si ppb	Al ppb	Fe ppb	Ti ppb	Mg ppb	Ca ppb	Na ppb	K ppb	Mn ppb	P ppb
20.-	20.-	10.-	5.-	50.-	20.-	50.-	500.-	1.-	100.-
Cu ppb	Zn ppb	Pb ppb	Ni ppb	Co ppb	V ppb	Mo ppb	Cd ppb	Cr ppb	Ba ppb
5.-	2.-	50.-	20.-	10.-	5.-	10.-	5.-	10.-	2.-
Sr ppb	Zr ppb	Ag ppb	B ppb	Be ppb	Li ppb	Sc ppb	Ce ppb	La ppb	Y ppb
1.-	5.-	10.-	10.-	1.-	5.0	1.-	50.-	10.-	1.-

**ANALYSEUSIKKERHET:** ± 20 rel. % for K, Pb, Cd, Li, Ce.  
± 10 rel. % for Si, Al, Na, Mo, Cr, Zr, Ag, B og La.  
± 5 rel. % for Fe, Ti, Mg, Ca, Mn, P, Cu, Zn, Ni, Co, V, Ba, Sr, Be, Sc, Y.

**PRESISJON :** Det kjøres rutinemessig kontrollprøver, som føres i kontrolldiagram (X-diagram). Disse kan forevises om ønskelig.

**ANTALL PRØVER: 1**

**ANMERKNINGER:** Be rapporteres ikke p.g.a. at Be-kanalen er ustabil

**Rapporten må ikke gjengis i utdrag uten skriftlig godkjenning fra NGU-Lab.**

Ferdig analysert	19. september 1996	Brit Inger Vongraven
	Dato	OPERATØR

Sv6A1F

Si	28.3ppm
Al	<20.0ppb
Fe	62.6ppb
Ti	< 5.0ppb
Mg	28.4ppm
Ca	94.8ppm
Na	283ppm
K	16.0ppm
Mn	13.4ppb
P	< 100ppb
Cu	< 5.0ppb
Zn	< 2.0ppb
Pb	<50.0ppb
Ni	<20.0ppb
Co	<10.0ppb
V	< 5.0ppb
Mo	<10.0ppb
Cd	< 5.0ppb
Cr	<10.0ppb
Ba	110ppb
Sr	1.1ppm
Zr	< 5.0ppb
Ag	<10.0ppb
B	311ppb
Li	443ppb
Sc	<1.00ppb
Ce	<50.0ppb
La	<10.0ppb
Y	<1.00ppb

7 ANIONER : F<sup>-</sup>, Cl<sup>-</sup>, NO<sub>2</sub><sup>-</sup>, Br<sup>-</sup>, NO<sub>3</sub><sup>-</sup>, PO<sub>4</sub><sup>3-</sup>, SO<sub>4</sub><sup>2-</sup>

INSTRUMENT TYPE : DIONEX IONEKROMATOGRAF 2120i

**NEDRE BESTEMMELSESGRENSER**

ION	F <sup>-</sup>	Cl <sup>-</sup>	NO <sub>2</sub> <sup>-*</sup>	Br <sup>-</sup>	NO <sub>3</sub> <sup>-</sup>	PO <sub>4</sub> <sup>3-</sup>	SO <sub>4</sub> <sup>2-</sup>
Nedre bestemmelsesgrense - mg/l	0.05	0.1	0.05	0.1	0.05	0.2	0.1

ANALYSEUSIKKERHET : 10 % rel. for alle ionene

5 PRECISION : Det kjøres rutinemessig kontrollprøver, som føres i kontrolldiagram (X-diagram). Disse kan forevises om ønskelig.

ANTALL PRØVER: 1

ANMERKNINGER:

\* NGU-LAB er ikke akkreditert for NO<sub>2</sub><sup>-</sup> \*

Rapporten må ikke gjengis i utdrag uten skriftlig godkjenning fra NGU-Lab.

Ferdig analysert	7. oktober 1996	Egil Kvam
	Dato	OPERATØR

Prøve Id.	F <sup>-</sup> [mg/l]	Cl <sup>-</sup> [mg/l]	NO <sub>2</sub> <sup>-</sup> [mg/l]	Br <sup>-</sup> [mg/l]	NO <sub>3</sub> <sup>-</sup> [mg/l]	PO <sub>4</sub> <sup>3-</sup> [mg/l]	SO <sub>4</sub> <sup>2-</sup> [mg/l]
188/96 - Sv6a1F	3.58	282	< 0.05	0.778	< 0.05	< 0.2	61.9



ANALYSEKONTRAKT NR.: 1996.0004  
 NGU PROSJEKT NR.: 2717.00

OPPDRAGSGIVER: NGU, Varme Kilder - Svalbard

ADRESSE:

TLF.:

KONTAKTPERSON: David Banks

PRØVETYPE: Vann

ANTALL PRØVER: 11

IDENTIFIKASJON AV PRØVER: Iflg. liste fra oppdragsgiver

PRØVER MOTTATT: 08.01.97

ANMERKNINGER: Ingen

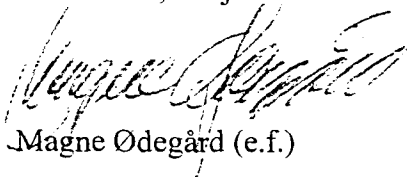
SPESIFIKASJON AV OPPDRAGET I HENHOLD TIL ANALYSEKONTRAKT:

METODE	DOKUMENTASJON *)	OMFATTES AV AKKREDITERING
IC	NGU-SD 3.4	Ja

Denne rapporten inneholder i alt 4 sider. Rapporten må ikke gjengis i utdrag uten skriftlig godkjenning fra NGU-Lab.

Alle forhold ved prøvetaking, behandling og transport av prøvene før innlevering til NGU-Lab er underlagt oppdragsgivers ansvar. Analyseresultater framlagt i denne rapporten refererer derfor kun til det prøvematerialet som er mottatt av NGU-Lab.

Trondheim, 13. januar 1997


 Magne Ødegård (e.f.)

\*) Fortegnelse over dokumentasjon finnes i NGU-Labs Kvalitetshåndbok, NGU-SD 0.1, som kan rekvireres fra NGU-Labs sekretariat.

7 ANIONER : F, Cl, NO<sub>2</sub><sup>-</sup>, Br, NO<sub>3</sub><sup>-</sup>, PO<sub>4</sub><sup>3-</sup>, SO<sub>4</sub><sup>2-</sup>

INSTRUMENT TYPE : DIONEX IONEKROMATOGRAF 2120i

NEDRE BESTEMMELSESGRENSER

ION	F <sup>-</sup>	Cl <sup>-</sup>	NO <sub>2</sub> <sup>-*</sup>	Br <sup>-</sup>	NO <sub>3</sub> <sup>-</sup>	PO <sub>4</sub> <sup>3-</sup>	SO <sub>4</sub> <sup>2-</sup>
Nedre bestemmelsesgrense - mg/l	0.05	0.1	0.05	0.1	0.05	0.2	0.1

ANALYSEUSIKKERHET : 10 % rel. for alle ionene

PRESISJON : Det kjøres rutinemessig kontrollprøver, som føres i kontrolldiagram (X-diagram). Disse kan forevises om ønskelig.

ANTALL PRØVER: 11

ANMERKNINGER: Ingen

\* NGU-LAB er ikke akkreditert for NO<sub>2</sub><sup>-</sup> \*

Rapporten må ikke gjengis i utdrag uten skriftlig godkjenning fra NGU-Lab.

Ferdig analysert	9. januar 1997	Egil Kvam
	Dato	OPERATØR

Prøve Id.	F <sup>-</sup> [mg/l]	Cl <sup>-</sup> [mg/l]	NO <sub>2</sub> <sup>-</sup> [mg/l]	Br <sup>-</sup> [mg/l]	NO <sub>3</sub> <sup>-</sup> [mg/l]	PO <sub>4</sub> <sup>3-</sup> [mg/l]	SO <sub>4</sub> <sup>2-</sup> [mg/l]
4/97 - 1	3.87	221	< 0.05	0.534	0.229	< 0.2	57.2
4/97 - 2	3.32	144	< 0.05	0.412	< 0.05	< 0.2	49.2
4/97 - 3	3.29	147	< 0.05	0.413	< 0.05	< 0.2	49.5
4/97 - 4	3.26	150	< 0.05	0.386	< 0.05	< 0.2	50.7
4/97 - 5	3.21	135	< 0.05	0.365	0.180	< 0.2	46.4
4/97 - 6	3.35	147	< 0.05	0.408	< 0.05	< 0.2	52.9
4/97 - 7	3.97	220	< 0.05	0.581	< 0.05	< 0.2	56.6
4/97 - 8	3.93	228	< 0.05	0.631	0.212	< 0.2	55.3
4/97 - 9	3.92	216	< 0.05	0.552	0.183	< 0.2	56.0
4/97 - 10	0.917	29.1	< 0.05	< 0.1	0.226	< 0.2	15.8
4/97 - 11	0.103	4.76	< 0.05	< 0.1	< 0.05	< 0.2	7.16

ID numbers as in Appendix 2

## **Appendix 2**

**Analytical results - anions-from Norges Landbrukshøyskole  
(filtered [0.45µm], unacidified water samples)**

**also laboratory determinations of  
pH, electrical conductivity and alkalinity at NLH**

Prøvemerkning

1: Sv5A4

2: Sv2A4i

3: Sv2A4ii

4: Sv3A1

5: Sv3Di

6: Sv1A5

7: Sv6A1

8: Sv5A3

9: Sv5Dii

10: SvK1

11: SvK2

LABORATORIUM FOR ANALYTISK KJEMI (LAK-KA) - NLH

Boks 5031

1432 ÅS

DATO:961205

ANALYSEBEVIS for rekvisisjon nr: 960430

JORD- OG VANNFAG (B Swensen)

Boks 5028

0 HER

Kontaktperson: BERIT SWENSEN

11 vann

Rekvisisjonens registreringsdato: 960912

Laboratoriet oppbevarer prøvene i en måned etter at oppdraget er fullført

Analyseresultatet er oppgitt med inntil tre gjeldende siffer

Merknader:

Journr	Mrk	Nitrat-N	Vannl-P	Sulfat-S	Vannl-Cl
438980	1	<0.1 mg/l	<0.1 mg/l	26. mg/l	210. mg/l
438981	2	<0.1	<0.1	22	140
438982	3	<0.1	<0.1	22	145
438983	4	<0.1	<0.1	21	145
438984	5	<0.1	<0.1	20	140
438985	6	<0.1	<0.1	21	150
438986	7	<0.1	<0.1	24	210
438987	8	<0.1	<0.1	25	215
438988	9	<0.1	<0.1	24	210
438989	10	<0.1	<0.1	7.6	34
438990	11	<0.1	<0.1	4.6	5.0

Frank Antons

J.nr. 210/96

Mottatt: 12.8 96

Rekvirent: B. Swensen

Institutt/Seksjon: .....

Ferdig: .....

Adresse: .....

Prove står på kjøleskap hos idrettskjelleren.

L.nr.	Pr.mrk.	Analyser som ønskes utført. (Kryss av på venstre side i rubrikken)										
		pH	ledn.	alk.								
1	Sv5A4											
2	Sv2A4i											
3	Sv2A4ii											
4	Sv3A1											
5	Sv3Di											
6	Sv1A5											
7	Sv6A1											
8	Sv5A3											
9	SvSDii											
10	SvK1											
11	SvK2											

Ark1

Berit Swensen				
Rekv. 210/96				
		pH	ms.	meq/l
		ledn.evne	alk.	
Pr.nr.	1	7.76	1.838	11.75
	2	7.86	1.496	10.75
	3	7.79	1.581	10.90
	4	7.45	1.630	11.75
	5	8.04	1.408	9.75
	6	7.54	1.607	11.75
	7	7.50	1.914	12.75
	8	7.32	1.942	12.90
	9	7.76	1.830	11.50
	10	8.05	1.612	5.25
	11	8.08	0.242	2.15

## **Appendix 3**

**Analytical results from BGR, Germany  
ICP-MS (filtered [0.45µm], acidified water samples)**



## 44155-W

φ1m			Ag	Al	As	B	Ba	Be	Bi	(Br)	(Ca)	Cd	Ce	Co
Probenidentifikation														
φ22m														
IM14360	44155/01	SV 1A5 FS	0.001	-0.1	0.25	188	112	0.33	0.001	476	103000	0.044	0.01	0.075
IM14361	44155/02	SV 2A41FS	-0.001	1.05	1.03	223	111	0.4	-0.001	487	120000	0.015	0.006	0.4
IM14362	44155/03	SV 3A1 FS	-0.001	-0.1	0.93	238	112	0.4	-0.001	486	126000	0.008	0.027	0.22
IM14363	44155/04	SV 5A3 FS	0.001	1.41	0.15	391	97	0.23	-0.001	751	111000	0.006	0.001	0.044
IM14364	44155/05	SV 5A4 FS	0.001	0.12	0.19	359	98.7	0.38	0.002	697	108000	0.011	0.001	0.088
IM14365	44155/06	SV 5Dii FS	0.001	0.73	0.26	377	99.3	0.2	-0.001	698	104000	0.012	0.001	0.029
IM14366	44155/07	SV 6A1 FS	0.001	2.3	0.036	372	115	0.56	0.002	697	120000	0.008	0.059	0.043
IM14367	44155/08	SV K1 FS	0.001	1.15	1.5	68.7	38.1	0.084	-0.001	103	87900	0.008	0.015	0.022
IM14368	44155/09	SV K2 FS	0.031	8.13	0.48	6.75	9.62	0.019	0.001	17.1	53000	0.03	0.025	0.034
IM14369	44155/10	SV 1JA3 FS	0.003	-0.1	0.42	969	97.7	0.65	0.002	2680	127000	0.011	0.001	0.062
IM14370	44155/11	SV 2JB1 FS	0.002	3.76	6.17	605	53.3	0.2	0.004	1420	116000	0.012	0.005	0.023
IM14371	44155/12	SV BLANK 1	-0.001	-0.1	0.014	-0.01	0.013	-0.002	-0.001	0.33	13.6	-0.002	0.001	-0.002
IM14372	44155/13	SV BLANK 2	-0.001	-0.1	-0.01	-0.01	-0.005	-0.002	-0.001	0.32	-1	-0.002	0.001	0.006

minus sign denote below detection limit: i.e. -1 means < 1

## 44155-W

φ1m			Cr	Cs	Cu	(Fe)	Ga	(Ge)	Hf	Hg	In	(I)	(K)	La
Probenidentifikation														
φ22m														
IM14360	44155/01	SV 1A5 FS	1.14	27.1	0.28	15.2	0.015	9.39	0.32	-0.005	-0.001	9.89	13700	0.008
IM14361	44155/02	SV 2A41FS	1.47	27	0.29	29.5	0.015	9.47	0.045	-0.005	-0.001	6.27	15600	0.007
IM14362	44155/03	SV 3A1 FS	1.32	27.6	0.24	22	0.016	9.62	0.16	-0.005	-0.001	11.1	15900	0.016
IM14363	44155/04	SV 5A3 FS	3.57	38.1	0.28	39.2	0.014	15	0.056	0.008	-0.001	10.7	20800	0.001
IM14364	44155/05	SV 5A4 FS	2.25	35.6	0.28	45.7	0.012	14	0.01	0.006	-0.001	17.2	20800	0.001
IM14365	44155/06	SV 5Dii FS	3.14	39.8	0.24	29.2	0.012	14.4	0.028	-0.005	-0.001	16.8	20700	0.001
IM14366	44155/07	SV 6A1 FS	2.25	37.9	0.64	61.6	0.018	14.2	0.19	-0.005	-0.001	30.6	20600	0.032
IM14367	44155/08	SV K1 FS	0.45	6.55	4.32	39.1	0.011	1.99	0.013	-0.005	-0.001	7.73	5310	0.028
IM14368	44155/09	SV K2 FS	0.23	0.017	11.8	50.5	0.016	0.002	-0.002	-0.005	-0.001	1.05	1130	0.015
IM14369	44155/10	SV 1JA3 FS	2.18	142	0.34	38.5	0.011	25.8	0.015	0.027	-0.001	76.3	54300	0.002
IM14370	44155/11	SV 2JB1 FS	0.9	48.8	0.25	25.4	0.013	9.5	0.073	-0.005	-0.001	33.2	30500	0.007
IM14371	44155/12	SV BLANK 1	-0.005	-0.001	0.48	-1	-0.001	-0.002	-0.002	0.012	-0.001	-0.1	7.98	-0.001
IM14372	44155/13	SV BLANK 2	0.019	-0.001	0.16	-1	-0.001	-0.002	-0.002	0.011	-0.001	-0.1	-1	-0.001

## 44155-W

φ1m			Li	(Mg)	Mn	Mo	(Na)	Nb	Ni	(PO4)	Pb	Rb	Sb	Sc
Probenidentifikation														
φ22m														
IM14360	44155/01	SV 1A5 FS	287	30900	11.1	0.31	348000	0.01	0.25	3.31	0.031	83.3	-0.002	7.13
IM14361	44155/02	SV 2A41FS	324	35700	16	0.36	350000	0.021	0.26	7.87	0.021	83	0.018	8.43
IM14362	44155/03	SV 3A1 FS	343	40300	15.1	0.32	352000	0.019	0.17	7.27	0.008	83.5	-0.002	8.63
IM14363	44155/04	SV 5A3 FS	484	36900	13.5	0.18	354000	0.093	0.2	21.2	0.002	110	0.064	12.1
IM14364	44155/05	SV 5A4 FS	467	38600	2.02	0.22	356000	0.019	0.25	12.7	0.006	104	0.026	11.5
IM14365	44155/06	SV 5Dii FS	468	35600	0.021	0.23	358000	0.081	0.22	16.3	0.006	108	0.035	11.8
IM14366	44155/07	SV 6A1 FS	468	35600	14.4	0.15	360000	0.084	0.18	11.6	0.07	105	0.005	10.8
IM14367	44155/08	SV K1 FS	89.8	14900	0.016	0.23	53200	0.014	0.3	7.54	0.018	19.7	0.055	0.91
IM14368	44155/09	SV K2 FS	1.52	4910	0.16	0.11	2930	0.011	0.29	7.3	0.36	0.21	0.039	0.45
IM14369	44155/10	SV 1JA3 FS	1030	33100	11.7	0.04	366000	0.011	0.32	22.4	0.002	341	0.019	20.6
IM14370	44155/11	SV 2JB1 FS	629	29500	-0.01	1.56	368000	0.022	0.39	16.4	0.15	168	0.078	11.1
IM14371	44155/12	SV BLANK 1	0.043	5.54	-0.01	0.001	-1	-0.002	-0.002	-1	0.047	0.002	0.007	0.017
IM14372	44155/13	SV BLANK 2	0.016	5.26	-0.01	0.002	-1	0.07	-0.002	-1	-0.002	-0.002	0.016	-0.005

## 44155-W

φ1m			Se	Sn	Sr	Ta	Te	Th	Ti	Tl	U	V	W	Y
Probenidentifikation														
φ22m														
IM14360	44155/01	SV 1A5 FS	2.26	0.022	1140	0.019	0.002	0.021	0.69	0.004	0.41	0.3	1.6	0.086
IM14361	44155/02	SV 2A41FS	0.89	0.017	1130	0.029	0.012	0.003	0.64	0.007	0.18	0.34	1.95	0.034
IM14362	44155/03	SV 3A1 FS	1.32	0.007	1110	0.004	0.006	0.014	0.84	0.006	0.51	0.36	1.41	0.11
IM14363	44155/04	SV 5A3 FS	1.04	0.003	1130	0.054	0.011	0.003	0.87	0.006	0.42	0.59	5.75	0.008
IM14364	44155/05	SV 5A4 FS	1.19	0.004	1130	0.026	0.01	0.001	0.89	0.007	0.12	0.65	5.03	0.009
IM14365	44155/06	SV 5Dii FS	0.57	0.004	1110	0.085	0.018	0.001	0.99	0.007	0.48	0.63	4.06	0.009
IM14366	44155/07	SV 6A1 FS	1.36	0.016	1150	0.018	0.009	0.021	1.1	0.004	0.33	0.58	4.37	0.16
IM14367	44155/08	SV K1 FS	0.63	-0.002	325	0.001	0.005	0.002	0.42	0.004	0.62	0.17	0.45	0.088
IM14368	44155/09	SV K2 FS	0.14	0.014	50.6	0.001	0.005	0.002	0.83	0.002	0.41	0.12	-0.002	0.036
IM14369	44155/10	SV 1JA3 FS	0.83	0.004	1430	0.009	0.005	0.001	1.34	0.005	0.005	0.45	14.3	0.017
IM14370	44155/11	SV 2JB1 FS	1.85	0.01	953	0.004	0.013	0.005	0.91	0.016	2.36	0.51	3.24	0.03
IM14371	44155/12	SV BLANK 1	0.13	0.003	-0.01	0.007	-0.001	-0.001	-0.1	-0.002	0.001	-0.01	0.002	-0.001
IM14372	44155/13	SV BLANK 2	0.53	-0.002	-0.01	0.006	-0.001	-0.001	0.18	-0.002	-0.001	-0.01	0.051	0.001

## 44155-W

φ1m			Zn	Zr	Pr	Nd	Sm	Eu	Gd	Tb	Dy	Ho	Er	Tm
Probenidentifikation														
φ22m														
IM14360	44155/01	SV 1A5 FS	1.55	0.38	0.001	0.008	0.004	0.004	0.004	0.001	0.006	0.001	0.004	-0.001
IM14361	44155/02	SV 2A41FS	1.44	0.043	-0.001	0.002	0.002	0.003	0.001	-0.001	-0.001	-0.001	0.003	-0.001
IM14362	44155/03	SV 3A1 FS	0.74	0.16	0.004	0.017	0.004	0.005	0.008	0.001	0.012	0.002	0.01	0.001
IM14363	44155/04	SV 5A3 FS	24.4	0.069	-0.001	-0.001	-0.001	0.002	0.001	-0.001	-0.001	-0.001	-0.001	-0.001
IM14364	44155/05	SV 5A4 FS	253	0.011	-0.001	0.001	-0.001	0.003	0.001	-0.001	-0.001	-0.001	0.001	-0.001
IM14365	44155/06	SV 5Di FS	7.92	0.034	-0.001	0.002	0.002	0.006	0.001	-0.001	-0.001	-0.001	-0.001	-0.001
IM14366	44155/07	SV 6A1 FS	1.68	0.21	0.007	0.029	0.009	0.011	0.01	0.002	0.016	0.003	0.014	0.002
IM14367	44155/08	SV K1 FS	38.6	0.019	0.006	0.03	0.007	0.001	0.008	0.001	0.009	0.002	0.006	0.001
IM14368	44155/09	SV K2 FS	57.2	0.01	0.004	0.017	0.006	0.001	0.006	0.001	0.006	0.001	0.003	-0.001
IM14369	44155/10	SV 1JA3 FS	2.25	0.008	-0.001	-0.001	0.001	0.008	-0.001	-0.001	-0.001	0.001	0.001	-0.001
IM14370	44155/11	SV 2JB1 FS	7.95	0.12	0.002	0.004	0.001	0.003	0.003	0.001	0.002	0.001	0.002	0.001
IM14371	44155/12	SV BLANK 1	1.63	-0.002	-0.001	-0.001	-0.001	-0.001	0.001	-0.001	-0.001	-0.001	-0.001	-0.001
IM14372	44155/13	SV BLANK 2	0.71	0.005	-0.001	-0.001	-0.001	-0.001	-0.001	-0.001	-0.001	-0.001	-0.001	-0.001

ø1m					
Probenidentifikation				Yb	Lu
ø22m					
IM14360	44155/01	SV	1A5 FS	0.005	0.001
IM14361	44155/02	SV	2A41FS	0.002	-0.001
IM14362	44155/03	SV	3A1 FS	0.009	0.001
IM14363	44155/04	SV	5A3 FS	0.001	-0.001
IM14364	44155/05	SV	5A4 FS	0.002	-0.001
IM14365	44155/06	SV	5Dii FS	-0.001	-0.001
IM14366	44155/07	SV	6A1 FS	0.01	0.002
IM14367	44155/08	SV	K1 FS	0.007	0.002
IM14368	44155/09	SV	K2 FS	0.005	0.001
IM14369	44155/10	SV	1JA3 FS	0.001	-0.001
IM14370	44155/11	SV	2JB1 FS	0.004	0.001
IM14371	44155/12	SV	BLANK 1	-0.001	-0.001
IM14372	44155/13	SV	BLANK 2	-0.001	-0.001

## **Appendix 4**

**Field analyses by David Banks and Ron S. Sletten**

Alkalinity was measured by field titration against 0.1 N acid, using a mixed indicator with end-point at 4.3. At each spring 3 titrations were performed and the average presented.								
Ron S. has the complete set of T, pH and Eh measurements								
Bubbling: Cont. = Continuously, Epi. = Episodically / Intermittently,								
Date	Spring	Alkalinity (meq/l) Indiv Readings	Alkalinity (meq/l) Average	Eh mV	pH	T °C	Bubbling ?	Runoff
14/07/96	1A5	14.0 13.0 12.7	13.2	-175		21.7	Epi.	No
14/07/96	2A4i	12.9 13.5 13.1	13.2				Epi.	to 2A4ii
14/07/96	2A4ii	13.8 12.9 13.0	13.2				Cont.	No
14/07/96	3A1	13.0 13.0 13.2	13.1				Epi.	Nasally detectable H2S 1 to 2 l/s
14/07/96	3Di	11.3 11.2 11.5	11.3				No	Spring reemergence at river level Yes
14/07/96	K1	5.3 5.4 5.5	5.4				No	Cold spr. between 3 & 4 0.3-0.5 l/s
14/07/96	K2	2.2 2.2 2.3	2.2		7.98	0.5	No	Cold spr. up slope between 4 & 5 Yes
14/07/96	5A1						Epi.	No
14/07/96	5A2						Epi.	No
14/07/96	5A3	13.6 13.5 13.5	13.5				Epi.	No
14/07/96	5A4	13.4 13.8 13.7	13.6				Epi.	Yes
14/07/96	5Dii	12.9 12.8	12.9				No	Spring reemergence at river level Yes
15/07/96	6A1	13.5 13.4 14.1	13.7	+ 70	6.69	20.0	No	Yes
15/07/96	6A2			+105-110	6.69	20.0		Very little
16/07/96	1JA1			- 170	6.71	21.9	Epi.	Jotunkilde 1, A position Very little
16/07/96	1JA2			- 160	6.69	21.8	No	Jotunkilde 1, A position Very little
16/07/96	1JA3			- 100	6.67	23.0	Epi.	Jotunkilde 1, A position Yes
16/07/96	1JA3 (after sediment sampled & disturbed)			- 250		23.3		
16/07/96	2JB			+ 200	7.08	13.8	No	Jotunkilde 2, B position c. 0.05 l/s
Data entry checked by DB 31/7/96								

Runoff = surface run-off from spring area. It is assumed that most springs have groundwater runoff by reinfiltration of water into travertine deposits.

Epi. = episodic



ldnr	*Spring	Type	Alk_field	Eh_DB	Eh_Ron	Eh_field	pH_DB	pH_Ron	pH_field	T_DB	T_Ron	T_field	*Bubbling	*Commen	*Runoff
1	1A1	1	-9999.99	-9999.99	178	178	-9999.99	6.91	6.91	-9999.99	12.9	12.9	-9999.99	-9999.99	-9999.99
2	1A2	1			117	117		8.11	8.11		9.7	9.7			
3	1A5	1	13.2	-175	-220	-220		6.80	6.80	21.7	21.6	21.6	Epi.		No
4	1B1	3									7.2	7.2			
5	1B2	3									11.3	11.3			
6	1B3	3									17.8	17.8			
7	1B3'	3									8.0	8.0			
8	1B4	3			120	120		6.95	6.95		17.2	17.2			
9	1C1	5			-235	-235		7.03	7.03		13.4	13.4			
10	2A2	1			-100	-100		7.78	7.78		10.2	10.2			
11	2A3	1			-150	-150		6.79	6.79		22.8	22.8			
12	2A4i	1	13.2		-240	-240		6.68	6.68		22.5	22.5	Epi.		to 2A4ii
13	2A4ii	1	13.2		-60	-60		7.19	7.19		13.4	13.4	Cont.		No
14	3A1	1	13.1		-273	-273		6.83	6.83		25.6	25.6	Epi.	Nasally d	1 to 2 l/s
15	3A2	1			-33	-33		7.60	7.60		19.6	19.6			
16	3A3	1			-24	-24		8.18	8.18		9.8	9.8			
17	3D1	6	11.3		216	216		7.65	7.65		9.5	9.5	No	Spring ree	Yes
18	3Dii	6			209	209		7.34	7.34		12.1	12.1			
19	K1	7	5.4		203	203		7.00	7.00		3.6	3.6	No	Cold spr.	0.3-0.5 l/s
20	K2	8	2.2		230	230	7.98	7.97	7.97	0.5	0.2	0.2	No	Cold sprin	Yes
21	5A1	1			259	259		6.94	6.94		18.2	18.2	Epi.		No
22	5A2	1			203	203		6.83	6.83		19.1	19.1	Epi.		No
23	5A3	1	13.5		133	133		6.78	6.78		19.4	19.4	Epi.		No
24	5A4	1	13.6		159	159		7.47	7.47		14.5	14.5	Epi.		Yes
25	5B1	3			160	160		7.10	7.10		12.2	12.2			
26	5Dii	6	12.9		160	160		7.18	7.18		11.8	11.8	No	Spring ree	Yes
27	6A1	1	13.7	70		70	6.69		6.69	20.0		20.0	No		Yes
28	6A2	1		107		107	6.69		6.69	20.0		20.0			Very little
29	1JA1	2		-170		-170	6.71		6.71	21.9		21.9	Epi.	Jotunkilde	Very little
30	1JA2	2		-160		-160	6.69		6.69	21.8		21.8	No	Jotunkilde	Very little
31	1JA3	2		-100		-100	6.67		6.67	23.0		23.0	Epi.	Jotunkilde	Yes
32	1JA3 (b)	2		-250		-250				23.3		23.3		(after sediment samp	
33	2JB	4		200		200	7.08		7.08	13.8		13.8	No	Jotunkilde	c. 0.05 l/s

This file was taken from the checked svalbalk main spreadsheet 15/1/97 and used for the creation of svmajdas.wks on the same date

-9999.99 = no data

Type: 1 = A spring Troll, 2 = A spring Jotun, 3 = B spring Troll, 4 = B spring Jotun, 5 = C spring Troll, 6 = D spring Troll, 7 = K1 spring Troll, 8 = K2 spring Troll

## **Appendix 5**

### **Rock analyses from BGR**

- (i) Summary and comparison of rock analyses**
- (ii) Rock sample list**
- (iii) XRF analysis results**
- (iv) ICP-MS analysis results**

**(i) Whole Rock Trace Geochemistry; Bockfjord, Svalbard. Summary and Comparison of XRF and ICP-MS Techniques.**

**Sample Selection**

The Bockfjord area of northern Spitsbergen is the site of what are believed to be the world's northernmost thermal springs (Hoel & Høltedahl 1913, Liestøl 1976, Banks et al. 1997, *in prep.*). The springs arise from a fault zone separating Upper Hecla Hoek Precambrian age mica schists and marbles from the Devonian "Old Red Sandstones" of the Wood Bay / Red Bay Series (Friend 1961, Hjølle 1993). Travertine basins have been formed associated with fossil (inactive) and current (active) thermal springs. Eleven samples of rock were collected from the Bockfjord spring areas in July 1996, as follows:

Sv1R - Grey, fine-grained, Hecla Hoek marble from Trollkildene

Sv3R - Travertine from an active travertine mound, Spring 5A4, Trollkildene

Sv5R - Buff brecciated Hecla Hoek marble from Trollkildene

Sv10R - Red-brown travertine from a fossil basin, spring area 1, Trollkildene.

Sv11Ra - Travertine from laminar layers in the wall of a fossil travertine basin, spring area 3, Trollkildene

Sv11Rb - Travertine from porous layers in the wall of a fossil travertine basin, spring area 3, Trollkildene

Sv14R - Fossil travertine, spring area 7, Trollkildene.

Sv15R - Travertine from active basin, spring area 3, Trollkildene

Sv17R - Greyish, coarser-grained marble, Hecla Hoek, Jotunkildene

Sv19R - Frost-shattered mica-schist the Hecla Hoek, Jotunkildene

Sv28R - Devonian Old Red Sandstone, near Trollkildene

XRD analysis of travertine samples Sv3R, Sv10R and Sv15R showed them to be almost exclusively composed of calcite.

**Sample Preparation and Analysis**

At the Geological Survey of Norway, samples were prepared by crushing and fine milling in a clean agate mill. Further sample preparation and analysis were performed at the Bundesanstalt für Geowissenschaften und Rohstoffe in Germany.

For X-Ray Fluorescence (XRF) analysis, 1000 mg of milled samples were heated to 1030°C for 15 minutes to determine loss on ignition. After mixing the residue with 5.0 g lithium metaborate and 100 mg iodine pentoxide, the samples were fused at 1150°C for 20 minutes in

Pt-Au crucibles to produce glassy XRF discs (Török & van Grieken 1992, Török et al. 1996). Analysis was performed using a wavelength dispersive X-ray spectrometer, with both chromium (PW 1480) and rhodium (PW 1400) tubes to achieve optimal results for different elements. Monitor samples and more than 50 certified reference materials were used for calibration and matrix correction (de Jongh's alpha-correction) of results.

For Inductively Coupled Plasma Mass Spectrometry (ICP-MS) analysis, 100 mg of milled sample was added to 5 ml of concentrated ultrapure HNO<sub>3</sub> and 2 ml HF. The mixture was heated at 170°C for 24 hours in a sealed Teflon bomb. The reaction vessel was then carefully opened and heated to near-dryness on a hotplate. 3 ml of concentrated HCl was added and samples were again heated in the sealed vessel at 150°C for 12 hrs. Finally 1 ml HNO<sub>3</sub> and 1 ml HCl were added to the residue and the solution topped up with pure water to 100 g total (i.e. a total dilution factor of 1000), prior to ICP-MS analysis using the instrument described in Reimann et al. (1996), who also cite detection limits for aqueous solutions of each element. For both XRF and ICP-MS analyses, results are cited as ppm relative to dry, non-ignited weight.

## Results and Discussion

The results are presented as x-y plots of XRF versus ICP-MS results for trace elements, all in ppm (Fig. 13). In many cases, XRF results fall below cited detection limits. In these cases, the values are plotted at half the detection limit and the detection limit is shown by an arrow. For the elements Bi, Mo, Sb and W, results obtained by the two techniques were of completely differing orders of magnitude. This is to be expected as such elements may be lost by volatilization, e.g. as volatile fluorides or chlorides during acid dissolution and heating for ICP-MS analysis (in fact, fluorides of W and Mo and a chloride of Sb have boiling points significantly under 100°C - Lide 1991).

Values obtained by the two techniques are of similar orders of magnitude for Cu, Hf, La, Sn and U, but correlations were poor. For As, Ce, Nb, Sc and V, good correlations are achieved at high concentrations, but at low concentrations the correlations are poorer. Good correlations between XRF and ICP-MS were obtained for Ba, Co, Cr, Ga, Mn, Ni, Pb, Rb, Sr, Th, Ti and Zn. In the case of Zr, a reasonable correlation was also obtained, though with generally lower values from ICP-MS, probably reflecting difficulties in total solution of, for example, zircon minerals in the ICP-MS digestion procedure.

The travertines, composed of almost pure calcite, show the lowest concentrations of most trace elements, followed by the marbles. The schist and sandstone generally yield the highest values. Exceptions include Ba, where the travertines are enriched compared to the marbles. In the case of Sr, the travertines reflect the Sr-enriched (though strontianite-undersaturated) nature of the thermal waters themselves (Fig. 14). The thermal waters are saturated with

respect to barite so it is possible that the travertines contain traces of this mineral. As XRD analysis did not indicate the presence of Sr or Ba minerals, however, it is regarded as more likely that these elements are substituting for chemically analogous Ca in the calcite lattice.

In conclusion, the results of this comparative study confirm that ICP-MS analysis of acid digests of rock samples is a satisfactory means of determining whole rock contents of most trace elements, in comparison with XRF and considering the wide range of parameters determined by both methods. In the majority of cases, ICP-MS is capable of achieving considerably lower detection limits than XRF. Discrepancies between the techniques may be due to (i) volatilisation during preparation, (ii) matrix effects or interference with other elements, (iii) incomplete digestion of some minerals (e.g. zircon) during preparation. Discrepancies could be minimised by tailoring preparation techniques and interference correction algorithms to the elements of most interest.

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**Note: In Figure 13 and 14, points below detection limit are plotted at a value corresponding to half the detection limit. This is a purely artificial means of registering the fact that the analysis returning a "<" value and has NO correspondence to an actual concentration. This fact may give a misleading visual impression of a correlation (or lack of correlation) where such is, in fact, not necessarily the case.**

Figure 13. Analyses of rock samples from Svalbard for trace elements by XRF (x-axis) and ICP-MS (y-axis). All cited as ppm relative to dry, non-ignited weight. For each element, two figures are cited near the top of each box: the upper represents the best fit linear correlation coefficient ( $r$ ); the lower is the slope of the best fit regression line. Arrows show analytical detection limit.

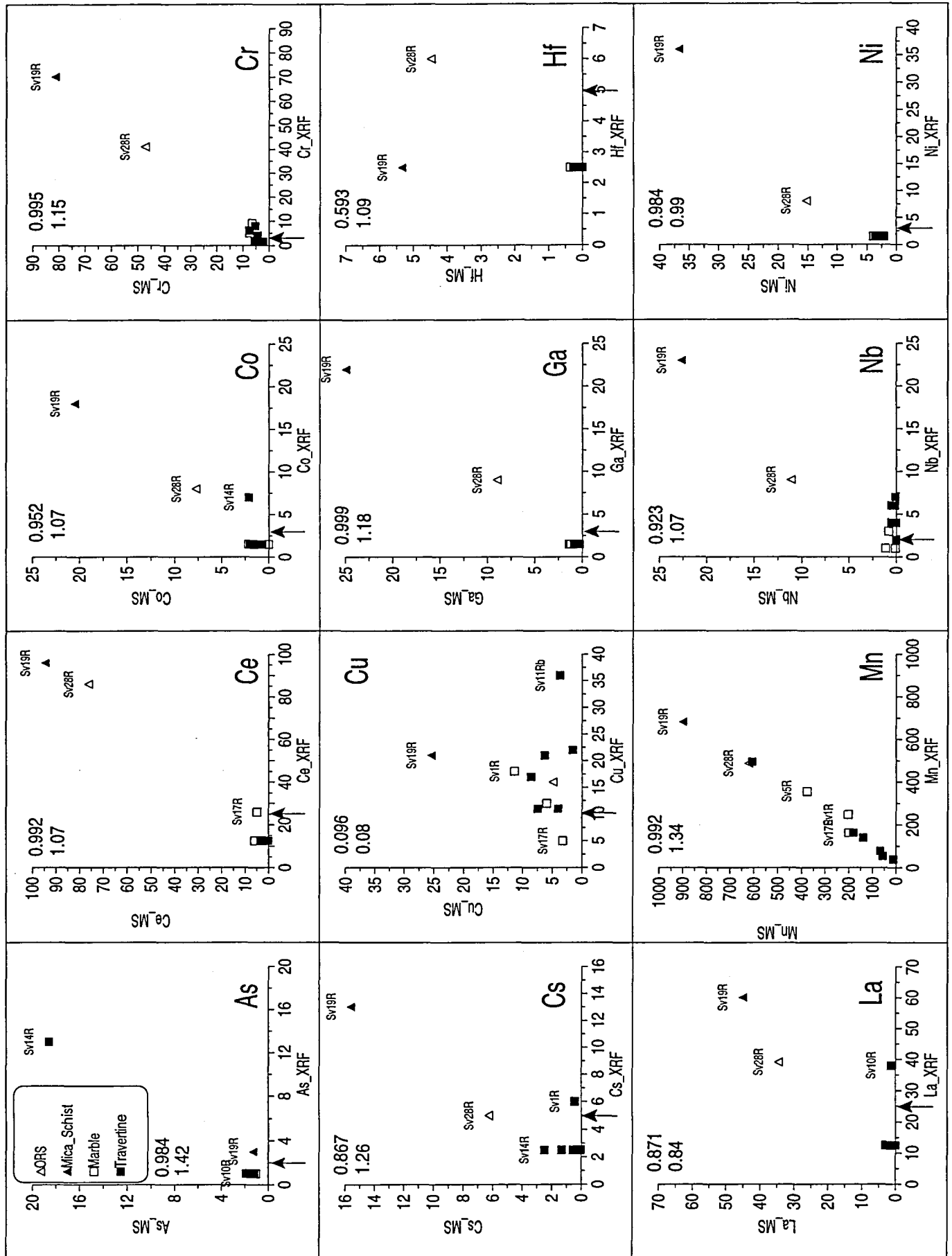


Figure 13 (continued).

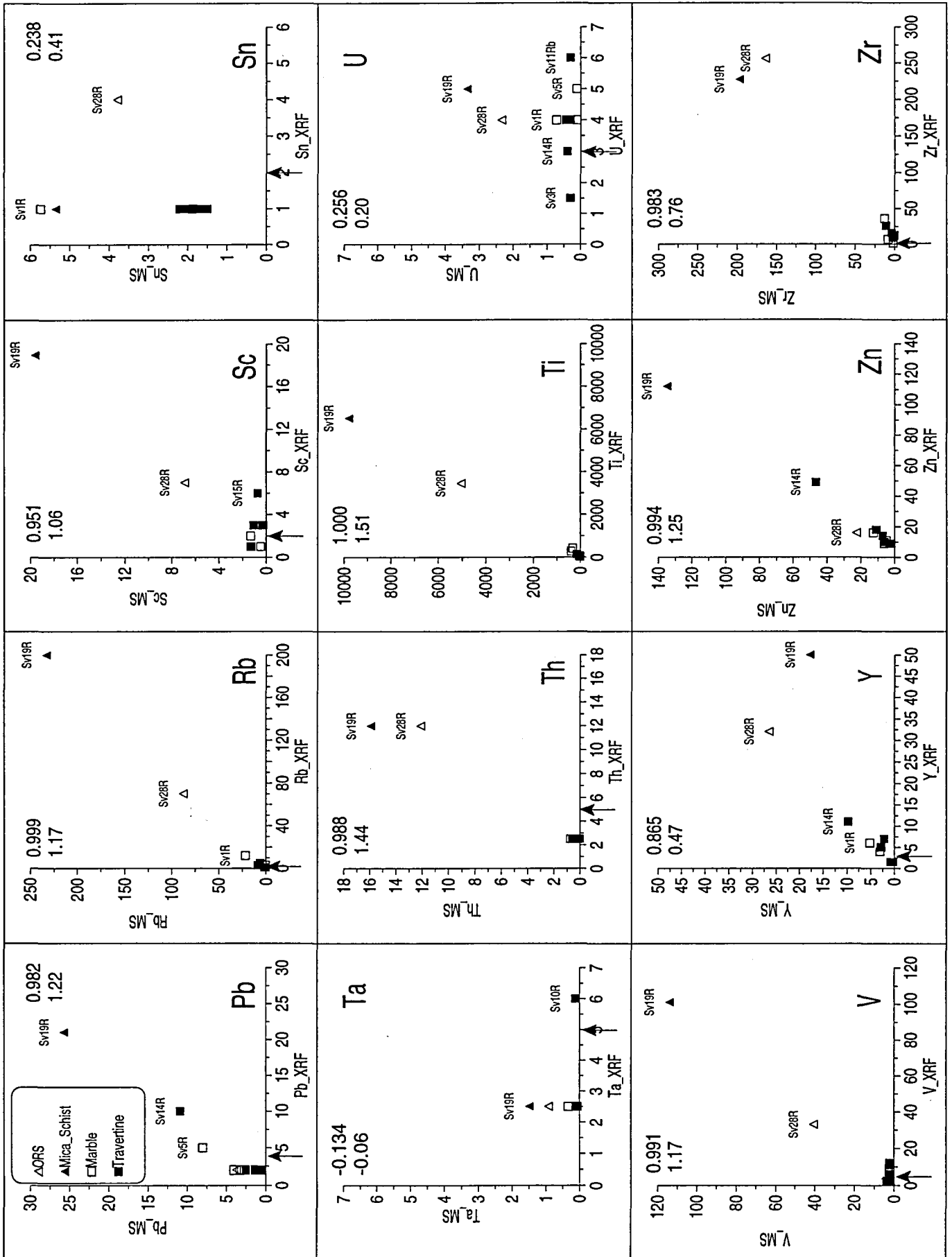
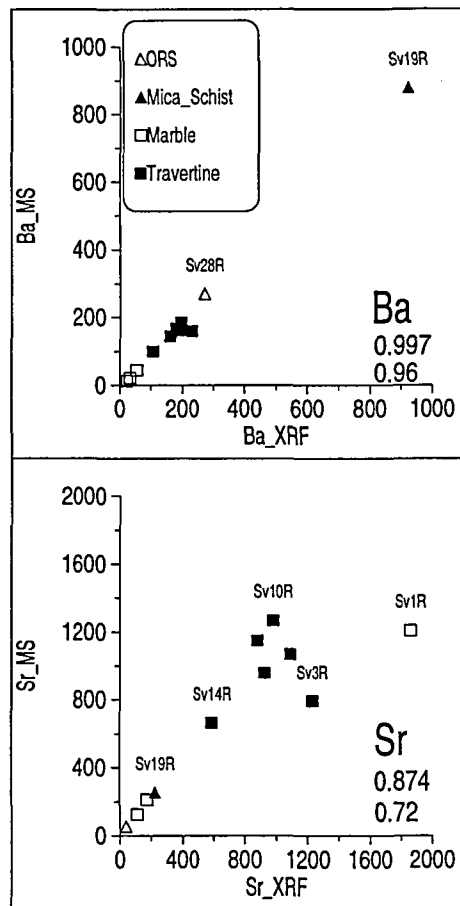


Figure 14. Analyses of rock samples from Svalbard for Ba and Sr by XRF (x-axis) and ICP-MS (y-axis), all cited as ppm relative to dry, non-ignited weight. For each element, two figures are cited near the bottom of each box: the upper represents the best fit linear correlation coefficient ( $r$ ); the lower is the slope of the best fit regression line.

Note: in Figures 13 and 14, the correlation coefficients are of little meaning where a significant proportion of samples lie below the detection limit.





## **Comments on Appropriateness of Analytical Techniques, from the Analyst (Ulrich Siewers, BGR, Germany)**

### **Uranium :**

The range of values for only eleven samples with XRF is < 3 to 6 ppm. At the detection limits it is generally agreed upon that an error of 100 per cent may appear, so one cannot rely on uranium in this range. Also, the limit of detection is, by definition, different to the limit of quantification. ICP-MS data are to be preferred.

### **Vanadium :**

The range of values for only eleven samples with XRF is < 5 to 101 ppm. It is thus only justifiable to compare two samples giving 33 ppm and 101 ppm respectively. for XRF with ICP-MS.

### **Tungsten :**

The range of values for only eleven samples with XRF is < 3 to 14 ppm. If scheelite ( $\text{CaWO}_4$ ) occurs in the sample, it will be difficult to attain total dissolution with acids, so the XRF data are to be preferred over ICP-MS. The variance between XRF and ICP-MS data may show this effect, but it may also be due to tungsten minerals being distributed inhomogeneously in the samples (for XRF analyses a 1000 mg sample is used, for ICP-MS only 100 mg).

### **Yttrium :**

The range of values for only eleven samples with XRF is < 3 to 50 ppm, with only three values which are justified for use in comparison, namely 11, 32 and 50 ppm. If xenotime ( $\text{YPO}_4$ ) occurs in the sample, it will be difficult to attain total dissolution with acids (this could for example explain the discrepancy in sample SV19R), so XRF data are to be preferred.

### **Zinc :**

No problem.

### **Zirconium :**

If  $\text{ZrSiO}_4$  occurs in the sample, it will be difficult to attain total dissolution with acids, particularly if the sample contains good crystals, rather than tiny weathered ones. XRF data are to be preferred.

### **Molybdenum :**

XRF data are to be preferred.  $\text{MoF}_4$  may be volatile and Mo may precipitate in an acid solution.

### **Niobium :**

The range of values for only eleven samples with XRF is < 2 to 23 ppm, with only three values which can be used for comparison, namely 7, 9 and 23 ppm. Some Nb-containing minerals cannot be totally dissolved in acid.

### **Nickel :**

The range of values for only eleven samples with XRF is < 3 to 36 ppm, with only two values which can be used for comparison, namely 8 and 36 ppm.

### **Lead :**

The range of values for only eleven samples with XRF is < 4 to 21 ppm, with only two values which can be used for comparison, namely 10 and 21 ppm.

### **Rubidium :**

the range of values for only eleven samples with XRF is < 2 to 200 ppm, with only three values which can be used for comparison, namely 12, 70 and 200 ppm.

### **Antimony :**

XRF data are to be preferred in this case. Maybe  $\text{SbF}_3$  is volatile or Sb precipitates in an acid solution.

### **Scandium :**

The range of values for only eleven samples with XRF is < 2 to 19 ppm, with only two values which can be used for comparison, namely 7 and 19 ppm. The range of concentration is not very favourable for XRF.

**Tin :**

The range of values for only eleven samples with XRF is < 2 to 4 ppm, with no value which can sensibly be used for comparison. The range of concentration is not very favourable for XRF.

**Strontium :**

XRF data are to be preferred where "major" concentrations are present (as in this case). ICP-MS to be preferred when present in trace concentrations.

**Tantalum :**

Comments as for several elements above. A difficult range of concentrations.

**Thorium :**

The range of values for only eleven samples with XRF is < 5 to 12 ppm, with only two values which can be used for comparison. Th sometimes occurs in zirconium minerals, see above under zirconium. XRF values are to be preferred above 12 ppm Th.

**Titanium :**

XRF values are to be preferred.

**Arsenic :**

The range of values for only eleven samples with XRF is < 2 to 13 ppm, , with only one value which can be used for comparison

**Barium :**

XRF values are to be preferred well above the detection limit. Otherwise, for trace quantities, use ICP-MS results.

**Bismuth :**

The range of values for only eleven samples with XRF is < 3 to 8 ppm, , with no values which can sensibly be used for comparison

**Cerium :**

The range of values for only eleven samples with XRF is < 25 to 96 ppm, , with only two values which can be used for comparison. Ce sometimes occurs in zirconium minerals, see above under zirconium. Prefer XRF values for concentrations above 40 ppm Ce.

**Cobalt :**

The range of values for only eleven samples with XRF is < 3 to 18 ppm, with only two values which can be used for comparison

**Chromium :**

The range of values for only eleven samples with XRF is < 3 to 70 ppm, with only two (or three) values which can be used for comparison

**Caesium :**

The range of values for only eleven samples with XRF is < 5 to 13 ppm, , with only one value which can be used for comparison

**Copper :**

The range of values for only eleven samples with XRF is < 10 to 36 ppm. XRF values should be preferred for concentrations above 20 ppm Cu. AAS analysis should ideally be taken as a control.

**Gallium :**

The range of values for only eleven samples with XRF is < 3 to 22 ppm, with only two values which can be used for comparison

**Hafnium :**

The range of values for only eleven samples with XRF is < 5 to 6 ppm, , with no values which can sensibly be used for comparison

**Lanthanum :**

The range of values for only eleven samples with XRF is < 25 to 60 ppm, with only one-three values which can be used for comparison. La is sometimes situated together with Ce in zirconium minerals, see above under zirconium. XRF values are to be preferred at concentrations above 40 ppm La.

**Manganese :**

XRF data are to be preferred where values are well above the detection limit. ICP-MS to be preferred when present in trace concentrations.

**(ii) Sample list****Bockfjord - Solid Samples: Preparation for ICPMS/XRF analysis****Prosjekt: 2717.00**

- Sv1R - several fragments of fine-grained grey marble, from outcrop in stream bed 1/3 of way between area 6 and moraine. Troll springs.
- Sv2R - several fragments of fine-grained buff-yellow marble, from outcrop in stream bed 1/3 of way between area 6 and moraine. Troll springs.
- Sv3R - several flakes of recently deposited travertine from the run-off surface of the travertine dome around spring 5A4, sample taken from within c. 2 m of spring outlet. Troll springs.
- Sv4R - secondary carbonate (?) mineralisation in the breccia outcrop near area 6 (see Sv6) Troll springs.
- Sv5R - brecciated and Qz-veined fine grained buff/yellow marble/dolomite, from outcrop area, near spring area 6. Troll springs.
- Sv6R - heavily brecciated and mineralised fine-grained grey marble, from outcrop area, near spring area 6. Troll springs.
- Sv7R - fragment of travertine from in-situ, topmost area 6 (RS has parallel sample). Troll springs.
- Sv8R - fragment of in-situ travertine from fossil travertine basins below active springs in area 6. The travertine fragment is from a horizontal bed, ca. 40 cm from the uppermost surface of the travertine basin. Troll springs.
- Sv9R - travertine basin bottom sediment from near core area in basin 2A4. Troll springs.
- Sv10R - fragment of deep red/brown travertine. Loose block from lowermost part of area 1's travertine basin complex. Troll springs.
- Sv11R - 2 fragments of travertine from a block fallen from «The Wall» (area 3), showing open and dense travertine layers. Troll springs.
- Sv12R - large fragment of travertine breccia/conglomerate from south of area 1, down towards river level. Troll springs.
- Sv13R - fragment of brecciated and Qz veined fine-grained buff/yellow marble/dolomite breccia from exposure scar behind spring area 3. Troll springs.
- Sv14R - two fragments of travertine from top of area 7 (RS has similar samples). Troll springs.
- Sv15R - samples of travertine from top of basin rim around basin 3A2 (RS has parallel sample). Troll springs.
- Sv16R - two loose fragments of copper mineralisation found as loose fragments in solifluxion or weathering deposits. Troll springs.
- Sv17R - large fragment of coarse grained marble. Loose block adjacent to marble outcrop below and slightly W of Jotunkilde 2. Shows solution features. (MH has other half of block). Jotun springs.
- Sv18R - loose block of travertine breccia found near base of travertine scree below Jotunkilde 2.
- Sv19R - fragment of mica schist thrown up in solifluxion furrow immediately above Jotunkilde 1
- Sv20R - fragment of mica schist thrown up in solifluxion furrow immediately above Jotunkilde 1 (MH has parallel sample).
- Sv21R - fragment of green (presumed Tertiary) sandstone in scree/minetip below mine entrance tunnel, higher mine tip, Mine 3, Bjørndalen, Longyearbyen
- Sv22R - fragments of (possibly somewhat weathered) coal in minetip below mine entrance tunnel, higher mine tip, Mine 3, Bjørndalen, Longyearbyen
- Sv23R - fragment of shale with yellow (iron sulphate ?) deposit on cleavage surfaces, lower tip area, Mine 3, Bjørndalen, Longyearbyen.
- Sv24R - fragments ditto

Sv25R - loose sandstone/conglomerate core from old drilling site near base of Bjørndalen valley, higher up valley from mine tips. Longyearbyen.

Sv26R - fragments of (possibly somewhat weathered) coal in minetip above Sverdrupbyen (Mine 1, new mine), Longyearbyen.

Sv27R - fragments of fresh coal from coal face, Mine 3, tourist stross (Longyear Coal). Longyearbyen.

Sv28R - Devonian Old Red Sandstone, Trollkildene. Collected by Michael Heim.

Samples Sv21R to 27R are samples collected from the mining area in Longyearbyen and have no relevance to this project.

**Sv 11R:** The sample is a layered travertine with two main types of layer:

a: dense, crystalline travertine

b: open porous travertine

Two subsamples were prepared from sample Sv11R (one from each type of layer) and marked Sv11Ra and Sv11Rb.

### **Prøveforberedelsen**

Prøvene, unntatt Sv19R og Sv28R er kalkstein av forskjellige typer.

Prøvene ønskes møllet til pulver i agatmølle. Minst 1 g homogenisert pulver legges i en liten plastflaske som skal sendes til Tyskland for ICPMS analyse.

Ved forberedelsen av prøvene, skal det i størst mulig grad unngås unødvendig kontakt med metall- eller andre potensielt forurensende gjenstander.

### **Spesielle betingelser:**

**Alle prøver:** støv, fremmed jord og tilvekst av alger (spesielt Sv15R) skal fjernes ved vasking i destillert vann før pulverisering. I størst mulig grad skal det tas den ferskeste delen av prøven til pulverisering. Dette vil være vanskelig i flere tilfeller siden travertinene er tildels fossile og dermed forvitrede.

## (iii) XRF analyses - Germany

	SiO <sub>2</sub>	TiO <sub>2</sub>	Al <sub>2</sub> O <sub>3</sub>	Fe <sub>2</sub> O <sub>3</sub>	MnO	MgO	CaO	Na <sub>2</sub> O	K <sub>2</sub> O*
SV 1 R	1.81	0.069	0.97	0.46	0.032	0.74	52.41	0.13	0.26*
SV 3 R	3.02	0.010	0.09	0.05	0.005	0.84	50.91	0.11	0.03*
SV 5 R	9.77	0.011	<0.05	0.68	0.046	17.99	29.27	0.01	0.01*
SV 10 R	2.06	0.023	0.43	0.86	0.018	1.13	51.77	0.13	0.10*
SV 11 Ra	0.14	0.006	<0.05	0.07	0.010	0.67	54.23	0.09	0.01*
SV 11 Rb	0.19	0.004	<0.05	0.05	0.007	1.01	53.79	0.10	<0.01*
SV 14 R	0.98	0.023	0.33	0.34	0.064	0.77	53.43	0.05	0.05*
SV 15 R	1.11	0.007	<0.05	0.15	0.021	0.55	52.58	0.20	0.02*
SV 17 R	2.80	0.046	0.71	0.36	0.021	0.75	52.62	0.29	0.07*
SV 19 R	61.89	1.088	16.30	7.53	0.088	2.13	1.79	2.37	4.71*
SV 28 R	80.36	0.577	6.44	2.80	0.063	0.87	2.05	0.34	1.51*
	P <sub>2</sub> O <sub>5</sub>	(SO <sub>3</sub> )	(Cl)	(F)	LOI	Summe	(As)	Ba	Bi*
	%	%	%	%	%	%	ppm	ppm	ppm*
SV 1 R	0.022	0.83	0.011	<0.020	41.92	99.66	<2	54	<3*
SV 3 R	0.009	1.19	0.003	0.310	43.29	99.87	<2	230	<3*
SV 5 R	0.015	0.89	0.036	0.045	41.10	99.87	<2	33	8*
SV 10 R	0.018	1.08	0.031	0.168	42.02	99.84	<2	179	4*
SV 11 Ra	0.008	0.94	0.005	0.217	43.47	99.87	<2	196	<3*
SV 11 Rb	0.009	0.96	0.006	0.161	43.59	99.88	<2	161	5*
SV 14 R	0.025	1.05	0.017	0.048	42.65	99.83	13	107	<3*
SV 15 R	0.013	1.33	0.014	0.239	43.60	99.83	<2	194	7*
SV 17 R	0.031	0.97	0.015	0.073	41.17	99.93	<2	23	5*
SV 19 R	0.111	0.06	0.049	<0.020	1.53	99.65	3	922	4*
SV 28 R	0.146	0.67	0.011	<0.020	3.96	99.80	<2	271	<3*
	Ce	Co	Cr	Cs	Cu	Ga	Hf	La	Mo*
	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm*
SV 1 R	<25	<3	5	6	18	<3	<5	<25	9*
SV 3 R	<25	<3	6	<5	11	<3	<5	<25	14*
SV 5 R	<25	<3	5	<5	12	<3	<5	<25	3*
SV 10 R	<25	<3	<3	<5	11	<3	<5	38	13*
SV 11 Ra	<25	<3	<3	<5	22	<3	<5	<25	17*
SV 11 Rb	<25	<3	<3	<5	36	<3	<5	<25	12*
SV 14 R	<25	7	8	<5	17	<3	<5	<25	16*
SV 15 R	<25	<3	4	6	21	<3	<5	<25	7*
SV 17 R	26	<3	9	<5	<10	<3	<5	<25	7*
SV 19 R	96	18	70	13	21	22	<5	<25	3*
SV 28 R	86	8	41	5	16	9	6	39	<2*

Brackets ( ) indicate elements for which the method is not particularly suited and the results are subject to a greater degree of uncertainty.

Loss on ignition (LOI) is measured at temperatures which may partially or totally remove CO<sub>2</sub> from carbonates. LOI thus includes loss of CO<sub>2</sub>, bound H<sub>2</sub>O, some sulphur and organic carbon. LOI is taken as part of the sum (summe).

	Nb	Ni	Pb	Rb	Sb	Sc	Sn	Sr	Ta*
	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm	ppm*
SV 1 R	<2	<3	<4	12	35	<2	<2	1861	<5*
SV 3 R	6	<3	<4	<2	22	3	<2	1232	<5*
SV 5 R	<2	<3	5	3	22	<2	<2	109	<5*
SV 10 R	4	<3	<4	3	6	3	<2	980	6*
SV 11 Ra	4	<3	<4	<2	15	3	<2	1090	<5*
SV 11 Rb	2	<3	<4	<2	28	3	<2	924	<5*
SV 14 R	6	<3	10	5	10	<2	<2	589	<5*
SV 15 R	7	<3	<4	<2	9	6	<2	880	<5*
SV 17 R	3	<3	<4	4	18	2	<2	174	<5*
SV 19 R	23	36	21	200	5	19	<2	227	<5*
SV 28 R	9	8	<4	70	7	7	4	39	<5*
	Th	U	V	W	Y	Zn	Zr		*
	ppm	ppm	ppm	ppm	ppm	ppm	ppm		*
SV 1 R	<5	4	<5	<3	6	9	36		*
SV 3 R	<5	<3	<5	6	<3	18	16		*
SV 5 R	<5	5	9	8	4	16	<3		*
SV 10 R	<5	<3	12	6	7	10	26		*
SV 11 Ra	<5	4	11	14	<3	9	13		*
SV 11 Rb	<5	6	<5	4	<3	9	12		*
SV 14 R	<5	3	6	3	11	49	9		*
SV 15 R	<5	4	<5	5	5	14	12		*
SV 17 R	<5	4	6	10	5	11	7		*
SV 19 R	12	5	101	4	50	112	227		*
SV 28 R	12	4	33	<3	32	16	256		*

\*\*\*\*\*

SAMPLE PREPARATION and ANALYSIS of ROCKS, ORE MINERALS, SOILS  
and STREAM SEDIMENTS by XRF

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The samples were milled to less than 40 5m particle size. 1000 mg were heated to 1030 OC for 15 min to determine loss on ignition (LOI). After mixing the residue with 5.0 g lithiummetaborate and 100 mg iodine pentoxide, it was fused at 1150 OC for 20 min for the rapid production of glasslike analytical disks with a good precision. It comprised the use of Pt-Au crucibles and a commercial automatic fluxer.

The advantages of fusion techniques are widely acknowledged to achieve homogeneous elemental distribution in glass disks and to reduce matrix effects (Tvrvk & van Grieken 1992, Tvrvk et al. 1996).

For quantitative analysis a wavelength-dispersive X-ray spectrometer (WDXRF) was used. For the determinations of TiO<sub>2</sub>, CaO, K<sub>2</sub>O and for the traces of Ba, Cs, Sb, Sc and Sn a Chromium-tube was used (PW 1480), for the determinations of As, Bi, Ce, Co, Cr, Cu, Ga, Hf, La, Mo, Nb, Ni, Pb, Rb, Sr, Ta, Th, U, V, W, Y, Zn and Zr and for SiO<sub>2</sub>, TiO<sub>2</sub>, Al<sub>2</sub>O<sub>3</sub>, Fe<sub>2</sub>O<sub>3</sub>, MnO, MgO, CaO, Na<sub>2</sub>O, K<sub>2</sub>O and P<sub>2</sub>O<sub>5</sub> a Rhodium-tube (PW 2400).

Monitor samples and more than 50 certified reference materials (CRM) were used in the correction procedures. Matrix corrections were done according to □(>

Literature :

Tvrvk, S.B. & van Grieken, R.E. (1992): X-ray Spectrometry.-  
Anal. Chem., 64: 180 R - 196 R.

Tvrvk, S.B., Labar, J., Injuk, J. & van Grieken, R.E. (1996):  
X-ray Spectrometry. Anal. Chem., 68: 467 R - 485 R.

## (iv) ICP-MS analyses on rock digests - Germany

	1	2	3	4	5	6	7	8*
identification	Ag	As	B	Ba	Be	Bi	Cd	Ce*
SV 1 R	0.027	1.65	2.36	44.3	0.30	.063	0.066	5.92*
SV 3 R	0.035	1.21	5.04	160	2.58	0.13	0.075	1.40*
SV 5 R	0.028	1.05	4.16	21.4	2.46	.034	0.073	2.66*
SV 10 R	0.020	1.91	6.83	168	6.10	.063	0.044	2.88*
SV 11 Ra	-.010	1.58	4.51	162	1.39	.020	0.035	0.083*
SV 11 Rb	-.010	1.37	5.14	145	1.57	.035	0.046	0.078*
SV 14 R	0.020	18.6	4.32	100	4.53	.050	0.44	3.48*
SV 15 R	-.010	1.18	6.91	185	3.60	.047	0.041	0.92*
SV 17 R	-.010	1.35	7.05	12.2	0.21	.077	0.052	4.91*
SV 19 R	0.11	1.29	8.71	879	2.36	0.13	0.089	94.5*
SV 28 R	0.048	1.32	47.4	270	1.85	.096	0.074	76.2*
identification	Ag	As	B	Ba	Be	Bi	Cd	Ce*
-----*								
	9	10	11	12	13	14	15	16 *
	Co	Cr	Cs	Cu	Ga	Ge	Hf	(Hg) *
SV 1 R	0.046	7.45	0.44	11.4	1.38	0.073	0.36	0.035*
SV 3 R	0.74	7.61	0.58	7.46	0.41	4.07	0.10	-.010*
SV 5 R	1.64	7.41	0.048	5.98	0.38	0.13	0.043	0.016*
SV 10 R	2.03	5.58	1.33	4.02	0.89	1.22	0.27	0.019*
SV 11 Ra	1.50	2.21	0.061	1.53	0.31	0.20	-.010	0.015*
SV 11 Rb	1.92	3.80	0.15	3.70	0.53	0.51	-.010	0.016*
SV 14 R	2.13	5.28	2.50	8.58	0.88	0.53	0.053	0.012*
SV 15 R	1.98	4.28	0.42	6.27	0.40	2.86	0.061	0.014*
SV 17 R	2.18	6.40	0.29	3.22	1.22	0.14	0.22	-.010*
SV 19 R	20.5	81.3	15.6	25.4	25.0	0.80	5.33	0.034*
SV 28 R	7.74	47.2	6.26	4.90	8.99	0.10	4.47	0.035*
	Co	Cr	Cs	Cu	Ga	Ge	Hf	(Hg) *
-----*								
	17	18	19	20	21	22	23	24*
	In	La	Li	Mn	Mo	Nb	Ni	Pb*
SV 1 R	0.005	2.94	1.53	201	0.14	1.16	3.82	4.06*
SV 3 R	-.005	0.66	7.40	11.7	0.17	0.15	2.60	2.53*
SV 5 R	0.007	1.20	4.45	378	0.22	0.14	2.46	8.02*
SV 10 R	-.005	1.36	7.28	139	0.097	0.55	3.00	1.36*
SV 11 Ra	-.005	0.034	3.03	67.1	0.058	0.033	2.67	-1.00*
SV 11 Rb	-.005	0.039	5.66	56.5	0.041	0.041	2.07	-1.00*
SV 14 R	-.005	2.59	2.77	607	0.094	0.52	3.54	10.9*
SV 15 R	-.005	0.45	6.52	180	0.057	0.11	3.09	1.28*
SV 17 R	0.006	2.65	2.59	199	0.055	0.83	3.90	3.26*
SV 19 R	0.085	45.0	63.9	900	0.30	22.7	36.7	25.7*
SV 28 R	0.032	34.6	18.6	621	0.15	11.1	15.3	3.70*
-----*								
	25	26	27	28	29	30	31	32*
	(Rb)	Sb	Sc	Se	Sn	Sr	Ta	Th*
SV 1 R	22.0	0.12	1.30	0.26	5.74	1210	0.34	0.74*
SV 3 R	2.15	0.030	0.34	0.61	2.20	793	0.10	0.31*
SV 5 R	0.62	0.27	0.48	1.39	1.88	125	0.097	0.15*
SV 10 R	8.95	0.047	1.08	0.11	1.85	1270	0.13	0.58*
SV 11 Ra	0.52	0.018	0.25	-.100	1.65	1070	0.056	0.015*
SV 11 Rb	0.51	0.014	0.35	0.13	1.50	959	0.066	0.019*
SV 14 R	6.41	0.17	1.32	0.22	1.69	665	0.11	0.40*
SV 15 R	2.02	0.032	0.73	0.14	2.01	1150	0.072	0.13*
SV 17 R	3.83	0.018	1.33	0.10	1.85	211	0.11	0.57*
SV 19 R	233	0.040	19.6	0.60	5.35	254	1.49	15.9*
SV 28 R	87.7	0.36	6.91	0.53	3.79	52.9	0.92	12.1*
-----*								



	33 (Ti)	34 Tl	35 U	36 V	37 W	38 Y	39 Zn	40 * (Zr)*
SV 1 R	328	0.087	0.72	3.78	1.08	5.18	6.27	12.8*
SV 3 R	24.7	0.013	0.29	1.69	0.28	0.84	10.9	3.86*
SV 5 R	45.9	0.015	0.11	2.77	0.31	3.08	12.6	1.79*
SV 10 R	142	-0.005	0.31	2.57	0.63	2.20	5.92	10.8*
SV 11 Ra	4.32	0.024	0.30	2.46	0.12	0.30	1.88	0.20*
SV 11 Rb	6.26	0.017	0.30	3.72	0.15	0.43	3.11	0.35*
SV 14 R	170	0.026	0.40	2.43	0.72	9.76	46.4	2.27*
SV 15 R	35.7	0.021	0.44	3.10	0.45	2.83	7.11	2.17*
SV 17 R	395	0.018	0.11	2.57	0.13	2.87	4.93	8.27*
SV 19 R	9810	0.71	3.35	114	1.96	17.8	135	197*
SV 28 R	5050	0.26	2.34	40.6	1.26	26.5	22.5	164*
-----*								
	41 Pr	42 Nd	43 Sm	44 Eu	45 Gd	46 Tb	47 Dy	48 * Ho*
SV 1 R	0.75	3.01	0.62	0.11	0.66	0.099	0.63	0.13*
SV 3 R	0.16	0.58	0.11	0.005	0.11	0.016	0.085	0.019*
SV 5 R	0.36	1.58	0.40	0.11	0.45	0.067	0.40	0.078*
SV 10 R	0.34	1.32	0.28	0.037	0.27	0.043	0.28	0.059*
SV 11 Ra	0.011	0.047	0.011	0.008	0.016	0.004	0.029	0.008*
SV 11 Rb	0.010	0.062	0.011	0.008	0.023	0.004	0.034	0.009*
SV 14 R	0.58	2.47	0.56	0.12	0.71	0.13	0.91	0.20*
SV 15 R	0.12	0.51	0.12	0.015	0.16	0.034	0.24	0.062*
SV 17 R	0.59	2.18	0.42	0.10	0.41	0.061	0.39	0.077*
SV 19 R	11.3	42.9	8.05	1.48	6.74	0.87	4.12	0.66*
SV 28 R	8.62	32.4	6.20	1.04	5.87	0.87	4.86	0.92*
-----*								
	49 Er	50 Tm	51 Yb	52 * Lu*				
SV 1 R	0.39	0.054	0.32	0.056*				
SV 3 R	0.064	0.009	0.063	0.013*				
SV 5 R	0.21	0.029	0.18	0.028*				
SV 10 R	0.17	0.026	0.18	0.029*				
SV 11 Ra	0.017	0.003	0.024	0.005*				
SV 11 Rb	0.027	0.003	0.029	0.006*				
SV 14 R	0.59	0.087	0.56	0.092*				
SV 15 R	0.19	0.030	0.18	0.034*				
SV 17 R	0.22	0.035	0.21	0.037*				
SV 19 R	1.65	0.22	1.40	0.23*				
SV 28 R	2.58	0.37	2.48	0.39*				
-----*								

"Feb" 7 1997

"QUANTITATIVE ANALYSIS: DATA SET SUMMARY"

"Gesteine, Svalbard/Norwegen"

100.0 mg + 5.0 ml HNO<sub>3</sub> + 2.0 ml HF (24 hrs at 170 OC) to dryness,  
+ 3.0 ml HCl (12 hrs at 150 OC),  
+ 1.0 ml HNO<sub>3</sub> + 1.0 ml HCl to 100.0 g (dilution factor = 1000.)

All results in ppm.

The minus sign signifies less than detection limit. -0.002 means <0.002

Brackets ( ) indicate elements for which the method is not particularly suited and the results are subject to a greater degree of uncertainty. This may be due to (a) high concentrations and out of range for a trace-element instrument, (b) low sensitivity to plasma excitation (counts per second per ppb/ppm) or (c) interferences.

## **Appendix 6**

### **XRD analysis of rocks by NGU**

ANALYSEKONTRAKT NR.: 1997.0043  
NGU PROSJEKT NR.: 2717.00

OPPDRAGSGIVER: NGU, Varme kilder Svalbard

ADRESSE:

TLF.: 310

KONTAKTPERSON: David Banks

PRØVETYPE: Bergart

ANTALL PRØVER: 3

IDENTIFIKASJON AV PRØVER: Iflg. liste fra oppdragsgiver

PRØVER MOTTATT: 13.02.97

ANMERKNINGER: Ingen

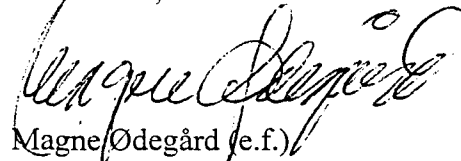
SPESIFIKASJON AV OPPDRAGET I HENHOLD TIL ANALYSEKONTRAKT:

METODE	DOKUMENTASJON *)	OMFATTES AV AKKREDITERING
XRD	NGU-SD 4.1	Ja

Denne rapporten inneholder i alt 5 sider + 3 vedlegg. Rapporten må ikke gjengis i utdrag uten skriftlig godkjenning fra NGU-Lab.

Alle forhold ved prøvetaking, behandling og transport av prøvene før innlevering til NGU-Lab er underlagt oppdragsgivers ansvar. Analyseresultater framlagt i denne rapporten refererer derfor kun til det prøvematerialet som er mottatt av NGU-Lab.

Trondheim, 17. februar 1997

  
Magne Ødegård (e.f.)

\*) Fortegnelse over dokumentasjon finnes i NGU-Labs Kvalitetshåndbok, NGU-SD 0.1, som kan rekvireres fra NGU-Labs sekretariat.

PROSJEKTNR.: 2717.00

OPPDRAKSGIVER: David Banks

PRØVEART: Bergart/mineral

ANTALL PRØVER: 3

MOTTATT: 13.02.97

BEHANDLINGSMETODE: XRD-opptak , scan 3-70°2θ , scan speed 0.02 °20/s

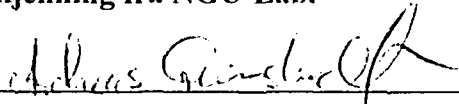
ARBEIDET UTFØRT AV: Andreas Grimstvedt

ANTALL SIDER: 3+3 vedlegg ( 3 stk. XRD-opptak for hver prøve med peak search og software basert identifisering-scorelist)

3  
PRESISJON : Det kjøres rutinemessig kontrollprøver (2θ-verdi/d-verdi), som føres i kontrolldiagram (X-diagram). Disse kan forevises om ønskelig.

ANMERKNINGER: Se de påfølgende sider

Rapporten må ikke gjengis i utdrag uten skriftlig godkjenning fra NGU-Lab.

Ferdig analysert	17.februar.1997	 Andreas Grimstvedt
------------------	-----------------	--

Dato

OPERATØR

## KOMMENTARER TIL XRD-ANALYSE

For hver prøve følger det med utskrift av diffraktogramet hvor beregnede topper er innsatt som vertikale linjer (peak search). I tillegg resultat fra software basert identifisering ("score list"), denne inkluderer et diffraktogram hvor topper for utvalgte mineraler er innsatt.

### PEAK SEARCH

Tabell med posisjonen (2 $\theta$ -verdi med tilhørende d-verdi) og høyden (intensitet) på beregnede topper. Videre kolonne med relative intensiteter (i forhold til høyeste topp)

### IDENTIFISERING

Som et hjelpemiddel for identifisering er de beregnede toppene benyttet i søk i database (JCPDS). Hovedprinsippet for søket er bestemmelse av antall topper i de ulike datakort som passer med topper i opptaket. Resultatet av et slikt søk er listet i en "score list".

#### Forklaring til "score list"

<u>Match score</u> :	Antall topper for det aktuelle mineral som passer med topper i opptaket
<u>Rel score</u> :	Fraksjon av topper for det aktuelle mineral som passer med opptaket, dvs. Match score/tot. antall topper(på datakortet) for mineralet
<u>Card-Id</u> :	Kortnr. for det aktuelle mineral i JCPDS-databasen

*NB! Navn på de listede mineraler er gitt på påfølgende side (sidenr. indeksert med .1) i samme rekkefølge.*

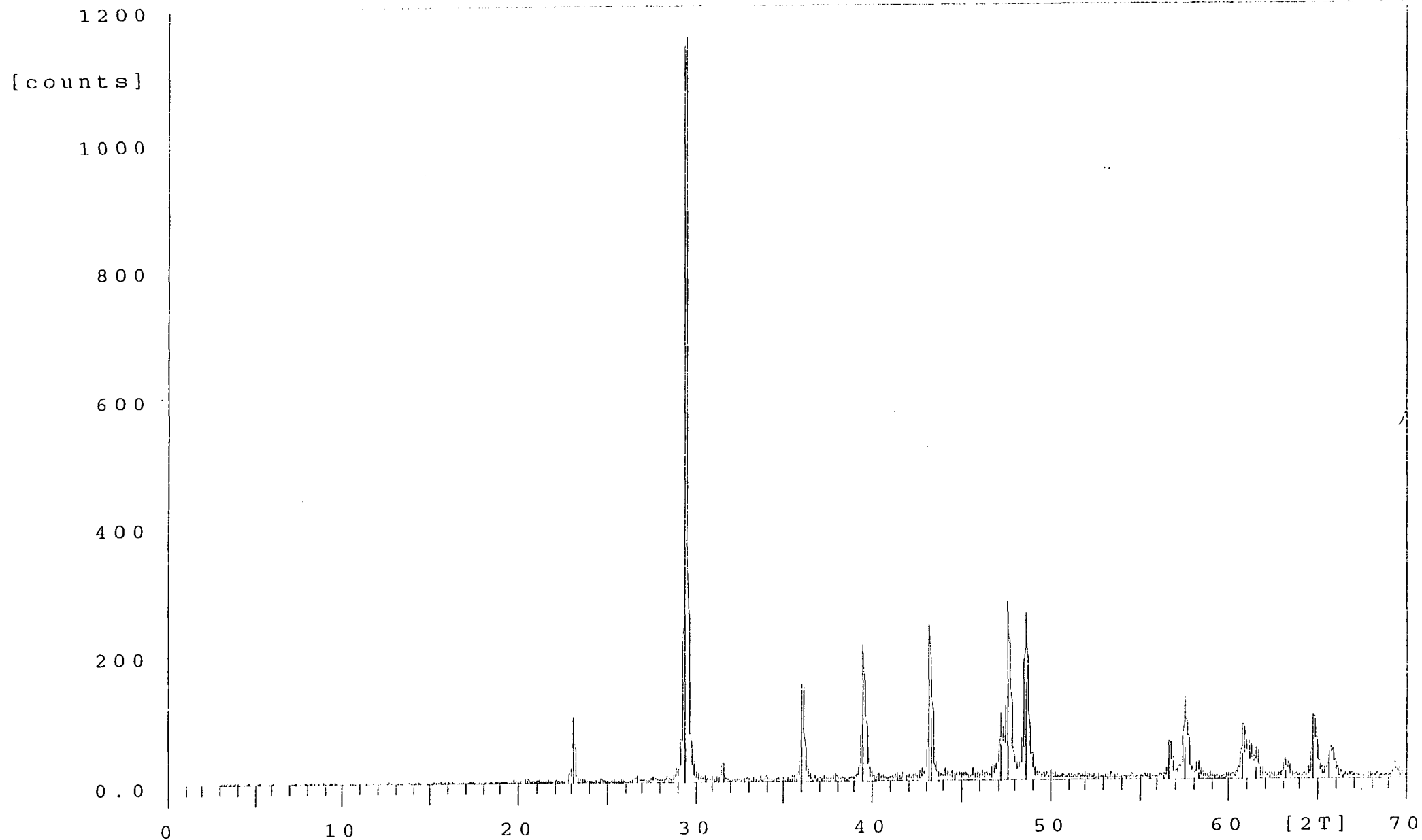
Forslag (innsatte) mineraler er merket med \* i kolonne 2.

I tillegg til "score listen" følger det med diffraktogramet hvor toppene til utvalgte mineraler er innsatt (vertikale linjer). Under dette diffraktogramet er det rubrikker med posisjon for topper. Den første rubrikken inneholder toppene til det aktuelle opptaket, mens de resterende er toppene på de innsatte mineralene (de samme som er merket med \* i "score list").

P.g.a. at en del mineraler har overlappende linjer behøver ikke alle forslagene å være tilstede i prøvene. Ta gjerne kontakt dersom mer informasjon om de vedlagte "score listene" ønskes.

Tabell 1 Veiledende oppsummering av identifisering mineraler. Se vedlagte scorelister for flere detaljer.

Prøve merket	Filnavn	Hoved mineral	Andre mulige mineraler	Merknader
Sv3R	DB01.sd	Calcitt	<i>Kvarts, monohydrat</i>	<i>Andre mulige mineraler foreligger i relativt sett små mengder. Tilstedeværelsen av disse er således usikker, og således er det mulig at andre enn de foreslåtte kan inngå.</i>
Sv10R	DB02.sd	Calcitt	<i>Glimmer(eks.zinvalditt)</i>	<i>Andre mulige mineraler foreligger i relativt sett små mengder. Tilstedeværelsen av disse er således usikker, og således er det mulig at andre enn de foreslåtte kan inngå.</i>
Sv15R	DB03.sd	Calcitt		



1475 counts  
width 1  
Sv3R  
1000

PEAK SEARCH:

Input file: DB01.SD  
 Start angle [°2θ]: 3.010  
 End angle [°2θ]: 69.990  
 Maximum d-value [Å]: 29.32891  
 Minimum d-value [Å]: 1.34314  
 Maximum number of counts: 338

Sample identification: Sv3R

Anode material: Cu  
 α1 Wavelength [Å]: 1.54060  
 α2 Wavelength [Å]: 1.54439

Intensity measured with FIXED slit

PEAK SEARCH:

d-value [Å]	Angle [°2θ]	Tip width [°2θ]	Height [counts]	Background [counts]	Rel. int. [%]	Significanc
6.18662	14.305	0.240	3	1	0.3	1.09
4.33518	20.470	0.160	6	2	0.5	0.82
3.84311	23.125	0.080	100	2	8.8	1.10
3.33671	26.695	0.160	8	3	0.7	0.88
3.02952	29.460	0.120	1142	3	100.0	9.64
2.83519	31.530	0.160	24	4	2.1	1.56
2.49074	36.030	0.100	123	4	10.8	2.51
2.28122	39.470	0.080	154	6	13.5	1.48
2.09135	43.225	0.080	151	8	13.2	2.05
2.08538	43.355	0.060	90	8	7.9	0.79
1.92484	47.180	0.080	49	9	4.3	1.29
1.90921	47.590	0.080	169	8	14.8	1.12
1.87150	48.610	0.080	149	8	13.0	0.91
1.82132	50.040	0.160	4	7	0.3	0.89
1.62336	56.655	0.100	30	5	2.6	1.27
1.60264	57.455	0.100	49	5	4.3	1.38
1.58277	58.245	0.200	6	5	0.5	0.83
1.52359	60.740	0.160	36	5	3.2	1.74
1.50580	61.535	0.160	15	5	1.3	0.82
1.47196	63.110	0.120	10	5	0.9	1.01
1.43818	64.770	0.080	45	6	3.9	1.44
1.42073	65.665	0.240	18	6	1.5	2.16
1.37940	67.895	0.040	4	6	0.4	0.84

SCORE LIST LONG:

Analysed DI file : C:\ANN\DATA\DB01.DI  
 Sample identification : Sv3R  
 Last update of results file : 14-feb-1997 14:31

Database used : C:\IDENTDB

MEASUREMENT PARAMETERS

Diffractometer : X'PERT  
 Start angle : 3.010  
 Final angle : 69.990  
 Step size : 0.020  
 Time per step : 1.0  
 Anode material : Cu  
 Focus : LFF

Date and time of measurement : 14-feb-1997 10:16

PEAK SEARCH PARAMETERS

Minimum peak width : 0.00  
 Maximum peak width : 1.00  
 Peak base width : 2.00  
 Minimum significance : 0.75

Number of peaks detected : 23

SEARCH-MATCH PARAMETERS

Number of strong lines of the reference patterns used in SEARCH : 5  
 Intensity threshold : 3.02  
 Confidence threshold : 10  
 Minimum specimen displacement : -100  
 Maximum specimen displacement : 100  
 Step size in specimen displacement: 25

Restrictions file : MINERAL

Results:

I	Card Id	#	Search score	Search displ	Match score	Rel m score	Conc	Displ	Formulas
1	* 24-0027	1	98	75	14.14	0.83	48	80	CaCO3
2	* 05-0586	1	96	100	16.29	0.81	58	127	CaCO3
3	06-0615	1	63	0	1.45	0.48	6	16	FeO
4	06-0680	1	42	-100	2.27	0.45	0	-77	Cu2-xSe



Name

Calcite  
Calcite, syn  
W\Plustite, syn  
Berzelianite

I	Card Id	#	Search score	Search displ	Match score	Rel m score	Conc	Displ	Formulas
5	18-0438	1	54	100	2.13	0.43	0	134	Cu <sub>1.4</sub> As <sub>0.9</sub> S <sub>2</sub>
6	12-0646	1	54	0	2.13	0.43	0	16	Cu <sub>3</sub> (As,V)S <sub>4</sub>
7	14-0567	1	40	-100	4.18	0.42	1	-77	(Sn,Fe)(O,OH) <sub>2</sub>
8	04-0835	1	50	75	1.23	0.41	3	109	NiO
9	04-0691	1	38	0	4.08	0.41	0	12	AgBiS <sub>2</sub>
10	* 33-1161	1	54	25	7.26	0.40	1	26	SiO <sub>2</sub>
11	37-1492	1	50	-100	6.29	0.39	1	-110	MoS <sub>2</sub>
12	34-0171	1	38	50	8.16	0.39	1	84	CuSe
13	06-0097	1	46	-75	5.39	0.39	1	-76	MoS <sub>2</sub>
14	26-1079	1	38	0	2.27	0.38	1	34	C
15	08-0012	1	42	-100	1.88	0.38	0	-84	AgSb <sub>2</sub> (O,OH,H <sub>2</sub> O) <sub>6</sub>
16	30-0093	1	54	0	10.69	0.37	1	34	Pd <sub>8</sub> Sb <sub>3</sub>
17	38-0346	1	38	25	2.93	0.37	0	37	Ni <sub>18</sub> Bi <sub>3</sub> AsS <sub>16</sub>
18	07-0007	1	42	100	2.55	0.36	1	110	NH <sub>4</sub> Cl
19	07-0321	1	42	0	2.18	0.36	0	16	PbxSb <sub>2</sub> -z(O,OH,H <sub>2</sub> O)
20	11-0121	1	38	-100	3.98	0.36	0	-90	Co <sub>3</sub> S <sub>4</sub>
21	27-0281	1	38	25	11.61	0.35	1	59	Pb <sub>2</sub> UO <sub>2</sub> (CO <sub>3</sub> ) <sub>3</sub>
22	18-0672	1	41	-75	6.68	0.35	0	-85	NaMnZn <sub>2</sub> (Ti,Fe) <sub>6</sub> Ti <sub>1</sub>
23	35-0582	1	42	100	4.90	0.35	13	134	Cu <sub>2</sub> FeSnS <sub>4</sub>
24	02-1231	1	63	-25	1.40	0.35	6	-26	CuZn
25	08-0002	1	47	-100	11.12	0.35	1	-119	UTi <sub>2</sub> O <sub>6</sub>
26	25-0597	1	38	100	9.01	0.35	0	105	Pd <sub>5</sub> Sb <sub>2</sub>
27	17-0546	1	47	0	11.33	0.34	3	16	KCa <sub>4</sub> F(Si <sub>2</sub> O <sub>5</sub> ) <sub>4</sub> !8H <sub>2</sub> O
28	12-0688	1	50	-100	9.88	0.34	9	-84	ZnS
29	39-1363	1	50	-100	9.88	0.34	9	-84	ZnS
30	30-0991	1	44	-75	5.09	0.34	0	-56	(K,Sr) <sub>2</sub> -xNb <sub>2</sub> O <sub>6</sub> (O,O
31	05-0418	1	5000	0	13.19	0.34	1	255	SrCO <sub>3</sub>
32	43-1469	1	40	100	3.72	0.34	0	134	Ni <sub>3</sub> S <sub>4</sub>
33	25-0298	1	42	-100	10.47	0.34	3	-130	Cu <sub>6</sub> Hg <sub>3</sub> As <sub>4</sub> S <sub>12</sub>
34	18-0972	1	75	-25	3.01	0.33	2	-20	PtS
35	24-0334	1	41	25	4.99	0.33	0	59	CoNi <sub>2</sub> S <sub>4</sub>
36	15-0445	1	46	-100	7.97	0.33	2	-88	Fe-Mg-Al-SiO
37	25-0264	1	46	-100	5.93	0.33	0	-84	Cu <sub>6</sub> As <sub>4</sub> S <sub>9</sub>
38	35-0581	1	38	-100	9.54	0.33	5	-84	Cu <sub>3</sub> SbS <sub>4</sub>
39	08-0237	1	38	-25	5.54	0.33	0	-37	WS <sub>2</sub>
40	29-0927	1	40	-25	7.10	0.32	0	-30	Ni <sub>9</sub> Sb <sub>2</sub> S <sub>8</sub>
41	14-0603	1	39	75	7.40	0.32	0	94	NaNbO <sub>3</sub>
42	16-0696	1	38	100	6.05	0.32	1	84	(Re,Ca,Fe) <sub>0.90</sub> NbTi
43	35-0706	1	39	0	5.69	0.32	0	16	(Na,Cs)Bi(Ta,Nb,Sb
44	29-0962	1	39	-100	11.68	0.32	1	-99	Pd <sub>10</sub> As <sub>4</sub> Bi
45	40-0492	1	38	0	29.74	0.31	3	10	AsTi <sub>3</sub> V <sub>4</sub> O <sub>13</sub> (OH)
46	05-0739	1	38	100	2.19	0.31	1	119	MgSiO <sub>3</sub> !H <sub>2</sub> O
47	29-0823	1	58	-100	10.55	0.31	1	-112	Na <sub>2</sub> KLi(Mn,Fe) <sub>2</sub> Ti <sub>2</sub> S
48	04-0728	1	41	25	5.56	0.31	1	59	Ca <sub>3</sub> Mg(SiO <sub>4</sub> ) <sub>2</sub>
49	13-0114	1	49	-100	6.42	0.31	1	-84	FeZn <sub>3</sub> Be <sub>3</sub> Si <sub>3</sub> O <sub>12</sub> S
50	33-0876	1	38	100	19.54	0.31	1	123	Mg <sub>3</sub> (PO <sub>4</sub> ) <sub>2</sub>
51	11-0508	1	39	25	11.53	0.30	0	15	Pd <sub>17</sub> Se <sub>15</sub>

SV3R

Name
Sinnerite
Arsenosulvanite
Cassiterite
Bunsenite, syn
Matildite
Quartz, syn
Molybdenite-2\ITH\RG, syn
Klockmannite, syn
Molybdenite-2H, syn
Graphite-3\ITR\RG, syn [NR]
Stetefeldtite
Mertieite-II, syn
Arsenochauchoecornite
Salammoniac, syn
Bindheimite
Linnaeite
Widenmannite
Landauite
Stannite
Zhanghengite, syn
Brannerite, heated
Stibiopalladinite
Fluorapophyllite
Wurtzite-10\ITH\RG
Wurtzite-8\ITH\RG
Kaliprochloro
Strontianite, syn
Polydymite
Aktashite
Cooperite, syn
Siegenite, syn
Unnamed mineral [NR]
Sinnerite, syn
Famatinite, syn
Tungstenite-2\ITH\RG, syn
Tucekite
Lueshite, syn
Aeschynite
Natrobismantite
Palladobismutharsenide, syn
Tomichite
Cerolite
Manganneptunite
Merwinite, syn
Genthelvite
Farringtonite, syn
Palladseite, syn

	I	Card Id	#	Search score	Search displ	Match score	Rel m score	Conc	Displ	Formulas
52		27-0250	1	38	0	21.53	0.30	1	34	Fe3(PO4)2
53		42-0423	1	38	50	12.42	0.30	3	40	CaTiO3
54		13-0550	1	39	0	5.68	0.30	0	16	Fe2Ti14O24
55		35-0525	1	44	100	4.48	0.30	0	119	(Cu, Zn, Ag)12(As, Sb)
56		19-0392	1	42	-50	5.08	0.30	1	-40	Cu2(WO4)(OH)2
57		25-0279	1	43	50	2.99	0.30	0	80	CuGaS2
58		35-1351	1	38	100	5.94	0.30	6	134	Cu2FeSnS4
59		03-0503	1	39	100	7.10	0.30	1	110	3NaAlSiO4:CaCO3
60		15-0682	1	52	-75	5.91	0.30	0	-65	PbSeO4
61		22-0349	1	40	0	2.65	0.29	5	23	(Fe, Mg)(Cr, Fe)2O4
62		38-0430	1	42	-100	2.94	0.29	1	-66	MnFe2O4
63		39-1469	1	39	75	12.25	0.29	0	109	Ba0.75Sr0.25SO4
64		29-0808	1	39	0	8.45	0.29	1	10	Li1-x(Fe, Mn)PO4
65		02-1257	1	46	-25	2.91	0.29	0	-20	CaCe2(CO3)3F2
66		21-1276	1	39	-25	3.48	0.29	3	-20	TiO2
67		42-1449	1	41	25	3.47	0.29	1	24	FeNi2S4
68		22-1135	1	50	-50	13.69	0.29	0	-55	Pb2(SO4)(SeO4)
69	*	29-0306	1	46	-25	10.24	0.28	0	-20	CaCO3:H2O
70		05-0283	1	54	50	9.10	0.28	1	62	CaAl(F, OH)5:H2O
71		05-0417	1	38	-75	9.66	0.28	2	-74	PbCO3
72		13-0583	1	40	-100	6.24	0.28	1	-110	CaHAsO4:2H2O
73		27-0034	1	38	0	10.96	0.28	0	27	BaCe2(CO3)3F2
74		38-0478	1	41	-100	5.01	0.28	1	-70	Mg6Al2(OH)18:4H2O
75		08-0003	1	55	-100	5.00	0.28	0	-90	Pb5Au(Te, Sb)4Sx
76		14-0345	1	38	0	8.60	0.28	0	34	Ce-La-Sr-Ca-C-O-H!
77		29-0384	1	38	0	8.60	0.28	0	34	(Sr, Ca)(La, Ce)(CO3)
78		12-0277	1	38	-100	3.88	0.28	0	-84	(La, Y)PO4:H2O
79		29-1420	1	39	0	6.35	0.28	0	34	Cu3PbTeO4(OH)6
80		11-0312	1	67	25	15.68	0.27	2	15	(K, Na)Cu7AlSi9O24(
81		41-1498	1	50	50	6.83	0.27	1	80	(Cu10Hg2)Sb4S12
82		05-0737	1	39	50	5.12	0.27	0	38	(Sr, Ca)Al3(PO4)(SO
83		21-0974	1	57	-100	11.01	0.27	1	-84	NiSO4:H2O
84		28-2002	1	38	-25	15.03	0.27	0	-35	C6H4(CO)2C6H4
85		12-0625	1	39	0	10.70	0.27	0	16	Mg2(Al, Fe)6(SiO4)4
86		39-0326	1	47	-50	8.48	0.27	1	-40	CuPb8(AsO3)2O3Cl5
87		20-0240	1	38	0	14.27	0.26	0	12	Ca5(SiO4)2(OH)2
88		27-0669	1	39	-50	15.86	0.26	1	-40	Na6(Ca, Mn)(Ti, Fe)S
89		41-1431	1	39	100	6.34	0.26	1	130	CaWO4
90		37-0414	1	38	-100	16.03	0.26	0	-105	Ca5(SiO4)2(OH, F)2
91		29-0322	1	39	-75	7.61	0.26	0	-59	Ca2F2CO3
92		31-0462	1	46	100	5.50	0.26	11	123	Cu4FeZnSn2S8
93		33-0188	1	43	0	18.65	0.26	1	27	Ba4(V, Ti)4ClSi8B2O
94		30-0263	1	39	50	12.65	0.26	0	40	(Ca, Ce)(Ti, Fe, Cr, M
95		11-0491	1	42	100	5.38	0.26	0	116	(Fe, Mn)4Be3Si3O12S
96		24-0730	1	43	50	8.70	0.26	1	77	(Mn, Fe)AlPO4(OH, O)
97		26-0292	1	38	-100	4.58	0.25	2	-81	Ca3Al2(SiO4)3
98		11-0695	1	46	0	7.39	0.25	8	16	SiO2

Sv 3R

VIII

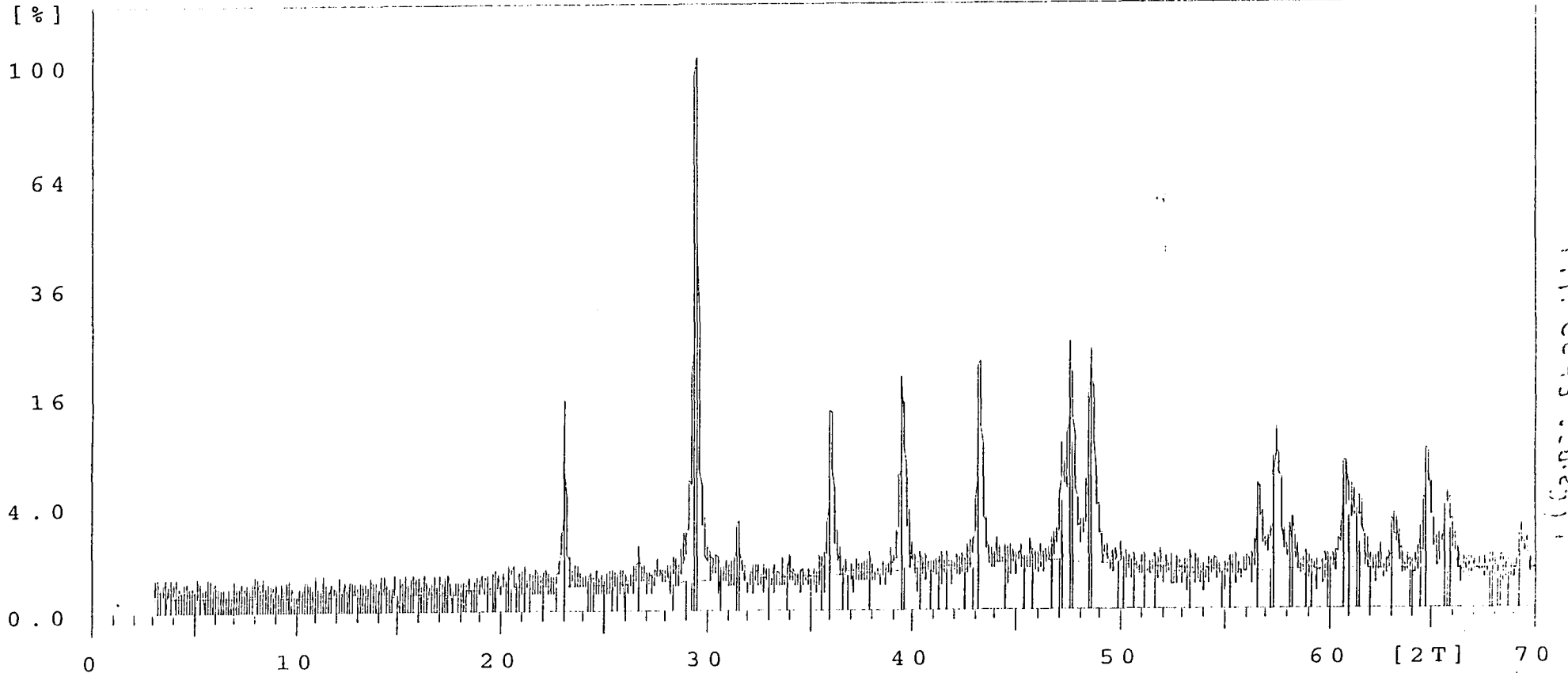
=====

NGU-lab XRD-analyser Philips X'Pert MPD Operator: Andreas Grimstvedt \*\*\*\*\*

+-----+  
Name+-----+  
Graftonite, syn  
Perovskite, syn  
Avicennite  
Giraudite  
Cuprotungstite, syn  
Gallite, syn  
Stannite, syn  
Cancrinite  
Kerstenite, syn [NR]  
Donathite  
Iwakiite  
Barite, strontian, syn  
Ferrisicklerite  
Parisite-(Ce)  
Rutile, syn  
Violarite  
Olsacherite  
Monohydrocalcite  
Gearksutite  
Cerussite, syn  
Pharmocolite  
Cordylite-(Ce)  
Meixnerite  
Nagyagite, syn  
Ancylite  
Ancylite-(Ce)  
Rhabdophane-(La)  
Parakhinite  
Ajoite  
Tetrahedrite, mercurian, syn  
Svanbergite, calcian  
Dwornikite, syn  
Hoelite, syn  
Yoderite  
Freedite  
Calcio-chondrodite, syn  
Koashvite  
Scheelite, syn  
Reinhardbraunsite, fluorian  
Brenkite  
Stannite, syn  
Nagashimalite  
Loveringite  
Danalite  
Ernstite  
Grossular  
Cristobalite, syn  
+-----+

Sample ident.: Sv3R

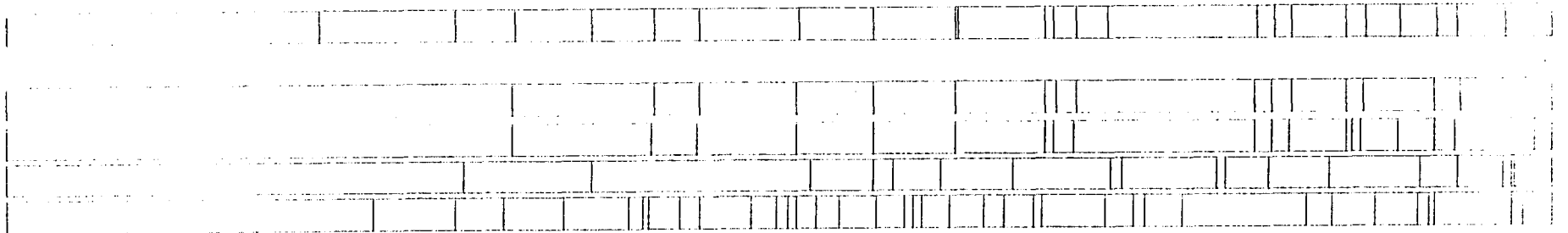
14-feb-1997 14:31

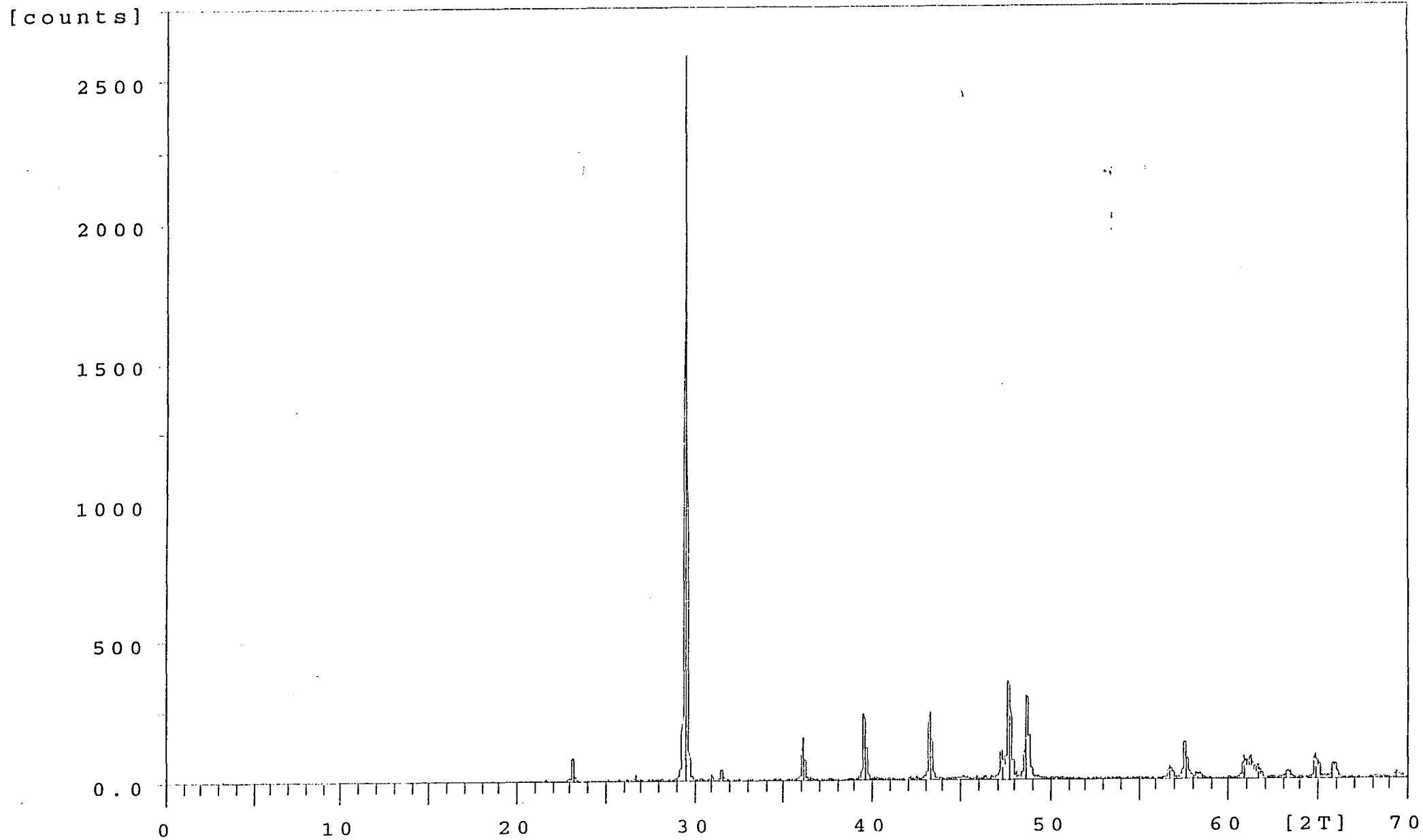


1997-02-14 14:31

DB01

Calcite  
Calcite,  
Quartz,  
Monohydr





DB02.SD

DB02.DI

1997-02-14 14:31  
Sv10R

xi

PEAK SEARCH:

Input file: DB02.SD  
 Start angle [°2θ]: 3.010  
 End angle [°2θ]: 69.990  
 Maximum d-value [Å]: 29.32891  
 Minimum d-value [Å]: 1.34314  
 Maximum number of counts: 508

Sample identification: Sv10R

Anode material: Cu  
 α1 Wavelength [Å]: 1.54060  
 α2 Wavelength [Å]: 1.54439

Intensity measured with FIXED slit

PEAK SEARCH:

d-value [Å]	Angle [°2θ]	Tip width [°2θ]	Height [counts]	Background [counts]	Rel. int. [%]	Significanc
9.92238	8.905	0.120	11	1	0.4	0.77
3.83656	23.165	0.120	79	2	3.1	3.09
3.34039	26.665	0.120	15	3	0.6	1.96
3.02551	29.500	0.160	2581	3	100.0	35.18
2.88699	30.950	0.080	15	3	0.6	0.76
2.83519	31.530	0.060	31	3	1.2	0.86
2.48773	36.075	0.100	104	4	4.0	2.83
2.27845	39.520	0.120	164	5	6.3	4.08
2.08928	43.270	0.080	164	7	6.3	1.85
1.92177	47.260	0.080	42	8	1.6	0.80
1.90751	47.635	0.100	219	7	8.5	2.85
1.87006	48.650	0.160	180	7	7.0	7.52
1.71840	53.265	0.100	3	5	0.1	0.85
1.62178	56.715	0.100	26	5	1.0	0.76
1.59984	57.565	0.080	67	5	2.6	1.22
1.58364	58.210	0.160	9	5	0.3	0.75
1.52144	60.835	0.100	42	5	1.6	1.37
1.50305	61.660	0.240	17	5	0.7	1.57
1.47784	62.830	0.060	3	5	0.1	0.81
1.47018	63.195	0.160	12	5	0.4	0.78
1.43779	64.790	0.140	40	5	1.5	2.87
1.41843	65.785	0.200	24	5	0.9	1.89
1.35311	69.400	0.060	11	6	0.4	0.84

SCORE LIST LONG:

Analysed DI file : C:\ANN\DATA\DB02.DI  
 Sample identification : Sv10R  
 Last update of results file : 14-feb-1997 14:45  
 Database used : C:\IDENTDB

MEASUREMENT PARAMETERS

Diffractionmeter : X'PERT  
 Start angle : 3.010  
 Final angle : 69.990  
 Step size : 0.020  
 Time per step : 1.0  
 Anode material : Cu  
 Focus : LFF

Date and time of measurement : 14-feb-1997 11:13

PEAK SEARCH PARAMETERS

Minimum peak width : 0.00  
 Maximum peak width : 1.00  
 Peak base width : 2.00  
 Minimum significance : 0.75

Number of peaks detected : 23

SEARCH-MATCH PARAMETERS

Number of strong lines of the reference patterns used in SEARCH : 5  
 Intensity threshold : 3.02  
 Confidence threshold : 10  
 Minimum specimen displacement : -100  
 Maximum specimen displacement : 100  
 Step size in specimen displacement: 25

Restrictions file : MINERAL

Results:

I	Card Id	#	Search score	Search displ	Match score	Rel m score	Conc	Displ	Formulas
1	* 05-0586	1	83	100	15.99	0.80	33	134	CaCO3
2	24-0027	1	94	100	13.56	0.80	23	110	CaCO3
3	* 29-1489	1	58	100	2.99	0.50	0	119	Al2Si2O5(OH)4·2H2O
4	06-0615	1	54	50	1.43	0.48	4	84	FeO

I	Card Id #	Search score	Search displ	Match score	Rel m score	Conc	Displ	Formulas
5	06-0680	1	38	-100	2.04	0.41	1	Cu2-xSe
6	04-0832	1	46	0	8.33	0.38	0	FeS
7	04-0763	1	48	100	2.62	0.37	1	NiCr2O4
8	04-0835	1	54	-25	1.12	0.37	0	NiO
9	30-0093	1	50	0	10.47	0.36	0	PdSb3
10	05-0739	1	63	100	2.52	0.36	0	MgSiO3.H2O
11	11-0151	1	46	0	5.70	0.34	0	FeS
12	02-1008	1	46	-25	4.54	0.32	0	(Mg,Fe)3Al2Si3O12
13	12-0688	1	42	50	9.35	0.32	3	ZnS
14	39-1363	1	42	50	9.35	0.32	3	ZnS
15	36-1451	1	46	-100	2.89	0.32	0	ZnO
16	29-0823	1	62	50	10.76	0.32	0	Na2KLi(Mn,Fe)2Ti2S
17	24-0080	1	46	100	5.74	0.30	2	FeS
18	43-0664	1	39	-25	2.08	0.30	1	Ni3Si4O10(OH)2.15H2
19	05-0664	1	44	-100	2.66	0.30	0	ZnO
20	09-0353	1	42	25	2.04	0.29	0	(Mg,Fe)(Cr,Al)2O4
21	33-0670	1	40	50	4.63	0.29	1	Fe2P
22	10-0469	1	42	0	9.73	0.29	1	Cu3(Ge,Fe)54
23	29-0718	1	42	-75	0.86	0.29	2	(Pt,Fe)
24	10-0496	1	55	100	6.53	0.28	0	KAlV2Si3O10(OH)2
25	15-0506	1	43	50	2.78	0.28	0	(UO2)2ALOH(VO4)2.18
26	19-0607	1	42	100	10.36	0.27	0	(Fe,Mg)SiO3
27	31-0635	1	42	100	10.36	0.27	0	(Fe,Mg)SiO3
28	25-0432	1	39	100	12.20	0.27	1	Pb10AgBi5S18
29	11-0695	1	54	0	7.86	0.27	4	SiO2
30	17-0148	1	52	100	6.21	0.27	0	Mg(UO2)2(AsO4)2.110
31	40-0492	1	40	0	25.42	0.27	1	AsTi3V4O13(OH)
32	42-0604	1	42	0	14.41	0.27	0	K(AlFeLi)(Si3Al)O1
33	02-1231	1	42	-75	1.06	0.27	1	Cu2
34	42-0565	1	42	50	4.21	0.26	5	Cu2FeGeS4
35	33-1144	1	67	-50	3.65	0.26	1	RuAs2
36	17-0546	1	49	75	8.50	0.26	1	KCa4F(Si2O5)4.8H2O
37	39-0359	1	58	25	3.34	0.26	1	Mg3(SO4)2(OH)2
38	43-0692	1	50	75	11.05	0.26	0	KLi(Al,Li)2(Si3Al)
39	31-0462	1	42	100	5.36	0.26	5	Cu4FeZnSn2S8
40	13-0550	1	40	25	4.85	0.26	1	Fe2Ti14O24
41	39-1425	1	58	0	7.00	0.25	4	SiO2
42	11-0515	1	54	100	1.48	0.25	2	Cu17.6Fe17.6S32
43	18-1085	1	44	-75	3.66	0.24	0	NaMgAlF6.H2O
44	22-1114	1	39	-75	4.81	0.24	0	(Nb,W,Fe,Mn)O2
45	16-0344	1	46	0	5.77	0.24	1	KMg3(Si3Al)O10F2
46	37-0477	1	42	100	8.92	0.23	1	FeS
47	35-0640	1	50	-50	13.84	0.23	1	Y2(SiO4)(CO3)
48	12-0291	1	39	75	5.83	0.23	0	FeSe2
49	25-0298	1	46	-100	7.20	0.23	1	Cu6Hg3As4S12
50	27-0166	1	46	-100	1.37	0.23	3	CuFe2S3
51	25-0843	1	39	-100	5.93	0.23	0	Mg3Fe2(SiO4)3

Name  
 Calcite, syn  
 Calcite  
 Halloysite-10A  
 W\Pluistite, syn

1470043 valley 2 6(9)

NGU-lab XRD-analyser Philips X'Pert MPD Operator: Andreas Grimstvedt \*\*\*\*\*

Name
Berzelianite
Troilite, syn
Nichromite, syn [NR]
Bunsenite, syn
Mertieite-II, syn
Cerolite
Troilite
Pyrope
Wurtzite-10\ITM\RG
Wurtzite-8\ITM\RG
Zincite, syn
Manganneptunite
Troilite-2\ITM\RG, syn
Pimelite
Zincite, syn
Magnesiochromite, ferroan
Barringerite, syn
Germanite
Platinum, ferroan
Roscoelite-1\ITM\RG
Vanuralite
Orthoferrosilite, magnesian
Ferrosilite, magnesian
Heyrovskiyite, syn
Cristobalite, syn
Novacekite-20A, syn
Tomichite
Zinnwaldite-1\ITM\RG
Zhanghengite, syn
Briartite, syn
Anduocite
Fluorapophyllite
Caminite
Lepidolite-2\ITM\RG#2
Stannite, syn
Avicennite
Cristobalite, syn
Talnakhite, syn
Ralstonite
Ashanite
Phlogopite-1\ITM\RG, syn
Troilite-2\ITM\RG
Ilmorite-(Y)
Ferroselite, syn
Aktashite
Isocubanite, syn
Majorite

1470043 valley 2 7(9)

NGU-lab XRD-analyser Philips X'Pert MPD Operator: Andreas Grimstvedt \*\*\*\*\*

I	Card Id	#	Search score	Search displ	Match score	Rel m score	Conc	Displ	Formulas
52	05-0453	1	5000	0	7.07	0.23	1	-134	CaCO3
53	35-0683	1	39	-100	3.17	0.23	0	-88	Cu6FeSn2S8
54	08-0028	1	42	50	1.58	0.23	1	62	PbTe
55	14-0657	1	42	50	7.37	0.22	2	51	Ba4Ti7NbS14O28Cl
56	16-0695	1	50	-100	5.88	0.22	0	-66	Ca5Al2(OH)4Si3O12
57	42-1392	1	49	75	2.39	0.22	0	94	Ni(As,S)2
58	41-0582	1	39	50	7.39	0.22	0	38	Na4Ca6Mn19Li2Al8(P
59	30-0444	1	47	-100	4.13	0.22	1	-88	(Co,Ni,Fe)9S8
60	18-0438	1	46	100	1.08	0.22	0	134	Cu1.4As0.9S2
61	05-0418	1	5000	0	8.43	0.22	1	-29	SrCO3
62	22-1121	1	45	-100	13.37	0.22	0	-77	SrTi21O38
63	26-1380	1	39	100	5.39	0.22	0	134	NbNbO3
64	11-0508	1	39	-25	8.11	0.21	0	-55	Pd17Se15
65	15-0800	1	40	100	17.07	0.21	1	119	Na2CO311OH2O
66	30-0486	1	42	-100	4.21	0.21	2	-84	Cu6Fe23SnS8
67	37-0471	1	54	100	4.20	0.21	1	119	CuFeS2
68	18-0889	1	48	-100	1.89	0.21	4	-123	Ni3Se4
69	20-0222	1	39	-100	8.53	0.21	1	-123	(Ca,Pb)ZnSiO4
70	20-0587	1	41	100	9.56	0.21	1	95	Pb2Mn2(Si2O7)O2
71	25-0597	1	42	100	5.38	0.21	0	95	Pd5Sb2
72	42-1414	1	42	75	6.55	0.20	0	74	K(Mg,Fe+2)3(Al,Fe+
73	18-0877	1	46	-75	3.48	0.20	1	-52	(Fe,Ni)
74	39-0368	1	39	0	5.50	0.20	1	27	Ca3Al2(SiO4)3
75	35-0558	1	39	0	3.87	0.20	0	12	(Ni,Fe)SO41H2O
76	22-0990	1	39	25	6.65	0.20	0	35	Y2FeBe2Si2O10
77	12-0646	1	63	25	0.99	0.20	2	26	Cu3(As,V)S4
78	35-0604	1	39	-75	5.14	0.20	1	-52	Na2Zr2Be(PO4)4
79	26-0832	1	47	25	9.09	0.20	1	48	Pb2CrO5
80	25-0288	1	50	25	2.37	0.20	1	59	CuFeS2
81	16-0155	1	39	-75	4.53	0.20	0	-87	Ca3ZrSi2O9
82	08-0001	1	56	100	2.84	0.19	1	81	(Cu,Co,Ni)3Se4
83	36-0395	1	46	0	5.01	0.19	6	16	Cu13Ge2Fe2S16
84	29-0322	1	39	0	5.37	0.19	0	12	Ca2F2CO3
85	29-0210	1	58	100	3.49	0.18	0	119	Ba(VO2)2(AsO4)211O
86	17-0644	1	40	50	2.73	0.18	0	80	CuRh2S4
87	39-1369	1	39	-100	6.32	0.18	0	-99	Ca2MgC2(SiO4)(Si2
88	38-0399	1	41	100	6.28	0.18	0	116	Hg3(AsO4)
89	41-0591	1	58	100	2.99	0.18	9	105	Cu10Fe4ZnGe2S16
90	25-0294	1	54	-25	6.28	0.17	0	-52	CuPbTlAs2S5
91	15-0378	1	42	-100	5.38	0.17	0	-88	Be2AsO4(OH)4H2O
92	15-0069	1	39	-75	2.60	0.17	0	-52	(Y,Ca,Na)F2
93	18-0670	1	42	100	5.40	0.17	2	123	MgFeAlSiBO
94	02-0520	1	39	25	2.84	0.17	0	20	(Mg,Fe)SiO3
95	21-0974	1	54	-50	6.67	0.16	2	-51	NiSO41H2O
96	39-0326	1	45	0	5.20	0.16	0	12	CuPb8(AsO3)2O3Cl5
97	07-0377	1	46	-50	4.71	0.16	0	-51	Fe-Te-O
98	25-0287	1	42	0	3.41	0.16	1	34	Cu9(Fe,Ni)8S16

SV10R



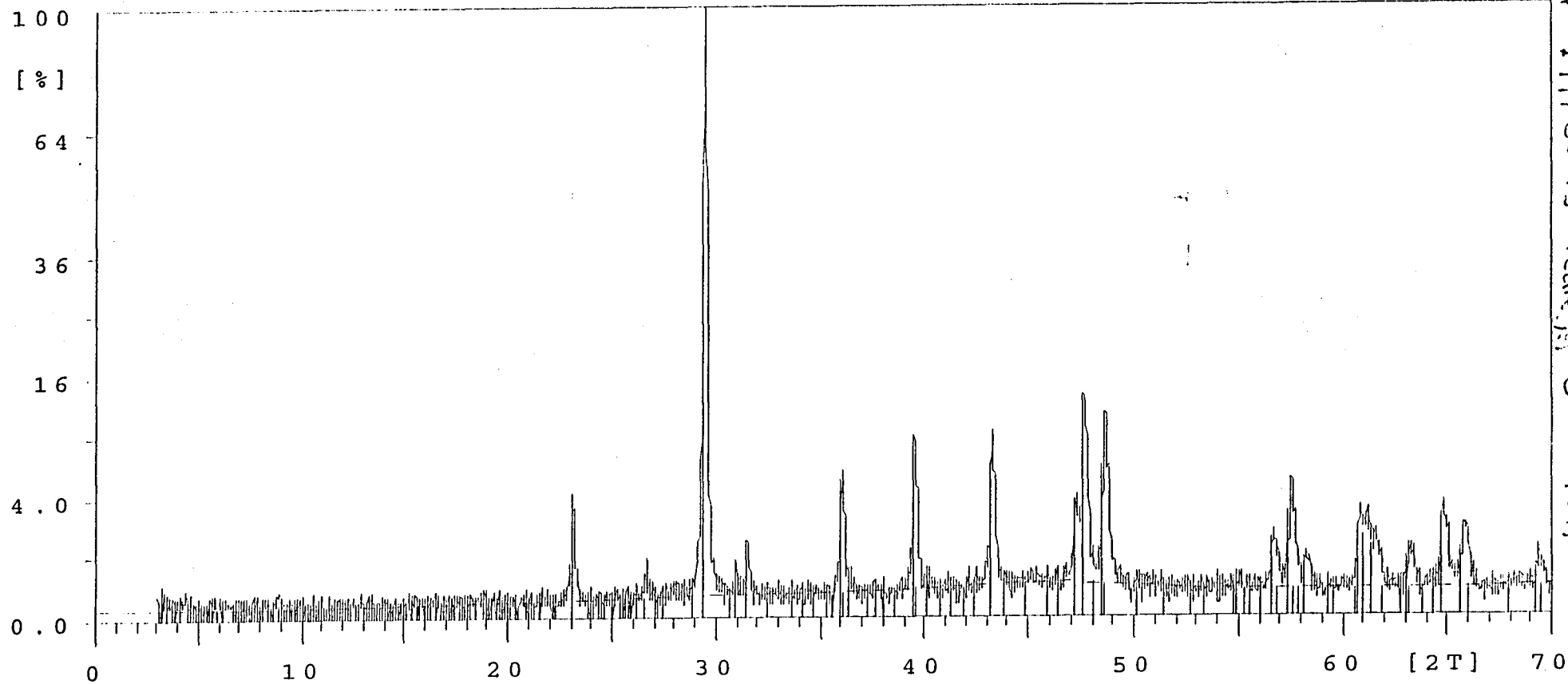
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Graphite
Cobaltpentlandite
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Crichtonite
Natroniobite
Palladseite, syn
Natron, syn
Mawsonite
Chalcopyrite
Tr\Plustedtite
Esperite, syn
Kentrolite, syn
Stibiopalladinite
Biotite-1\ITM\RG
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Grossular
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Gadolinite-(Y), heated
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Gainesite
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Baghdadite, syn
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Germanite
Brenkite
Heinrichite, syn
Cuprorhodsite, syn
Shuiskite
Chursinite
Renierite
Wallisite
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Freedite
Blakeite
Talnakhite

SV10R

XV

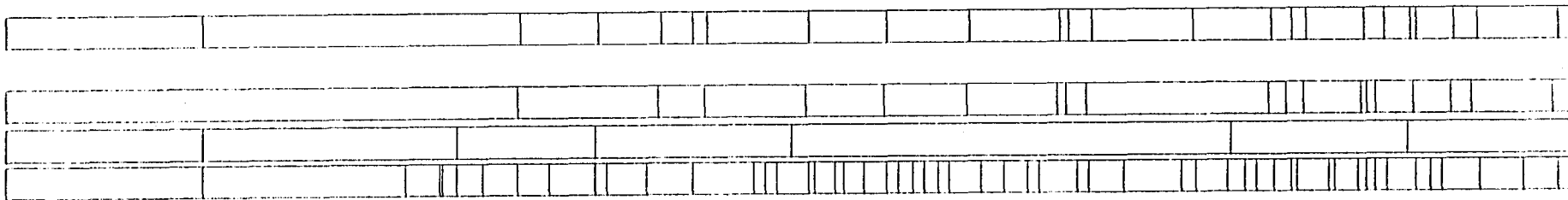
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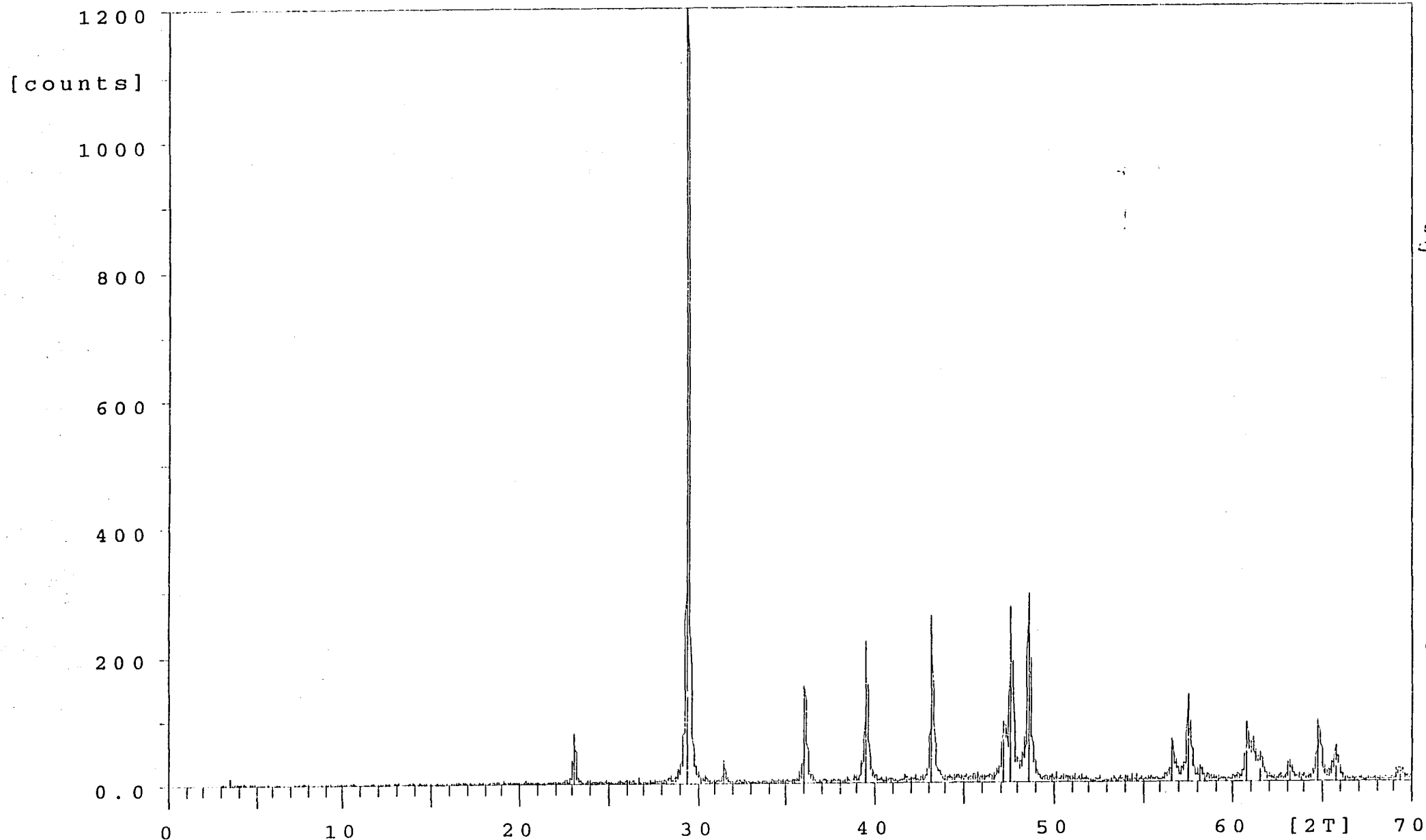


DB02

Calcite,  
Halloysi  
Zinnwald



Sv10R  
xvi



19970213 Vedlegg 3

1 (1)

Sv15R

DB03.SD

DB03.DI

File: C:\ANN\DATA\DB03.DI 14-feb-1997 14  
 NGU-lab XRD-analyser Philips X'Pert MPD Operator: Andreas Grimstvedt \*\*\*\*\*

PEAK SEARCH:

Input file: DB03.SD  
 Start angle [°2θ]: 3.010  
 End angle [°2θ]: 69.990  
 Maximum d-value [Å]: 29.32891  
 Minimum d-value [Å]: 1.34314  
 Maximum number of counts: 347  
 Sample identification: Sv15R  
 Anode material: Cu  
 λ1 Wavelength [Å]: 1.54060  
 λ2 Wavelength [Å]: 1.54439

Intensity measured with FIXED slit

PEAK SEARCH:

d-value [Å]	Angle [°2θ]	Tip width [°2θ]	Height [counts]	Background [counts]	Rel. int. [%]	Significanc
25.15205	3.510	0.060	12	0	1.0	0.91
3.84721	23.100	0.120	79	2	6.6	2.44
3.03003	29.455	0.140	1204	4	100.0	14.93
2.84002	31.475	0.060	31	3	2.6	1.11
2.49174	36.015	0.080	125	4	10.4	1.53
2.28150	39.465	0.080	164	5	13.6	1.37
2.09158	43.220	0.060	180	7	14.9	0.91
1.92369	47.210	0.120	58	8	4.8	1.03
1.91015	47.565	0.080	164	8	13.6	1.65
1.87241	48.585	0.100	161	8	13.4	2.91
1.62388	56.635	0.120	26	5	2.2	0.79
1.60251	57.460	0.080	62	5	5.2	1.43
1.58401	58.195	0.200	7	5	0.6	0.96
1.52336	60.750	0.120	42	5	3.5	2.12
1.50679	61.490	0.200	18	5	1.5	0.80
1.47143	63.135	0.160	13	5	1.1	1.14
1.43868	64.745	0.100	44	5	3.6	1.27
1.41987	65.710	0.160	19	5	1.6	1.06
1.35457	69.315	0.320	6	5	0.5	1.60

File: C:\ANN\DATA\DB03.IDN 14-feb-1997 15  
 NGU-lab XRD-analyser Philips X'Pert MPD Operator: Andreas Grimstvedt \*\*\*\*\*

SCORE LIST LONG:

Analysed DI file : C:\ANN\DATA\DB03.DI  
 Sample identification : Sv15R  
 Last update of results file : 14-feb-1997 15:00

Database used : C:\IDENTDB

MEASUREMENT PARAMETERS

Diffractometer : X'PERT  
 Start angle : 3.010  
 Final angle : 69.990  
 Step size : 0.020  
 Time per step : 1.0  
 Anode material : Cu  
 Focus : LFF

Date and time of measurement : 14-feb-1997 12:09

PEAK SEARCH PARAMETERS

Minimum peak width : 0.00  
 Maximum peak width : 1.00  
 Peak base width : 2.00  
 Minimum significance : 0.75

Number of peaks detected : 19

SEARCH-MATCH PARAMETERS

Number of strong lines of the reference patterns used in SEARCH : 5  
 Intensity threshold : 3.02  
 Confidence threshold : 10  
 Minimum specimen displacement : -100  
 Maximum specimen displacement : 100  
 Step size in specimen displacement: 25

Restrictions file : MINERAL

Results:

I	Card Id	#	Search score	Search displ	Match score	Rel m score	Conc	Displ	Formulas
1	* 05-0586	1	97	100	17.47	0.87	58	119	CaCO3
2	24-0027	1	97	25	13.86	0.82	48	35	CaCO3
3	06-0615	1	63	-25	1.46	0.49	8	-20	FeO
4	18-0438	1	54	100	2.13	0.43	0	110	Cu1.4As0.9S2

Name
Calcite, syn
Calcite
W\Piustite, syn
Sinnerite

I	Card Id	#	Search score	Search displ	Match score	Rel m score	Conc	Displ	Formulas
5	07-0007	1	42	100	2.51	0.36	1	110	NH4Cl
6	08-0012	1	38	-100	1.76	0.35	0	-84	AgSb2(O,OH,H2O)6
7	35-0582	1	42	100	4.81	0.34	11	95	Cu2FeSnS4
8	41-0602	1	50	100	1.03	0.34	1	134	(Ir,Os,Ru)
9	20-0133	1	44	100	6.09	0.34	0	105	(Na,K)2(Ba,Ca,Sr)2
10	04-0836	1	33	-75	0.65	0.33	16	-74	Cu
11	42-0423	1	38	0	13.26	0.32	3	34	CaTiO3
12	04-0728	1	43	-50	5.67	0.31	1	-31	Ca3Mg(SiO4)2
13	16-0696	1	38	100	5.96	0.31	1	105	(Re,Ca,Fe)0.90NbTi
14	38-0430	1	46	-100	3.12	0.31	1	-66	MnFe2O4
15	05-0739	1	42	100	2.18	0.31	1	119	MgSiO3!H2O
16	06-0598	1	42	-50	0.93	0.31	3	-27	Ir
17	04-0835	1	42	0	0.89	0.30	4	23	NiO
18	14-0603	1	39	25	6.71	0.29	0	35	NaNbO3
19	33-0126	1	39	-50	7.29	0.29	0	-38	NiAs2
20	18-1484	1	41	100	4.07	0.29	0	105	ZnMn2O4
21	25-1426	1	38	-75	4.06	0.29	1	-41	(Cu,Fe)8(Pb,Cd)S8
22	11-0014	1	38	0	6.89	0.29	1	16	NiAs2
23	24-0518	1	40	0	6.59	0.29	0	16	NiAs2
24	21-1276	1	39	-100	3.43	0.29	4	-66	TiO2
25	18-0658	1	38	-100	2.53	0.28	1	-66	Fe2TiO4
26	04-0763	1	43	25	1.95	0.28	1	35	NiCr2O4
27	05-0737	1	36	-25	5.26	0.28	0	-20	(Sr,Ca)Al3(PO4)(SO
28	17-0546	1	43	0	9.05	0.27	3	16	KCa4F(Si2O5)4!8H2O
29	19-0392	1	38	0	4.66	0.27	0	27	Cu2(WO4)(OH)2
30	30-0093	1	54	0	7.86	0.27	1	34	Pd8Sb3
31	31-0462	1	46	50	5.60	0.27	11	66	Cu4FeZnSn2S8
32	41-1431	1	38	50	6.36	0.26	1	84	CaWO4
33	11-0695	1	54	-75	7.68	0.26	8	-85	SiO2
34	19-0750	1	38	50	13.65	0.26	0	84	(Mg4Al4)(Al4Si2)O2
35	29-0133	1	42	-25	5.75	0.26	0	-48	(Pt,Pd)4Sb3
36	38-0478	1	38	75	4.68	0.26	0	65	Mg6Al2(OH)18!4H2O
37	39-0326	1	43	-100	8.30	0.26	0	-123	CuPb8(AsO3)2O3Cl5
38	08-0001	1	41	100	3.89	0.26	1	130	(Cu,Co,Ni)3Se4
39	39-0329	1	42	-100	3.11	0.26	0	-99	CuIr2S4
40	25-0264	1	42	-100	4.63	0.26	0	-84	Cu6As4S9
41	18-0672	1	37	-100	4.89	0.26	0	-123	NaMnZn2(Ti,Fe)6Ti1
42	* 31-1082	1	39	100	7.20	0.26	0	116	KNaCaMg5Si8O22(OH)
43	09-0287	1	43	100	3.06	0.26	1	112	((Pt,Pd,Ni)S)
44	18-1362	1	38	25	3.06	0.25	1	59	Th(Ca,Ce)(CO3)2F2!
45	39-1425	1	58	0	7.04	0.25	11	16	SiO2
46	29-0322	1	38	-75	7.28	0.25	0	-59	Ca2F2CO3
47	25-1379	1	37	50	7.23	0.25	1	77	MgAlBO4
48	09-0423	1	33	-25	1.99	0.25	0	9	CuFeS2
49	25-0598	1	38	-50	6.67	0.25	0	-27	Pd11(Sb,As)4
50	29-0914	1	39	75	3.71	0.25	1	85	MoSe2
51	35-1351	1	38	100	4.93	0.25	5	105	Cu2FeSnS4

Name
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Stannite
Iridium, osmian
Barytolamprophyllite
Copper, syn
Perovskite, syn
Merwinite, syn
Aeschynite
Iwakiite
Cerolite
Iridium, syn
Bunsenite, syn
Lueshite, syn
Rammelsbergite
Hetaerolite
Shadlunite
Rammelsbergite, syn
Rammelsbergite
Rutile, syn
Ulvospinel, syn
Nichromite, syn [NR]
Svanbergite, calcian
Fluorapophyllite
Cuprotungstite, syn
Mertieite-II, syn
Stannite, syn
Scheelite, syn
Cristobalite, syn
Sapphirine-2\ITM\RG
Genkinite
Meixnerite
Freedite
Tyrrellite
Cuproiridsite
Sinnerite, syn
Landauite
Richterite, potassian
Cooperite
Thorbastn\PIasite
Cristobalite, syn
Brenkite
Sinhalite, syn
Chalcopyrite
Mertieite-I
Drysdallite-2\ITH\RG
Stannite, syn

	I	Card Id	#	Search score	Search displ	Match score	Rel m score	Conc	Displ	Formulas
52		26-0292	1	37	-100	4.43	0.25	2	-81	Ca3Al2(SiO4)3
53		22-1092	1	37	25	14.08	0.24	0	59	CuTeO3
54		22-0349	1	37	-75	2.18	0.24	4	-74	(Fe,Mg)(Cr,Fe)2O4
55		35-0706	1	40	0	4.36	0.24	0	16	(Na,Cs)Bi(Ta,Nb,Sb
56		03-0801	1	34	-50	3.86	0.24	0	-40	Ca3Al2(SiO4,CO3,OH
57		22-0729	1	40	-25	1.93	0.24	1	9	Hg(S,Se)
58		21-0974	1	38	-100	9.73	0.24	1	-116	NiSO4·H2O
59		05-0623	1	35	100	5.93	0.24	0	88	Fe-La-U-Ca-Ti-O-OH
60		29-1344	1	38	0	4.94	0.24	11	16	Tl2S
61		34-1316	1	37	25	4.70	0.23	0	59	(Ca,Ce)(Nb,Ta,Ti)2
62		11-0515	1	46	0	1.40	0.23	5	34	Cu17.6Fe17.6S32
63		40-0492	1	40	-100	22.14	0.23	2	-99	AsTi3V4O13(OH)
64		08-0003	1	43	-100	4.13	0.23	0	-84	Pb5Au(Te,Sb)4Sx
65		25-0774	1	34	50	7.06	0.23	1	40	Mn6CaFeAl(PO4)6F2
66		08-0028	1	42	-25	1.56	0.22	2	-24	PbTe
67		29-0823	1	55	-100	7.53	0.22	1	-66	Na2KLi(Mn,Fe)2Ti2S
68		42-1345	1	53	-25	5.25	0.22	1	-24	CoAsS
69		05-0453	1	5000	0	6.71	0.22	1	296	CaCO3
70		35-0525	1	43	100	3.12	0.21	0	134	(Cu,Zn,Ag)12(As,Sb
71		41-1498	1	46	0	5.18	0.21	1	34	(Cu10Hg2)Sb4S12
72		05-0418	1	5000	0	8.03	0.21	1	-100	SrCO3
73		20-0669	1	39	0	13.38	0.21	6	16	MgCO3·3H2O
74		18-0972	1	83	-50	1.83	0.20	11	-51	PtS
75		11-0306	1	39	50	5.81	0.20	2	49	CaCuAsO4(OH)
76		25-0298	1	38	-100	6.20	0.20	2	-105	Cu6Hg3As4S12
77		02-1282	1	34	-50	3.97	0.20	1	-55	Li3Na3Al2F12
78		15-0682	1	39	0	3.96	0.20	0	16	PbSeO4
79		39-0368	1	40	-50	5.12	0.19	3	-40	Ca3Al2(SiO4)3
80		16-0155	1	35	-100	4.31	0.19	2	-88	Ca3ZrSi2O9
81		43-1468	1	42	75	2.25	0.19	1	80	CuCr2S4
82		33-0332	1	34	100	8.72	0.19	1	105	CeNbO4
83		42-0590	1	38	-25	20.94	0.19	0	-48	CaV2O6·4H2O
84		35-0672	1	42	100	3.47	0.18	1	119	(Cs,Na)SbTa4O12
85		33-0895	1	34	100	9.12	0.18	4	95	MnH(AsO4)·H2O
86		41-0582	1	40	50	6.12	0.18	1	38	Na4Ca6Mn19Li2Al8(P
87		20-0240	1	39	0	9.72	0.18	0	12	Ca5(SiO4)2(OH)2
88		32-0680	1	40	25	6.11	0.18	0	15	NdNbO4
89		16-0609	1	58	-75	4.47	0.18	1	-65	NaAl8V10O38·30H2O
90		22-0990	1	38	0	5.89	0.18	0	27	Y2FeBe2Si2O10
91		39-0195	1	39	-50	11.37	0.18	3	-16	Ca3Zr(Si2O9)O2
92		29-1338	1	37	25	5.97	0.18	0	26	CuTlHgAsS4
93		22-0539	1	38	-75	5.37	0.17	1	-41	Ca3Si2O7
94		35-0640	1	37	-50	10.08	0.17	1	-16	Y2(SiO4)(CO3)
95		12-0688	1	46	-100	4.93	0.17	9	-84	ZnS
96		39-1363	1	46	-100	4.93	0.17	9	-84	ZnS
97		14-0189	1	39	0	3.15	0.17	2	16	(Sn,Pb,Zn)S
98		18-0791	1	39	75	3.61	0.16	0	94	(Mn,Mg)(Mn,Fe)2O4

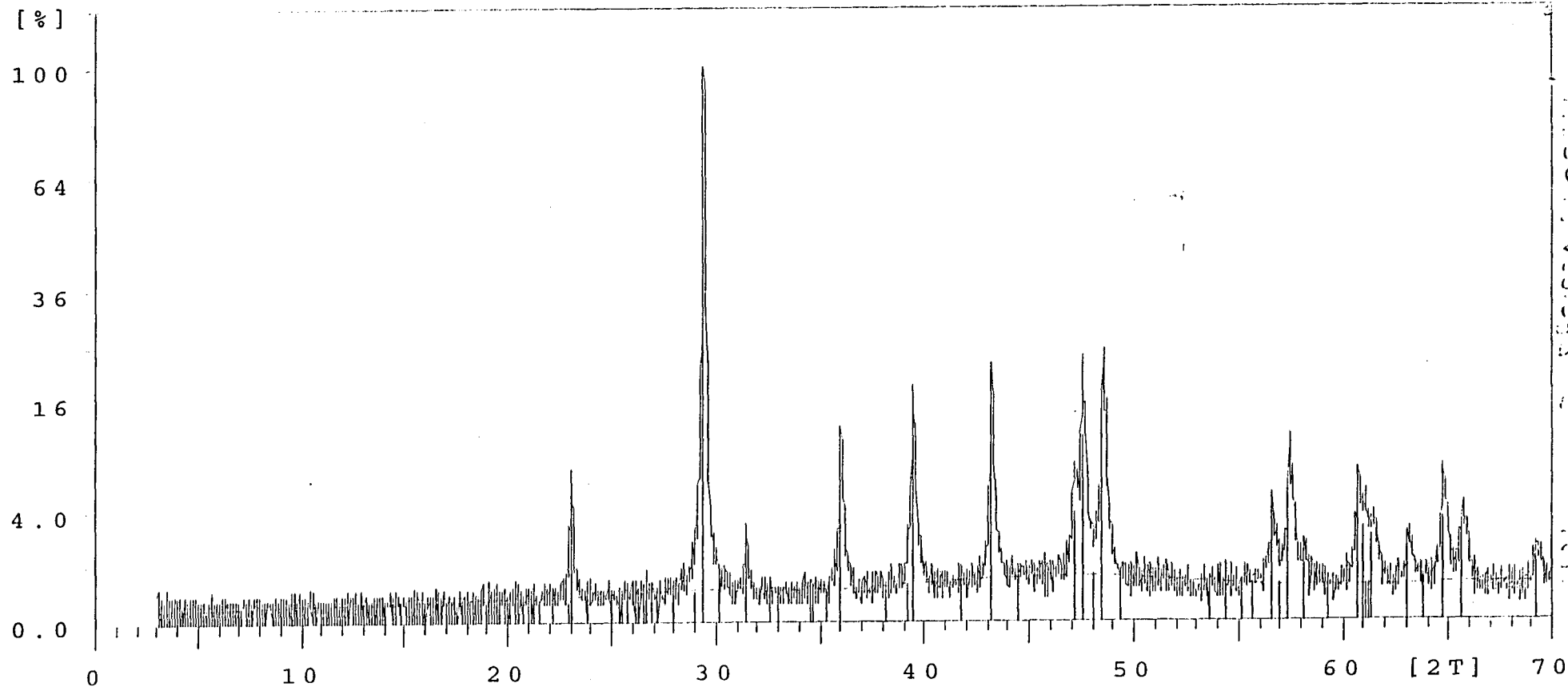
SV15R  
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Natrobistantite
Grossular, hydroxylan
Metacinnabar, selenian
Dwornikite, syn
Davidite
Carlinite, syn
Vigezzite
Talnakhite, syn
Tomichite
Nagyagite, syn
Griphite
Altaite
Manganneptunite
Cobaltite
Aragonite, syn
Giraudite
Tetrahedrite, mercurian, syn
Strontianite, syn
Nesquehonite, syn
Cooperite, syn
Conichalcite
Aktashite
Cryolithionite
Kerstenite, syn [NR]
Grossular
Baghdadite, syn
Florensovite
Fergusonite- $\beta$ -(Ce), syn
Rossite
Cesttibtantite
Krautite, syn
Griphite
Calcio-chondrodite, syn
Fergusonite- $\beta$ -(Nd)
Vanalite
Gadolinite-(Y), heated
Baghdadite, syn
Routhierite
Rankinite
Iimoriite-(Y)
Wurtzite-10\ITH\RG
Wurtzite-8\ITH\RG
Teallite, zincian
Hausmannite

SV15R

Sample ident.: Sv15R

14-feb-1997 15:00



DB03

Calcite,  
Richteri

Sv15R

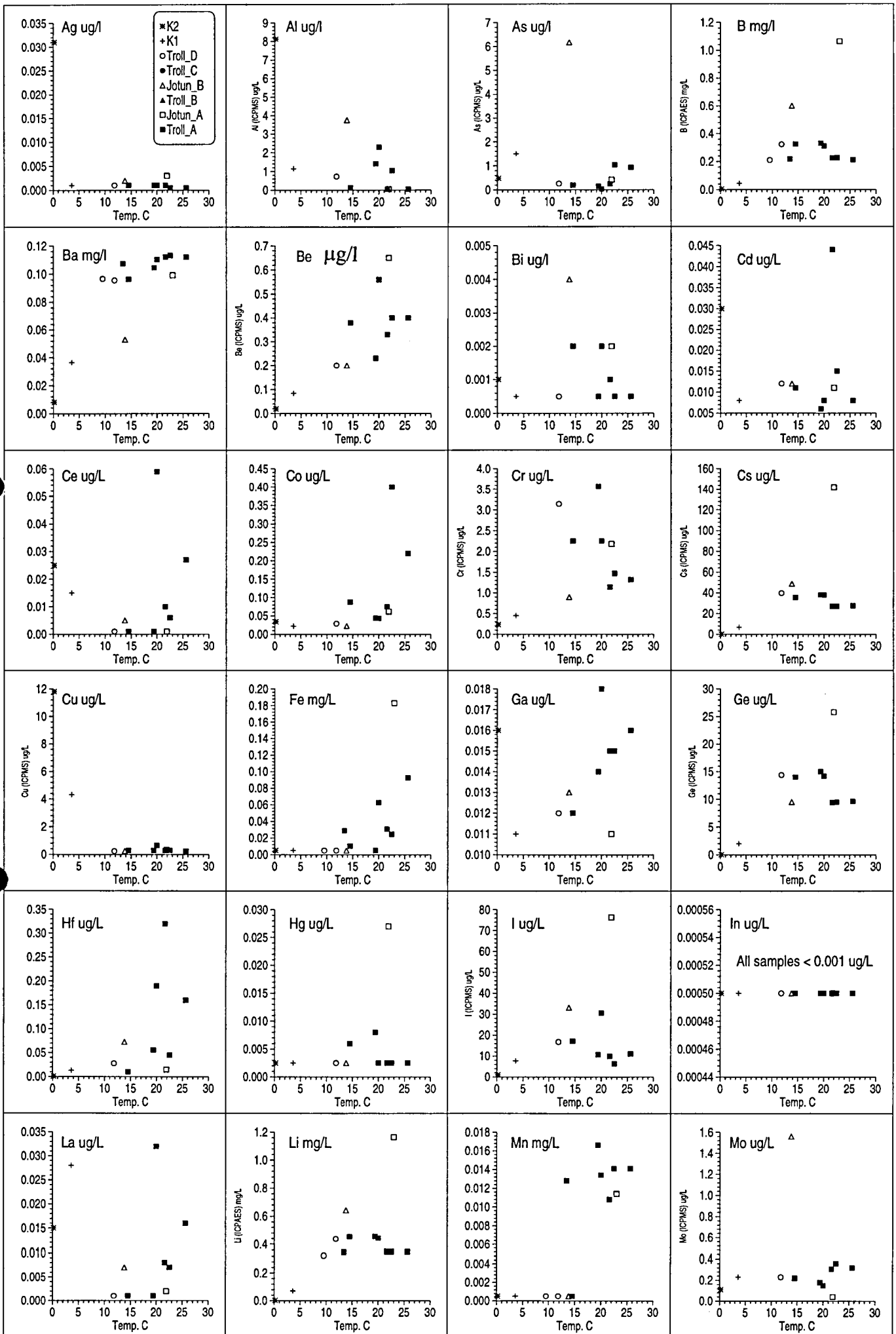
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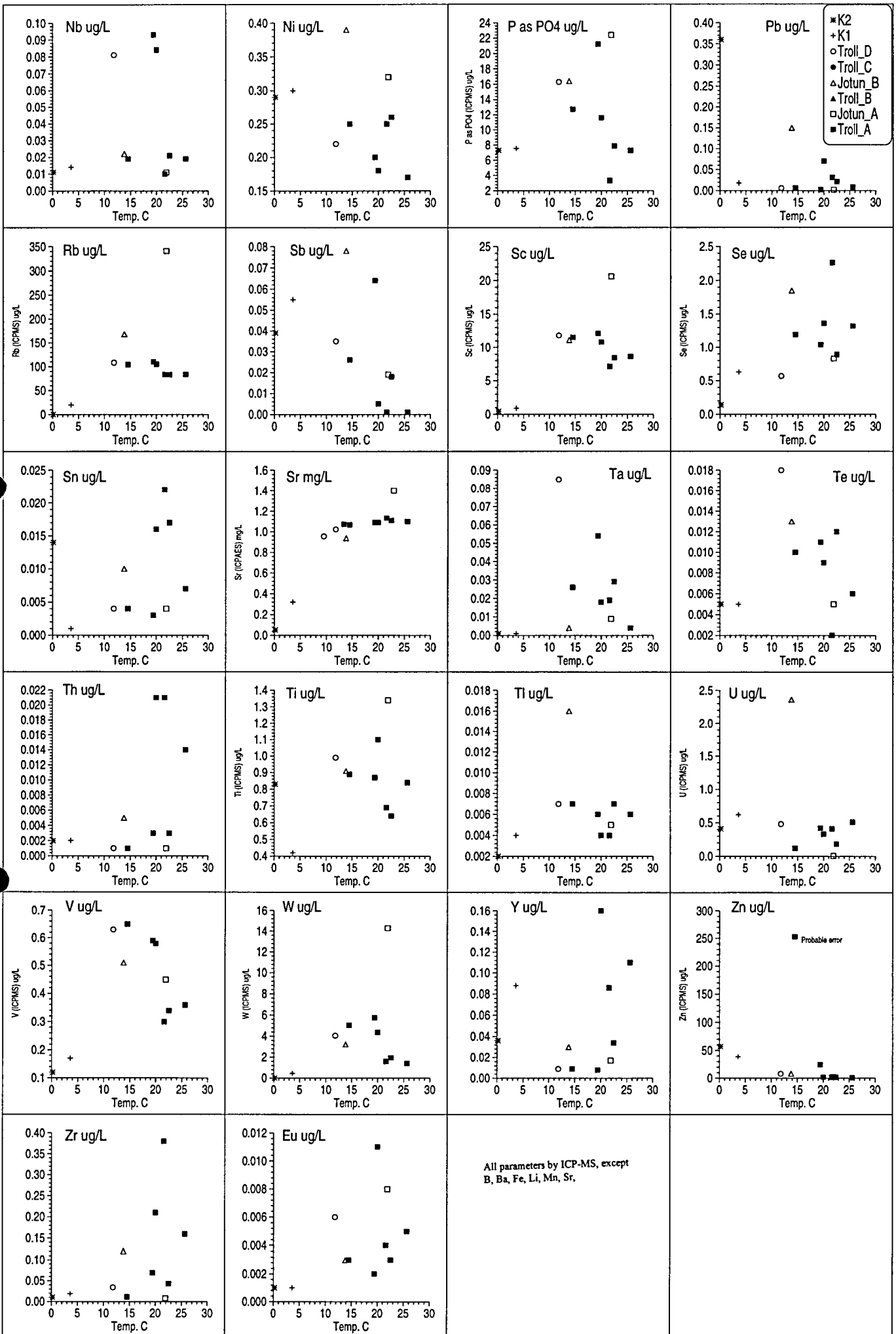
## **Appendix 7**

### **Boxplots / x-y plots illustrating distribution of analytical parameters versus temperature and as a function of spring type**

Samples below analytical detection limit are plotted at half the detection limit.



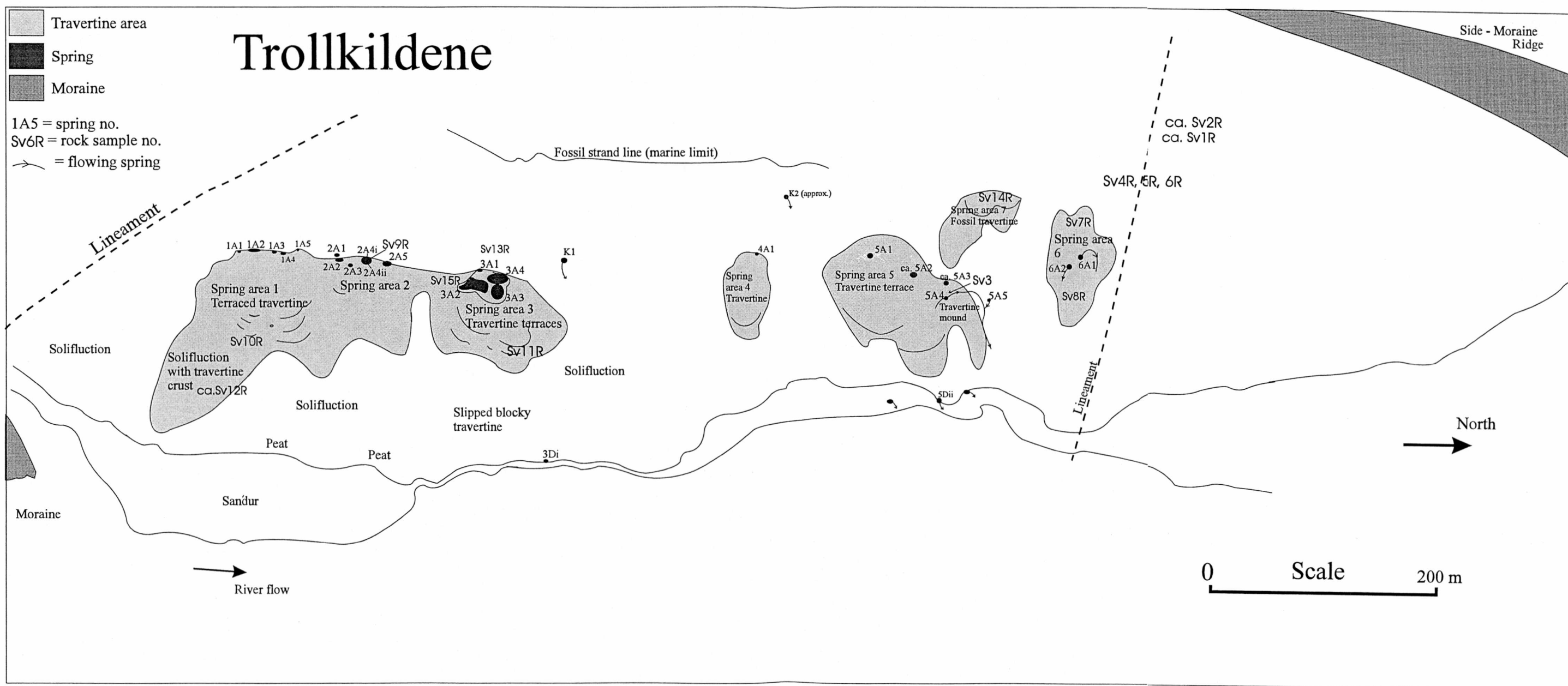
In cases where the parameter is below the analytical detection limit, a value of half the detection limit is plotted.



## **Appendix 8**

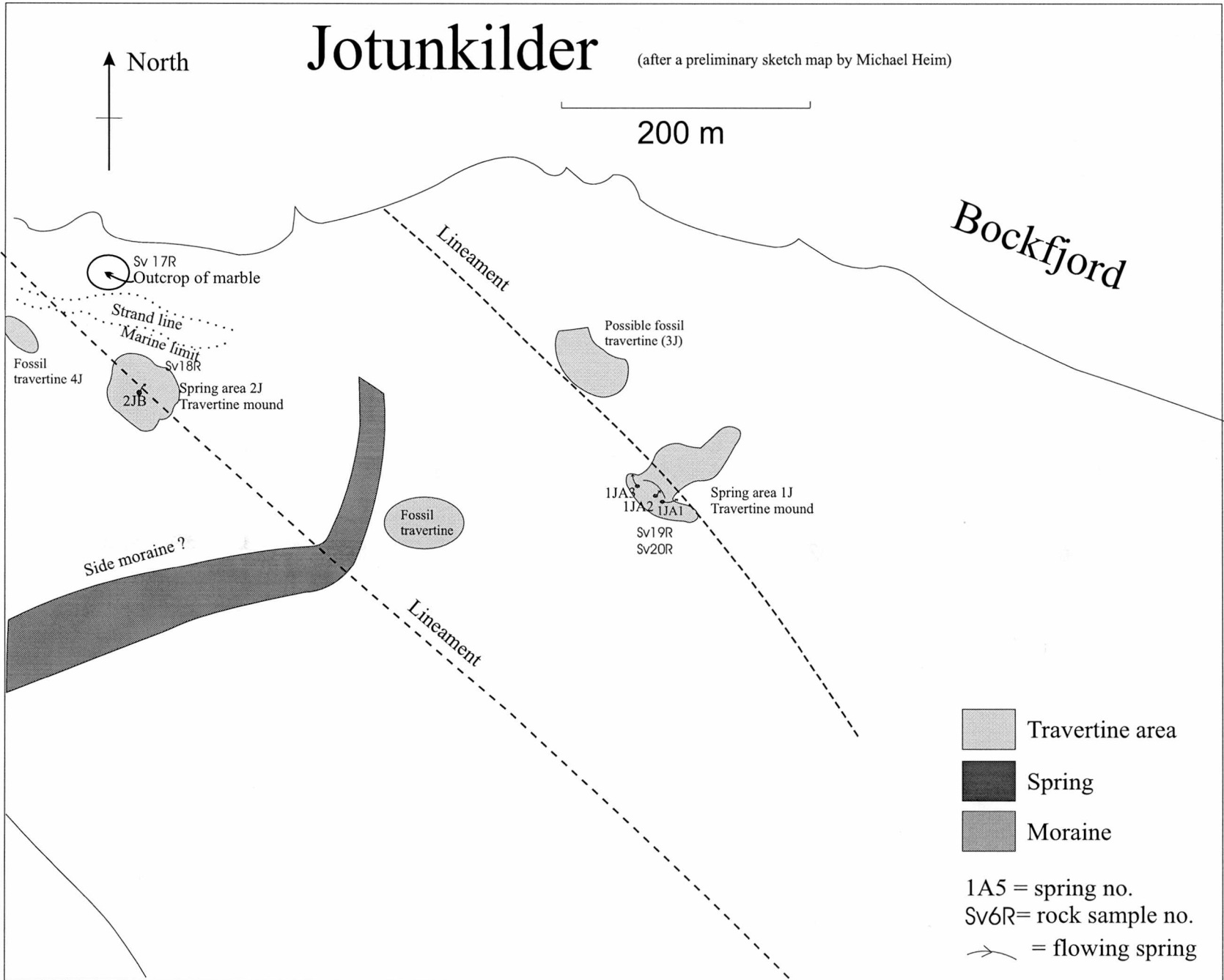
### **Sketch maps illustrating spring location**

Simplified versions of preliminary maps drawn by Michael Heim on the basis of field observations and interpretation of aerial photographs.



# Jotunkilder

(after a preliminary sketch map by Michael Heim)



## Appendix 9

### Results of radon analyses from Statens Strålevern (Prof. Terje Strand) in Bq/l

Sample numbers:	971-96	Spring 5A3	11:30 am 13/7/96
	972-96	Spring 5A3	11:33 am 13/7/96
	974-96	Spring K1	11:53 am 13/7/96
	975-96	Spring K1	11:55 am 13/7/96
	976-96	Spring 3A1	12:00 noon 13/7/96
	977-96	Spring 2A4ii	12:09 pm 13/7/96
	979-96	Spring 3A1	12:02 pm 13/7/96

#### Banks David

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**From:** Terje Strand  
**To:** David.Banks  
**Subject:** vannprover - Svalbard  
**Date:** 05. August 1996 09:07

Hei David!

Her er resultatene fra de 8 pr=F8vene fra Svalbard som vi har analysert:

971-96: 7 kBq/m<sup>3</sup>  
 972-96: 7 kBq/m<sup>3</sup>  
 974-96: 65 kBq/m<sup>3</sup>  
 975-96: 69 kBq/m<sup>3</sup>  
 977-96: 2 kBq/m<sup>3</sup>  
 976-96: 9 kBq/m<sup>3</sup>  
 979-96: 8 kBq/m<sup>3</sup> =20

Som du ser er resultatene lave.

Med hilsen

Terje Strand